Soil $\delta^{15}N$ is a better indicator of ecosystem nitrogen cycling than plant $\delta^{15}N$: A global meta-analysis

批注 **[f1]:** Response to the comment 1 from RC2.

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Submitted to: ***Soil***

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Abstract. The nitrogen-15 (^{15}N) natural abundance composition (δ^{15}N) in soils or plants is a useful tool to indicate the openness of ecosystem N cycling. This study was aimed to evaluate the influence of the experimental warming on soil and plant δ^{15}N. We applied a global meta-analysis method to synthesize 79 and 76 paired observations of soil and plant δ^{15}N from 20 published studies, respectively. Results showed that the mean effect sizes of the soil and plant δ^{15}N under experimental warming were -0.524 (95% CI: -0.987 to -0.162) and 0.189 (95% CI: -0.210 to 0.569), respectively. This indicated that soil δ^{15}N had negative response to warming at the global scale, where warming had no significant effect on plant δ^{15}N. Experimental warming significantly ($p < 0.05$) decreased soil δ^{15}N in Alkali and medium-textured soils, in grassland/meadow, under air warming, for 4-10 yr warming period and for an increase of > 3 ^{o}C in temperature, whereas it significantly ($p < 0.05$) increased soil δ^{15}N in neutral and fine-textured soils and for an increase of 1.5-3 ^{o}C in temperature. Plant δ^{15}N significantly ($p < 0.05$) increased with increasing temperature in neutral and fine-textured soils and significantly ($p < 0.05$) decreased in alkali soil. Latitude did not affect the warming effects on both soil and plant δ^{15}N. However, the warming effect on soil δ^{15}N was positively controlled by the mean annual temperature, which is related to the fact that the higher temperature can strengthen the activity of soil microbes. The effect of warming on plant δ^{15}N had weaker relationships with environmental variables compared with that on soil δ^{15}N. This implied that soil δ^{15}N was more effective than plant δ^{15}N in indicating the openness of global ecosystem N cycling.

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批注 **[f3]:** Response to the comment 6 from RC1.

批注 **[f4]:** Response to the comment 7 from RC1.

批注 **[f5]:** Response to the comment 1 from RC2.

批注 **[f6]:** Response to the comment 8 from RC1.

1 Introduction

Nitrogen (N) is one of the most important nutrient elements for plant growth and the key limiting factors for vegetation productivity (McLay et al., 2001; Zhu et al., 2018; Lu et al., 2020). On the one hand, if the available N in the soil is insufficient, it will damage and weaken the ecosystem service function, including the supply of primary material products, water conservation, climate regulation, etc. (Averill and Waring, 2018). On the other hand, if the available N in the soil is over supplied, it will also damage the structure and function of the ecosystem, resulting in a series of environmental problems such as soil acidification and imbalance of ecosystem nutrient (Schrijver et al., 2008). The intermediate products of the N cycling processes, such as nitrate nitrogen ($NO_3^- - N$), nitrous oxide (N_2O) and nitric oxide (NO), may also cause eco-environmental pollution such as eutrophication of water body and aggravation of climate-related issues (Liao et al., 2019). Therefore, it is of great significance to reveal the openness of the ecosystem N cycle process for understanding the plant N fixation and long-term trend of N cycling and protecting the eco-environment (Wang et al., 2014; Wu et al., 2019). Openness is a measure of both N inputs and outputs relative to internal cycling and determines both the potential rate of N accumulation in the ecosystem and the potential for N losses following a disturbance (Rastetter et al., 2021).

The ^{15}N natural abundance composition ($\delta^{15}N$) in soils or plants (leaves, shoots, fine roots and litter) is often used to indicate the openness of ecosystem N cycling (Robinson, 2001). This is because the lighter isotope of ^{14}N always preferentially

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批注 [f11]: Response to the comment 2 from RC2.

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loses from the ecosystem. Thus, the isotopic fractionation effect results in gradual ^{15}N enrichment in the ecosystem (Aranibar et al., 2004). The larger the $\delta^{15}N$ value, the higher degree of openness of N cycling. In addition, soil $\delta^{15}N$ also appears to reflect the degree of decomposition of the organic matter, showing that $\delta^{15}N$ increases with processing (Craine et al., 2015). A large number of studies have confirmed that climate was the main factor regulating the soil and plant $\delta^{15}N$ (Craine et al., 2015; Soper et al., 2015). Previous studies have demonstrated that precipitation had a negative effect on soil and plant $\delta^{15}N$ from in-situ evidences to cross-sites syntheses (Swap et al., 2004; Soper et al., 2015). However, the influence of temperature on soil and plant $\delta^{15}N$ remained controversial. Some studies have showed that soil and plant $\delta^{15}N$ increased with temperature (Amundson et al., 2003; Craine et al., 2015), while others have indicated that $\delta^{15}N$ decreased with temperature (Cheng et al., 2009; Sheng et al., 2014) or even there was no correlation between them (Yang et al., 2013). The various studies suggested that the responses of soil and plant $\delta^{15}N$ to warming were very complex and not well understood. In addition to climate factor, soil and plant $\delta^{15}N$ are affected by a variety of other environmental factors, such as vegetation type, topography, soil properties and management practices (Gurmesa et al., 2017; Wang et al., 2019). However, we know little about the influences of environmental factors on the warming effect on ecosystem N cycling, in terms of soil and plant $\delta^{15}N$.

Soil and air warming experiments have often been conducted to study the effect of warming on the ecosystem N cycling at site scale (Schindlbacher et al., 2009). At present, the effect of experimental warming on soil and plant $\delta^{15}N$ has not been

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studied on a global scale. The objectives of this study were to: (i) detect the effect of experimental warming on the soil and plant $\delta^{15}N$ based on a global meta-analysis of 20 studies; and (ii) identify the main factors influencing the warming effect on the soil and plant $\delta^{15}N$. Specifically, we hypothesized that soil $\delta^{15}N$ is a better indicator of ecosystem N cycling than plant $\delta^{15}N$.

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2 Materials and methods

2.1 Source of data and selection criteria

Peer-reviewed journal articles and dissertations related to soil and plant $\delta^{15}N$ under experimental warming were searched using Web of Science and China National Knowledge Infrastructure (CNKI, http://www.cnki.net) until March 31, 2020 (Tab. 1). The keywords used for the literature search were related to: "nitrogen isotope composition", "experimental warming" and "ecosystems nitrogen cycling".

Our criteria were as follows: at least one of the target variables was contained, including soils (different fractions, e.g., sand, silt, clay, aggregate and bulk soil) and plants (leaves, shoots, roots and litters) $\delta^{15}N$; studies with climate gradients (space-time substitution) were excluded and only field warming experimental studies were included; only data from control and warming treatments were applied for multifactor experiments; means, standard deviations (SD) (or standard errors (SE)) and sample sizes were directly provided or could be calculated from the studies; if one article contained soil or plant $\delta^{15}N$ in multiple years, only the latest results were applied since the observations should be independent in the meta-analysis (Hedges et al., 1999).

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2.2 Data extraction and statistical analysis

In total, 20 published papers were selected from 54 published papers. The locations of warming experiments were presented and their site information is listed in Tab. 1. For each study, the means, the statistical variation (SE or SD) and the sample size values for treatment and control groups were extracted for each response variable ($\delta^{15}N$). In addition to $\delta^{15}N$, the latitude, longitude, altitude, soil pH, organic matter content, vegetation type, mean annual precipitation (MAP) and mean annual temperature (MAT) were also extracted if they were provided (Tab. 1). All data were extracted from tables or digitized from graphs with the software GetData v2.2.4 (http://www.getdata-graph-digitizer.com). A total of 79 and 76 paired observations for soil and plant $\delta^{15}N$ were obtained, respectively.

The METAWIN 2.1 software (Sinauer Associates Inc., Sunderland, MA, USA) (Rosenberg et al., 2000) was used to perform meta-analysis in this study. The Hedges' *d* value was used as the effect size (Hedges et al., 1999). The absolute *d* value indicated the magnitude of the treatment impact. Positive or negative *d* values represented an increase or decrease effect of the treatment, respectively. Zero meant no difference between treatment and control groups. Resampling tests were incorporated into our meta-analysis using the bootstrap method (999 random replicates). The mean effect size (calculated from 999 iterations) and 95% bootstrap confidence intervals (CI) were then generated. If the 95% CI values of *d* did not overlap zero, the effects of experimental warming on $\delta^{15}N$ were considered significant at $p < 0.05$. We used a random effects model to test whether warming had a significant

批注 **[f22]:** Response to the comment 12 from RC1.

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批注 **[f26]:** Response to the comment 15 from RC1.

effect on δ^{15}N. To examine whether experimental conditions alter the response direction and magnitude of soil and plant δ^{15}N, observations were further divided into subgroups according to the soil acidity-alkalinity (acid (pH < 6.5), neutral (6.5 < pH < 7.5), and alkali (pH > 7.5)), vegetation types (forest/shrub, moss/lichen, and grassland/meadow), warming treatments (soil warming, air warming, and both soil and air warming), soil texture (fine-, medium-, and coarse-textured soil), length of warming (< 4 yr, 4-10 yr, and > 10 yr), and increase in temperature (< 1.5 ^{o}C, 1.5-3 ^{o}C, and > 3 ^{o}C). A random effects model with a grouping variable was used to compare responses among different subgroups. Linear regression analyses were applied to assess the relationships between the Hedges' *d* values and environmental factors (i.e., latitude, altitude, MAT and MAP).

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3 Results

Across all sites, the mean effect sizes of the soil and plant δ^{15}N under experimental warming were -0.524 (95% CI: -0.987 to -0.162) and 0.189 (95% CI: -0.210 to 0.569), respectively (Fig. 1). Experimental warming significantly (p < 0.05) decreased soil δ^{15}N in Alkali (mean effect size = -2.484; 95% CI: -2.931 to -2.060) and medium-textured (mean effect size = -0.676; 95% CI: -1.153 to -0.249) soils, in grassland/meadow (mean effect size = -0.609; 95% CI: -1.076 to -0.190), under air warming (mean effect size = -0.652; 95% CI: -1.081 to -0.273), for 4-10 yr warming period (mean effect size = -0.652; 95% CI: -1.081 to -0.273) and for an increase of > 3 ^{o}C in temperature (mean effect size = -0.652; 95% CI: -1.081 to -0.273). However, it significantly (p < 0.05) increased soil δ^{15}N in neutral (mean effect size = 0.359; 95%

CI: 0.078 to 0.620) and fine-texture soils (mean effect size = 2.394; 95% CI: 1.770 to 3.735), and for an increase of 1.5-3 ^{o}C in temperature (mean effect size = 0.409; 95% CI: 0.070 to 0.707) (Fig. 2). Experimental warming did not significantly ($p > 0.05$) change soil δ^{15}N under other experimental conditions.

In addition, experimental warming significantly ($p < 0.05$) increased plant δ^{15}N in neutral (mean effect size = 3.157; 95% CI: 1.529 to 6.967) and fine-textured soils (mean effect size = 1.202; 95% CI: 1.042 to 1.360), whereas it significantly ($p < 0.05$) decreased plant δ^{15}N in alkali soil (mean effect size = -1.930; 95% CI: -2.325 to -1.573) (Fig. 2). Experimental warming did not significantly ($p > 0.05$) change plant δ^{15}N under other experimental conditions.

For soil and plant δ^{15}N, their responses to experimental warming did not correlate well with latitude ($p = 0.268$ and $p = 0.160$, respectively) (Fig. 3ab). However, the Hedges' *d* values of soil δ^{15}N decreased significantly with altitude ($p < 0.001$) (Fig. 3c) and increased significantly with MAT ($p < 0.001$) and MAP ($p < 0.001$) (Fig. 3eg). In addition, the Hedges' *d* values of plant δ^{15}N were also found to increase significantly with MAP ($p < 0.001$) (Fig. 3h). However, the responses of plant δ^{15}N to experimental warming did not correlate well with altitude ($p = 0.109$) and MAT ($p = 0.002$) (Fig. 3df).

4 Discussion

A significant decreasing trend in soil δ^{15}N and no significant trend in plant δ^{15}N were found in this study. This is somewhat inconsistent with previous findings. Chang et al. (2017) observed that soil and plant δ^{15}N values decreased under warming in the

批注 **[f28]:** Response to the comment 1 from RC2.

Tibetan permafrost. However, Zhang et al. (2019) found that the warming treatment significantly increased soil and plant $\delta^{15}N$ in a subtropical forest. The various studies suggest that soil and plant $\delta^{15}N$ are controlled by interactive effects of N fixation and mineralization. At the global scale, $\delta^{15}N$ of N input (~ 0) is generally lower than that of soil, so greater N fixation or higher N input (deposition and fertilization) under warming can result in a lower soil $\delta^{15}N$ (Sorensen and Michelsen, 2011; Rousk and Michelsen, 2017; Wang et al., 2018).

Soil pH has an important influence on nitrification, denitrification and N_2O emissions from soils (Kyveryga et al., 2004). The results in this study showed that when the soil was alkaline, the mean effect sizes of soil and plant $\delta^{15}N$ under warming were negative, while when the soil was neutral, they were positive (Fig. 2ab). Compared with alkaline condition, the near neutral conditions are more suitable for the biological activities of heterotrophic denitrifying bacteria (Simek and Cooper, 2002). Therefore, the denitrification activity is usually higher under neutral conditions, resulting in an enrichment of soil and plant N pools with ^{15}N (Kyveryga et al., 2004). Vegetation type has limit effects on $\delta^{15}N$ under warming, except for soil $\delta^{15}N$ in grassland/meadow (Fig. 2cd). This may be related to the differences in altitude, MAP and MAT among three vegetation types (Tab. 1). Warming treatment was found to have a substantial effect on soil $\delta^{15}N$, showing that the mean effect size of soil $\delta^{15}N$ under air warming was negative and less than that under soil warming (Fig. 2ef). Salmon et al. (2016) have found that soil warming can increase N availability by stimulating mineralization of organic matter in the warmed active layer. In addition,

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批注 [f33]: Response to the comment 10 from RC2.

批注 [f34]: Response to the comment 1 from RC2.

批注 [f35]: Response to the comment 7 from RC2.

air warming directly impacts aboveground temperatures and has an indirectly effect on soil δ^{15}N (Pardo et al., 2006). From Fig. 2gh, the finer the soil texture, the more significant the positive effect of warming on soil and plant δ^{15}N. The possible reason is that the finer the soil texture, the stronger the adsorption of various ions on the soil and the smaller the leaching loss of the soil, resulting in the greater the residual amount of ^{15}N in the soil (Webster et al., 1986). In addition, the longer warming period and the greater increase in temperature resulted in the more negative effect of warming on soil δ^{15}N (Fig. 2ik). Chang et al. (2017) deduced that N fixation was greater under warming and consequently resulted in a lower soil δ^{15}N.

In the study of Mayor et al. (2015), who found that soil and plant δ^{15}N were significantly ($p < 0.001$) and negatively correlated with latitude at the global scale. However, the Hedges' *d* values of soil and plant δ^{15}N had weak correlations with latitude in this study (Fig. 3). The warming effect on soil δ^{15}N was significantly ($p < 0.001$) influenced by altitude, MAT and MAP. Among these, the strongest correlation was observed for MAT. Temperature has been demonstrated to be a key factor to regulate the soil δ^{15}N by influencing the processes of N mineralization, nitrification and denitrification (Craine et al., 2015). The higher temperature can strengthen the activity of soil microbes and thereafter increase the N uptake for plants and soil N loss from ammonia volatilization and gas N emissions, and thereby more ^{15}N-enriched retains in soils (Wang et al., 2019). Craine et al. (2015) also proposed that warmer sites have soil N that is elevated in ^{15}N, but has lower C:N. Once C:N is controlled, there is little pattern in ^{15}N across temperature gradients. In other words, the

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批注 [f40]: Response to the comment 21 from RC1.

relationship between soil δ^{15}N and climate is indirect, and mediated through climate effects on soil properties (e.g., the concentrations of organic carbon and clay). High *d* values of soil δ^{15}N corresponded to MAT of about 20 ^{o}C, which was the most suitable temperature for nitrification and denitrification. However, warming had a substantial negative impact on soil δ^{15}N when MAT decreased to around -5 ^{o}C. Recently, Rousk et al. (2018) also found that the increase of temperature in the Arctic promoted the biological N fixation, which can decrease the soil δ^{15}N. The decrease of *d* values of soil δ^{15}N with increasing altitude and decreasing MAP in this study might be caused by the positive response of *d* values to MAT.

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批注 [f42]: Response to the comment 4 from RC1.

The relationships between the *d* values and environmental variables for plant δ^{15}N were weaker than those for soil δ^{15}N (Fig. 3). The possible reason is that several other factors (e.g., plant N concentrations and species richness) might co-regulate plant δ^{15}N (Wu et al., 2019). This is consistent with the study of Craine et al. (2009), who found different inflection points in soil and plant δ^{15}N relationships with MAT. In addition, plants are generally depleted in ^{15}N relative to soils. Above results implied that soil δ^{15}N was more efficient in indicating the openness of ecosystem N cycling than plant δ^{15}N at the global scale. Although the present study provided a global meta-analysis of the responses of δ^{15}N to experimental warming, the magnitude of these responses might be uncertain. For example, a small number of observations were obtained in moss/lichen under soil warming and both soil and air warming treatments, which would affect the results of meta-analysis. Future research should take more experimental data into account in order to better investigate the warming

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effects on δ^{15}N.

6 Conclusions

Our global meta-analysis indicated a significant decreasing trend in soil δ^{15}N and no significant trend in plant δ^{15}N under experimental warming. Latitude did not affect the warming effects on δ^{15}N. However, the warming effect on δ^{15}N was related to soil acidity-alkalinity, texture, vegetation type, warming treatment and period, increase in temperature, altitude, MAT and MAP. The effect of warming on soil δ^{15}N was better correlated with environmental variables compared with that on plant δ^{15}N. Our findings should be useful for understanding the underlying mechanisms of the response of ecosystem N cycling to global warming.

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Data availability. The data that support the findings of this study are available from the corresponding author upon request.

Author contributions. KL and QZ designed this study, KL and XL performed the meta-analysis, KL and QZ obtained funding, and KL and XL wrote the paper with contributions from QZ.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. We thank two anonymous reviewers and editor for their efforts on this paper. Support for this research was provided by the National Natural Science Foundation of China and by Chinese Academy of Sciences.

Financial support. This study was financially supported by the National Natural Science Foundation of China (41771107 and 42171077), the Key Research Program of Frontier Sciences, Chinese Academy of Sciences (QYZDB-SSW-DQC038), and the Youth Innovation Promotion Association, Chinese Academy of Sciences (2020317).

Review statement. This paper was reviewed by two anonymous referees.

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Table 1: Site characteristics from a global meta-analysis of 20 studies.

References	Country/Region	Vegetation types	Soil pH	Soil type	OMC[a] (%)	Latitude	Longitude	Altitude (m a.s.l)	MAT[b] (°C)	MAP[c] (r
Anadon-Rosell et al. (2017)	Spain	Subalpine shrub	4.49~4.63	Mineral soil	13.15~14.04	41.39 °N	2.17 °E	2250	3	1146.4
Zhang et al. (2019)	China	*C. lanceolata* seedlings	5.07	Oxisol	-	26.32 °N	117.6 °E	300	19.1	1670
Lim et al. (2019)	Sweden	Boreal forests	5.92~6.44	A thin, podzolic, sandy soil	-	64.12 °N	19.45 °E	310	2.4	600
Deane-Coe et al. (2015)	USA	Tundra mosses	-	Gelisol	-	63.88 °N	149.23 °W	700	-2.7~-1	138~228
Bijoor et al. (2008)	USA	Turfgrass lawn	-	Alkaline alo clay	-	33.7 °N	117.7 °W	30	18.6	352
Chang et al. (2017)	China	Alpine meadow	9.1~9.3	Gelisols	5.5	34.73 °N	92.89 °E	4750	-5.3	269.7
Gonzalez-Meler et al. (2017)	Brazil	Grasslands	5.0	Dystrophic red latosols	-	21.17 °S	47.86 °W	578	21.5	1100
Natali et al. (2012)	USA	Shrubs, sedges and mosses	-	Gelisol	-	63.88 °N	149.23 °W	700	-1	178~250
Munir et al. (2017)	Canada	Shrubs, mosses and trees	-	-	-	55.27 °N	112.47 °W			
Salmon et al. (2016)	USA	*Eriophorum vaginatum*	-	Gelisols	-	63.88 °N	149.23 °W	700	-1.45	200
Rui et al. (2011)	China	Alpine	-	-	-	37.62 °N	101.2 °E	3200	-2	500

批注 **[f46]:** Response to the comment 1 from RC1 and comment 9 from RC2.

		meadow								
Aerts et al. (2009)	Sweden	Shrubs, mosses and trees	-	-	-	68.35 °N	18.82 °E	340	0.5	303
Cheng et al. (2011)	USA	Tallgrass prairie	Neutral pH	Nash-Lucien complex	-	34.98 °N	97.52 °W		16	911.4
Dawes et al. (2017)	Switzerland	Alpine treeline	-	Sandy Ranker and Podzols	-	46.77 °N	9.87 °E	2180	9.2	444
Schaeffer et al. (2013)	Greenland	Prostrate dwarf-shrub herb tundra	-	Turbic cryosols	-	76 °N	68 °W		4~8	<200
Schnecker et al. (2016)	Austria	Spruce forest	Near neutral pH	A mosaic of shallow Chromic Cambisols and Rendzic Leptosols	8.55~14.96	47.58 °N	11.64 °E	910	6.9	1506
Hudson et al. (2011)	Canada	Heath, willow and meadow	-	-	-	78.88 °N	75.78 °W		8.6~10.4	
Lv et al. (2018)	China	*Cunninghamia lanceolata* juveniles	-	Red soil	2.21	26.32 °N	118 °E		19.1	1585
Zhao et al. (2016)	China	Alpine meadow	-	Alpine meadow soils	-	37.48 °N	101.2 °E	3200~3250	-1.7	600
Peng (2017)	China	Alpine meadow	-	Alpine meadow soils	-	34.73 °N	92.89 °E	3200~4800	-5.03	267.4~426.3

[a]Soil organic matter content; [b]Mean annual temperature; [c]Mean annual precipitation. If the soil organic carbon content was provided in the literature, soil

organic matter content was determined by multiplying the organic carbon content by a coefficient of 1.724.

List of Figures:

Figure 1: Effect sizes of the experimental warming on soil and plant $\delta^{15}N$ from a global meta-analysis of 20 studies. The error bars indicate effect sizes and 95% bootstrap confidence intervals (CI). The warming effect was statistically significant if the 95% CI did not bracket zero. The sample size for each variable is shown next to the bar.

Figure 2: Factors influencing the effect sizes of the soil and plant $\delta^{15}N$ under experimental warming from a global meta-analysis of 20 studies, including (a-b) soil acidity-alkalinity, (c-d) vegetation types, (e-f) warming treatments, (g-h) soil texture, (i-j) length of warming and (k-l) increase in temperature. The error bars indicate effect sizes and 95% bootstrap confidence intervals (CI). The warming effect was statistically significant if the 95% CI did not bracket zero. The sample size for each variable is shown next to the bar.

Figure 3: Relationships between the Hedges' *d* values of soil and plant $\delta^{15}N$ with the latitude, altitude, mean annual temperature (MAT) and mean annual precipitation (MAP) under experimental warming.

Figure 1

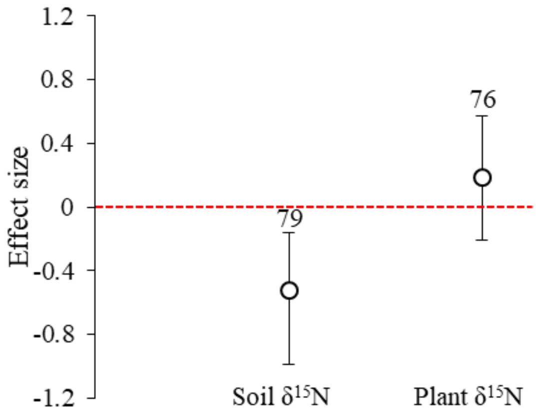
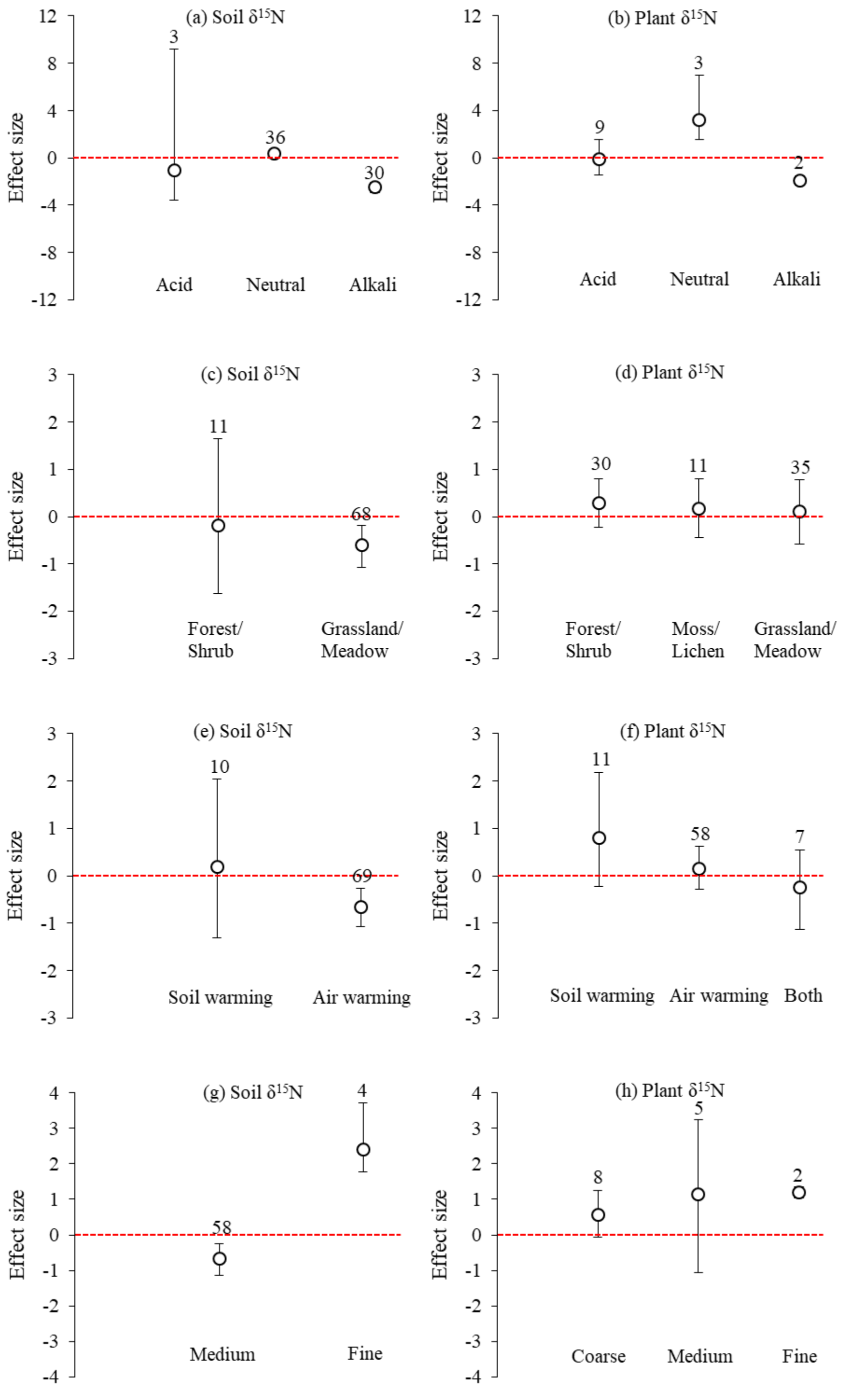


Figure 2


(a) Soil δ15N
Effect size
12
8
4
0
-4
-8
-12
3
36
30
Acid
Neutral
Alkali
(b) Plant δ15N
Effect size
12
8
4
0
-4
-8
-12
9
3
2
Acid
Neutral
Alkali
(c) Soil δ15N
Effect size
3
2
1
0
-1
-2
-3
11
68
Forest/
Shrub
Grassland/
Meadow
(d) Plant δ15N
Effect size
3
2
1
0
-1
-2
-3
30
11
35
Forest/
Shrub
Moss/
Lichen
Grassland/
Meadow
(e) Soil δ15N
Effect size
3
2
1
0
-1
-2
-3
10
69
Soil warming
Air warming
(f) Plant δ15N
Effect size
3
2
1
0
-1
-2
-3
11
58
7
Soil warming
Air warming
Both
(g) Soil δ15N
Effect size
4
3
2
1
0
-1
-2
-3
-4
58
4
Medium
Fine
(h) Plant δ15N
Effect size
4
3
2
1
0
-1
-2
-3
-4
8
5
2
Coarse
Medium
Fine

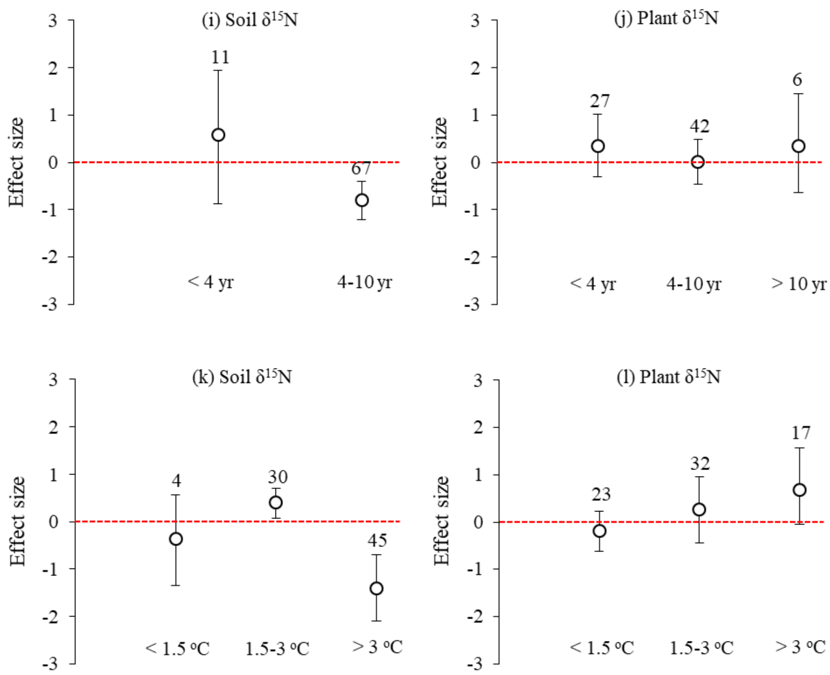
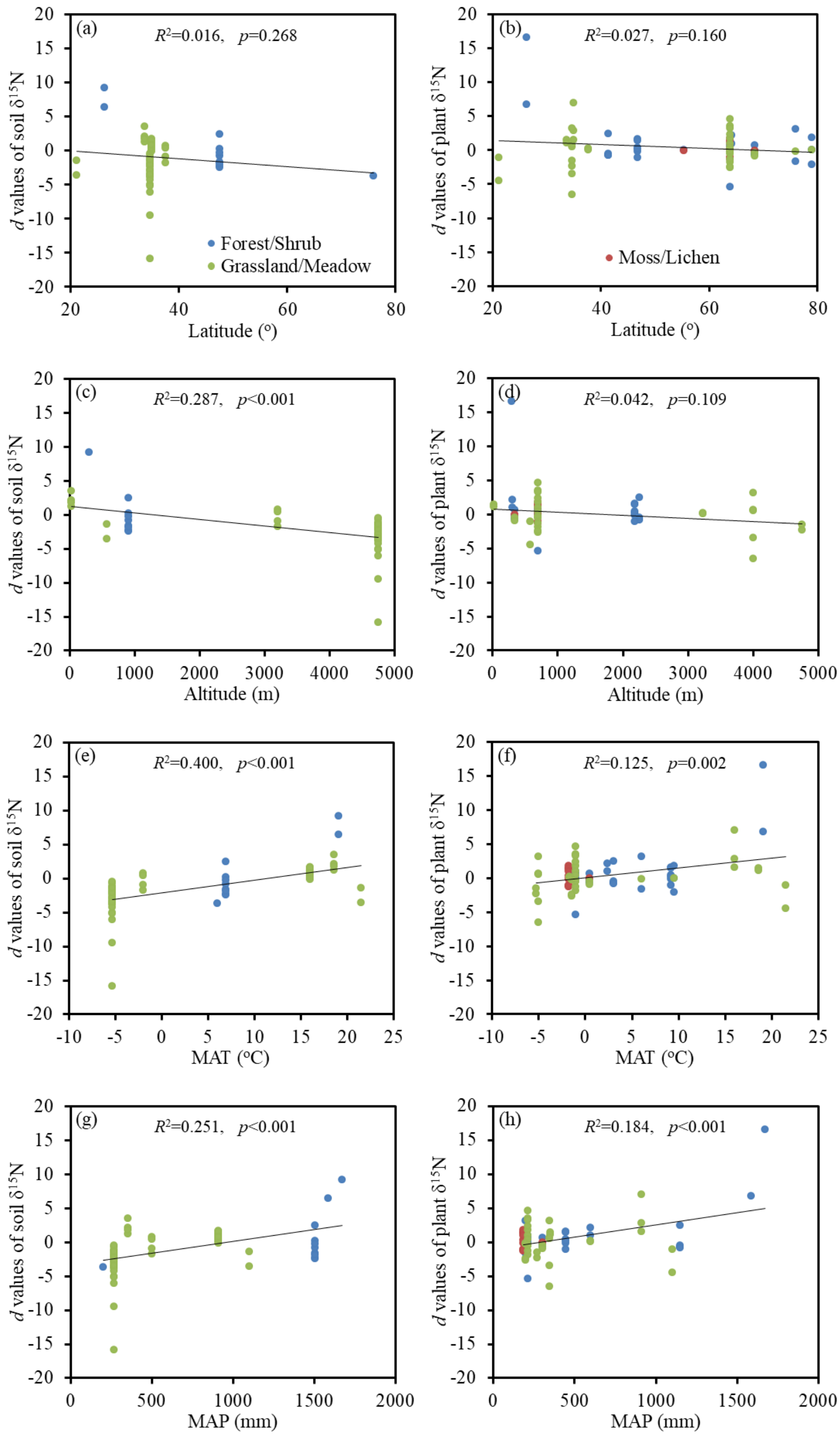

(i) Soil δ15N
Effect size
11
67
< 4 yr
4-10 yr
(j) Plant δ15N
27
42
6
< 4 yr
4-10 yr
> 10 yr
(k) Soil δ15N
4
30
45
< 1.5 ºC
1.5-3 ºC
> 3 ºC
(l) Plant δ15N
23
32
17
< 1.5 ºC
1.5-3 ºC
> 3 ºC

Figure 3



(a)
R^2=0.016, p=0.268
d values of soil δ15N
Latitude (°)
Forest/Shrub
Grassland/Meadow
(b)
R^2=0.027, p=0.160
d values of plant δ15N
Latitude (°)
Moss/Lichen
(c)
R^2=0.287, p<0.001
d values of soil δ15N
Altitude (m)
(d)
R^2=0.042, p=0.109
d values of plant δ15N
Altitude (m)
(e)
R^2=0.400, p<0.001
d values of soil δ15N
MAT (°C)
(f)
R^2=0.125, p=0.002
d values of plant δ15N
MAT (°C)
(g)
R^2=0.251, p<0.001
d values of soil δ15N
MAP (mm)
(h)
R^2=0.184, p<0.001
d values of plant δ15N
MAP (mm)