

Author's response to editors decision (in blue)

“Topical Editor Decision: Publish subject to minor revisions

Comments raised by the referee, public and AE have all been properly addressed. The in depth reply on some of the raised points is much appreciated. The most crucial issues: our doubts on the relation between NO₃ and kiso, and effect of NH₄⁺ on kiso have been resolved. Also foreseen improvements of the readability of the text are expected to be sufficient. We welcome submission of a final revised MS that implements the suggested changes by the authors.”

Following the editors decision of we have implemented the changes outlined in our intial responses to the online discussion. We have also edited the text and re-written the introduction following the good advice of both the editor and the reviewer to improve the readability. We would like to thank Steven Sleutel, Gerbrand Koren and Reviewer 1 for their time and input.

Author's response to review comments

We have listed our responses to the reviewer comments (reproduced in blue) along with the final changes made below. In some cases the exact content of our responses to comments about writing style or clarity now differ to those made during the dicussion phase because we edited and re-written some sections.

EC1

1) SOIL-2020-44 presents on an interesting research into how soil factors would affect O₁₈ exchange between CO₂ and H₂O. As the 18O levels of emitted CO₂ are used for estimating land-atmosphere carbon exchange it presents a well-delineated and timely topic. The experimental assessment of exchange of 18O between soil water and CO₂ is by no means trivial, yet as many as 44 soils collected from a wide geographical area (Eurasia and Australia) were assessed for their potential to oxygen isotope exchange.

We would sincerely like to thank the editor for taking the time to handle the review process of this manuscript and for providing their own useful comments that have improved the manuscript. We have addressed the points raised (reproduced in blue) below.

2) The description was mostly excellent but I concur with the referee that on instances the text is lengthy. For instance the description of the statistical approach and reporting of fitted general linear models are longer than usual. However, here such an elaborate formulation appeared warranted given the clear collinearity of some of the predictor variables.

We agree that as the interpretation of the data is dependent on the statistical tests used it is important to be explicit about the steps taken.

3) Still, in particular the introduction can surely be further condensed. Also the lengthy description of the setup to administer air with CO₂ with contrasting δ¹⁸O requires shortening (e.g. by just referring to or some part of it should be moved to supplementary material).

We have rewritten the introduction and edited the text elsewhere to remove unnesserary information.

4) The description of preparation of soil mesocosms could also be condensed.

Please see the previous comment.

5) Please also use a different term for available NO₃⁻ and available NH₄⁺ (or‘availability of’). Both were measured in typical 1M KCl extracts and are best referred to exchangeable NO₃⁻ and NH₄⁺.

Thanks for this suggestion, we have modified the text accordingly to use exchangeable in place of available.

6) The entire text is filled with long complex sentences,which on their own are well crafted but do not always allow fluid reading. I would recommend the authors to read their manuscript once more and if they see fit do try to subdivide some of the longer sentences.

We have worked to improve the writing style throughout the abstract and main text (see also comment 3 and 4).

7) I concur with the referee that some of the hypotheses require rephrasing and attention needs to be given to his/her possibly justified concern on the NH₄NO₃ administration's ineptitude to robustly prove that NO₃-led to inhibition of carbonic anhydrase.

Please see responses to comments 1) and 4) of Reviewer 1. In brief we have rephrased the hypotheses as suggested and provided caveats about the power of the treatment experiment.

8) The interpretation of the relationship between NH_4^+ and k_{iso} : On L408 following statement is made 'Notably the weak relationship between changes in k_{iso} and NH_4^+ availability identified in this experiment (Table 2) suggests the relationship between these variables across the untreated soils (Table 1) does indeed reflect the pH sensitivity of ammonia speciation rather than a direct causal link.'. I would agree that the link is indeed direct, but strongly doubt that $\text{NH}_4^+/\text{NH}_3$ speciation is important here. Below pH 7 there is virtually no (toxic) NH_3 . Most soils had acidic pH so this explanation seems wrong. Instead: high NH_4^+ levels in upland soils rather suggest nitrification is impeded in some way. There is an obvious link between pH and NH_4^+ levels: at low pH, nitrification is well known to be slowed: indeed there was a strong negative correlation between both (Table 1) and so the negative correlation between NH_4^+ and k_{iso} may just be indirect through their mutual relation with pH. Alternatively, higher NH_4^+ levels point at impeded autotrophic nitrification – a rather energetically unfavourable process and therefore sensitive to environmental constraints. Inhibited nitrification may also point at unfavourable conditions for other microbial processes: perhaps also production of anhydase activity? In Table 2 the artificially elevated NH_4^+ levels are no longer the resultant from slow or fast pH-dependent nitrification and therefore do not display any relation with k_{iso} . The referee also commented on this matter – please do take that comment into account.

Following this good advice we have removed the unfounded reference to the role of ammonia speciation in explaining the apparent co-correlation between k_{iso} , pH and exchangeable ammonium identified in the Spearman's rank analysis presented in Table 1. We maintain that the relationship between k_{iso} and exchangeable NH_4^+ is an artefact of its relationship with soil pH for two reasons. Firstly, the Spearman's rank correlation between soil pH and exchangeable ammonium is strongly influenced by the co-occurrence of low exchangeable ammonium in the high pH soils that exhibit greater k_{iso} (Figure 2 a) & e); groups Csa, Bsh and Cfa). Secondly, we do not find support for a role of exchangeable ammonium in explaining variability in k_{iso} in the subsequent analyses (Tables S1 & S3). Please also see the response to Reviewer 1, comment 1).

9) It is unfortunate that the authors opted for an enormous dose of NH_4NO_3 , viz 0.7 mg NH_4NO_3 g⁻¹ (L168). No doubt such a large addition of NH_4^+ would have led to serious soil acidification following its partial nitrification during the 1 week pre-incubation at 23°C. And on the other hand the obtained NH_4^+ and NO_3^- levels in treated soil would not reflect environmentally realistic exchangeable N levels. That would severely limit the relevance of Fig. 5. Or is the 0.7 mg NH_4NO_3 g⁻¹ a typing error? In any case in their rebuttal the authors will need to present (not necessarily in the revised manuscript) the absolute increase in NH_4^+ and NO_3^- levels alongside changes in pH brought forth by the NH_4NO_3 administration as based on the factorial changes in NO_3^- – now presented it is impossible to judge whether or not excessive amounts of N had been added. This firstly needs to be clarified.

The fertilisation rate of 0.7 mg NH_4NO_3 per gram of dry soil (0.25 mg of N per gram of dry soil) was adopted from Ramirez, Craine and Fierer (2012). The justification for this value (also used in Kaisermann et al., 2018) was that it approximates typically applied fertilizer loads in field studies. We have added a figure to the supplementary material (Figure S3) to provide information about the absolute values of measured parameters in both the untreated controls and treated soils and allow comparison with the wider dataset in Figure 2. The median increase in exchangeable nitrate and ammonium in the treated over untreated soils was 22 and 10 times, respectively.

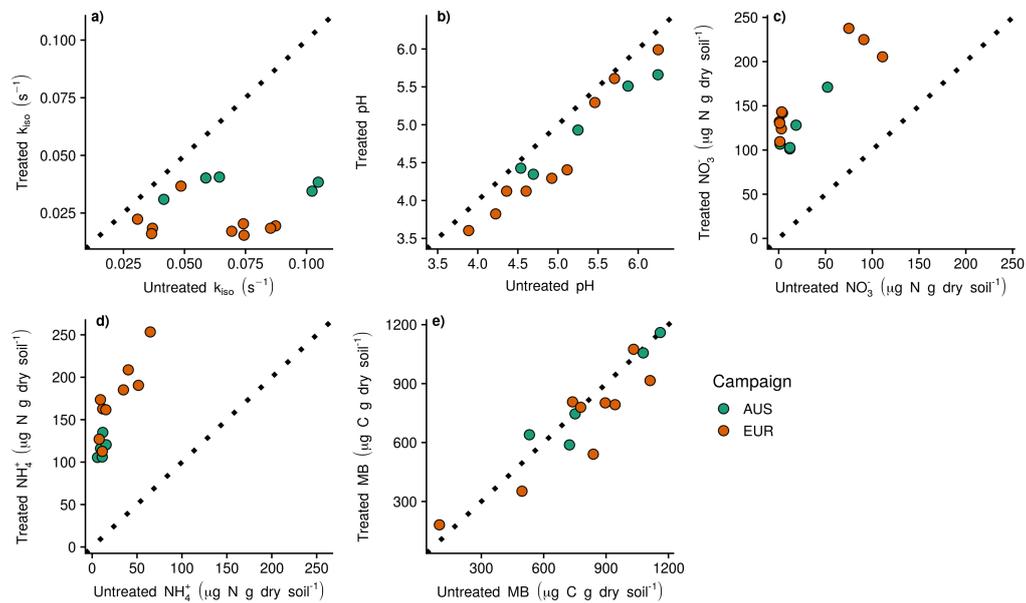


Figure S3: Mean a) k_{iso} , b) pH, c) exchangeable nitrate (NO_3^-), d) exchangeable ammonium (NH_4^+) and e) microbial biomass (MB) for the untreated control and the corresponding treated soils. Dashed lines indicating the 1:1 line with points below the line representing a decrease in treated relative to untreated soils and points above the line representing an increase. Points falling along the line indicate no change.

Ramirez, K. S., Craine, J. M. and Fierer, N.: Consistent effects of nitrogen amendments on soil microbial communities and processes across biomes, *Global Change Biology*, 18(6), 1918–1927, <https://doi.org/10.1111/j.1365-2486.2012.02639.x>, 2012.

Kaisermann, A., Jones, S. P., Wohl, S., Ogée, J. and Wingate, L.: Nitrogen Fertilization Reduces the Capacity of Soils to Take up Atmospheric Carbonyl Sulphide, *Soil Systems*, 2(4), 62, <https://doi.org/10.3390/soilsystems2040062>, 2018.

10) Minor comment: Remove commas when the citation is part of the sentence throughout the entire text: Jones et al. (2017), not Jones et al., (2017).

Corrected.

11) L21 suggest to omit ‘future’

This has been removed as part of the re-write and editing implemented in response to comment 3.

12) L31 ‘the $\delta^{18}O$ of leaf-atmosphere CO_2 exchange’ reads strange, sounds clearer with-out ‘exchange’

We have modified this as part of our response to comment 3) but we maintain the use of the word “exchange” in reference to leaf-atmosphere and soil-atmosphere exchanges through-out the text as it is important to do so.

13) L 35 CO_2 that interacts with a leaf – interacts seems too vague, could this be reworded?

This section has been rewritten as part of our response to comment 3).

14) L36 ‘undergo considerable enrichment’ of what?

This section has been rewritten as part of our response to comment 3).

15) L40 ‘alters the $\delta^{18}O$ of atmospheric CO_2 ’ was not really clear to me. Perhaps write ‘alters the $\delta^{18}O$ of CO_2 in soil vs. atmospheric CO_2 ’?

This section has been rewritten as part of our response to comment 3).

16) L51 is ‘of soil atmospheric CO_2 ’ not better?

This section has been rewritten as part of our response to comment 3).

17) L56 ‘abiotically invades’ sounds awkward, perhaps write ‘diffuses into the soil porenetwork’

We specifically used the term “invade” as we were referring to the invasion flux of CO₂ (also known as the piston velocity of the overlying air-column; L52).

18) L65 delete ‘true’, is confusing here

This section has been rewritten as part of our response to comment 3).

19) L99 better ‘agricultural soils’?

‘agriculture’ changed to ‘agricultural’ (L103)

20) L166 ‘an additional three replica incubations’ correct English?

We have re-written this paragraph to improve the clarity of our methodology.

L169 – L174: “An NH₄NO₃ addition experiment was also conducted in both campaigns. This involved the preparation of three additional 170 replicated incubations as described above, for nine of the EUR sites and five of the AUS sites. Prior to the pre-incubation step, 0.7 mg of NH₄NO₃ g dry soil⁻¹ was dissolved in water and used to adjust the water content of these additional replicate incubations. These were then incubated alongside the three other ‘control’ incubations prepared as part of the spatial survey described above.”

21) 2.1 Not clear if soils were kept cool during transport

We have clarified that EUR samples were transported at ambient temperatures (L135) and that AUS samples were returned to the laboratory on the day of sampling (L154).

22) L273 remove ‘presumably reflecting the pH dependency of NH₄⁺ and ammonia speciation’ this does not belong in the M&M section

Thanks, this has been deleted.

23) L293-294 do not seem entirely accurate: there seems to be a higher k_{iso} for the Csa group.

It’s true there are two sites associated with the Csa group that have high values above the 0.75 quantiles (i.e. the upper hinge of the box). However, these values are lower than the unique case for Bsh and an individual from the Cwa group. Furthermore the median of the Csa group clearly overlaps with the 0.50 to 0.75 quantile of both the Cfa and Cwa groups. On balance we feel that this does not support the conclusion that there is a robust pattern associated with these climate and landcover classifications. We do not attempt to test this statistically simply because despite our efforts the number of replicates for land-cover and climate still remains relatively low and their distribution is unbalanced.

24) L370-375 is quite unclear: are the 0.04 to 13 s⁻¹ field values derived from your labmeasured 0. To 0.4 s⁻¹ k_{iso} estimates? + which older studies. The whole part starts quite sudden. In fact for the sake of clarity the apparent issue of comparing lab and field estimates is perhaps best omitted from the paper entirely – also further on.

We feel it is important to include the comparison of k_{iso} observed in this study with those from the literature. This type of information has not typically been presented in recent publications in part because it is not always easy to compare reported values on an equal basis. However, by attempting to do so we do observe that there are potentially significant discrepancies that should not be ignored. We have adjusted and rephrased this paragraph.

L365 – 388: “This study aimed to reveal the drivers of variations in the oxygen isotope exchange rate, k_{iso}, to make it possible to predict the influence of different soil characteristics on the δ¹⁸O of atmospheric CO₂ and improve our understanding of soil CA activity. To do so, controlled incubation experiments were conducted to estimate k_{iso} from soils collected across western Eurasia and northeastern Australia. Estimates of k_{iso} for untreated soils in this study ranged from 0.01 to 0.4 s⁻¹ (Fig. 2 a). In all cases these rates exceeded theoretical uncatalysed rates (from 0.00008 to 0.008 s⁻¹ depending on soil pH, Uchikawa & Zeebe, 2012), indicating the presence of active CAs. The median k_{iso} of 0.07 s⁻¹ reported here is in the range of previously published values for sieved soils incubated in the dark (between 0.03 and 0.15 s⁻¹, Jones et al., 2017; Sauze et al., 2018, 2017) but lower than those reported by Meredith et al. (2019) with a median and range of 0.46 s⁻¹ and 0.08 to 0.88 s⁻¹, respectively. These greater k_{iso} values reported by Meredith et al. (2019) are more comparable to values (between 0.01 to 0.75 s⁻¹) reported by Sauze et al. (2017) for soils with well-developed phototroph communities. Direct comparison of our estimates of k_{iso} with those observed in the field is challenging because these older studies (Seibt et al., 2006; Wingate et al., 2008, 2009, 2010) estimated soil CA activity as a range of enhancement factors over a temperature sensitive uncatalysed rate of hydration. However, using the mid-

point of the enhancement factors and soil temperatures reported by Wingate et al. (2009), we estimate that k_{iso} varied between 0.04 and 13 s^{-1} with a median of 0.31 s^{-1} across the seven ecosystems considered in their analysis. Understanding why k_{iso} can be orders of magnitude greater in the field compared to values observed in laboratory incubations is a key question for further studies. Potentially, the abundance and activity of CAs may be reduced during the process of sieving soils and incubating them for prolonged periods in the dark. For example, the exclusion of intact roots and mycorrhizal fungi interacting within the rhizosphere might reduce k_{iso} (Li et al., 2005). Equally the suppression of phototrophic community members by incubating soils in the dark (Sauze et al., 2017) may also contribute to differences in k_{iso} between the field and such experiments. Furthermore, we cannot exclude the possibility that determining k_{iso} accurately under field conditions is less reliable. For example, the calculation of k_{iso} relies on determining the $\delta^{18}O$ of the soil water pool in equilibrium with CO_2 . Given the potential for increased heterogeneity in the soil water pool in natural conditions this may make it more challenging to determine k_{iso} robustly in the field (Jones et al., 2017). “

25) L431 ‘between untreated and treated’ is not a very clear phrasing – spell out better what you are referring at. We have rephrased this to hopefully make the meaning clearer.

L442-444: “Indeed, the ability of this model to reasonably predict fractional changes in k_{iso} between untreated control soils, that were used to build the model, and their fertiliser treated counterparts, that were not used to ‘train’ the model selection process, is encouraging (Fig. 4 b).”

26) L434 ‘A significant challenge to using this relationship to predict k_{iso} is likely the availability of suitable pedotransfer functions, particularly for NO_3^- availability and microbial biomass, to estimate patterns in the proposed drivers (Van Looy et al., 2017).’ Is an understatement. It will be impossible to predict the ephemeral soil NO_3^- levels with a simple pedotransfer function. The authors best refer to use soil N models to predict NO_3^- levels and use these as inputs into Eq 6.

We have rephrased this section in response to Reviewer 1, comment 22) and the point raised here.

L445 – L460: “A significant challenge to using this statistical relationship to predict k_{iso} is underpinned by our capacity to describe the spatial and temporal variations in the important drivers of k_{iso} , namely soil pH, microbial biomass and exchangeable NO_3^- . For this reason we also considered whether more readily available parameters such as soil texture, carbon content and nitrogen content might provide an alternative basis for empirical predictions of k_{iso} (Van Looy et al., 2017). However, relationships between these variables and k_{iso} were relatively weak and could only explain a marginal amount of the observed variability. Fortunately, a number of promising spatial databases are evolving for soil characteristics such as pH and microbial biomass (Serna-Chavez et al., 2013; Slesserev et al., 2016). Likewise a number of land surface models can now estimate the spatial and temporal dynamics of the biosphere nitrogen cycle convincingly (Zaehle, 2013). Predictions of soil nutrient dynamics will likely depend on the use of such advanced soil nitrogen cycle models. Given the interaction between soil pH and exchangeable NO_3^- (Fig. 3 a & b), the absence of such data may not seriously compromise predictions for fertilised agricultural soils as typically they are not strongly acidic. However, accurately predicting natural spatial and seasonal variability and the influence of future changes in atmospheric NO_3^- deposition (DeForest et al., 2004) may be more problematic. Nonetheless, the data reported in this study now lay the foundations for an empirical approach to predicting k_{iso} for a wide range of soils using readily available maps of key soil traits. This represents an important breakthrough in predicting how variations in soil community CA activity impacts the $\delta^{18}O$ of atmospheric CO_2 .”

RC1

1) My main concern is with the interpretation of the results of ammonium nitrate (NH₄NO₃) treatment. The authors attributed the decrease of k_{iso} following NH₄NO₃ addition to the inhibition of carbonic anhydrase caused by NO₃⁻. However, other possible mechanisms, namely, inhibition through increased ammonium content or decreased pH cannot be ruled out by the experimental design, nor by the statistical analysis that follows.

In essence, NH₄NO₃ addition may affect k_{iso} through these causal pathways:

- NH₄NO₃ addition → [NH₄⁺] increase → k_{iso} decrease
- NH₄NO₃ addition → [NH₄⁺] increase → pH decrease → k_{iso} decrease
- NH₄NO₃ addition → [NO₃⁻] increase → k_{iso} decrease

To accept Hypothesis 3, the authors must show evidence that after controlling for all confounding variables, including pH and [NH₄⁺], there is still a robust decrease of k_{iso} with the increase of [NO₃⁻]. Given the absence of a randomized design and the small sample size (n=14) for NH₄NO₃ addition treatment, it is difficult to identify [NO₃⁻] as the unique cause for carbonic anhydrase inhibition. One possible solution could be to treat pH and [NH₄⁺] as instrumental variables, but this would require them to show strong correlation with [NO₃⁻]. The best way would be to separate different causes through experimental design.

We agree that the experimental design of the ammonium nitrate treatment is not sufficient to fully tease apart the combined systematic effects (i.e. increased nitrate and ammonium availability and decreased soil pH) of the treatment on k_{iso}. As the reviewer states, a more extensive controlled factorial experiment would be required to achieve this. However, the results of this experiment (Section 3.2; Figure 5; Table S3), that show changes in k_{iso} are most strongly linked to changes in nitrate availability (pathway 3 above) and to a lesser degree soil pH (pathway 2 above) but do not appear related to changes in ammonium availability (pathway 1 above), are still informative to the interpretation of the wider study. Across the untreated soils we clearly identify soil pH, nitrate availability and microbial biomass as explaining variations in k_{iso} (Section 3.1; Figure 3; Table S1). Agreement between the results of both these analyses helps reinforce the importance of pH and nitrate (pathways 2 and 3) but not, directly at least, a role for ammonium (pathway 1). We have adjusted the text in the abstract and Section 4 to acknowledge that limitations of the experimental treatment prevent the definite conclusion that only nitrate, and not some combination of effects, influences the decrease in k_{iso} observed following the fertilisation treatment.

L25 - L26: *“This effect appears to be supported by a supplementary ammonium nitrate fertilisation experiment conducted on a subset of the soils”*

L426 – L430: *“It is important to note that whilst the relationship between the changes in k_{iso} and exchangeable NO₃⁻ are supported by observations from the untreated dataset, the experimental design used in this addition experiment is not sufficient to fully test the influence of the combined changes in soil pH, exchangeable NO₃⁻, exchangeable NH₄⁺ and microbial biomass on k_{iso}. Further controlled, factorial experiments are needed for this purpose.”*

2) A minor concern I have is that this study was partially motivated by the use of δ¹⁸O of CO₂ to estimate terrestrial photosynthesis. While the validity of this method has been demonstrated at the global scale by Welp et al. (2011), I would caution that it is un-clear whether the current in-situ observational network would provide sufficient data to resolve regional-scale photosynthesis. Nevertheless, in my opinion, soil–atmosphere CO₂ isotope exchange is an interesting topic for its own sake, regardless of whether δ¹⁸O–CO₂ can provide constraints on terrestrial photosynthesis with accuracy and spatio-temporal resolution as high as those of other photosynthetic tracers in vogue (e.g., solar-induced chlorophyll fluorescence).

We agree with the reviewer that the current in-situ observational network of δ¹⁸O in atmospheric CO₂ is rather coarse, at least compared to the network for total CO₂ mixing ratio. However, there are still more than 50 atmospheric stations measuring δ¹⁸O in CO₂, spread across all latitudes and continents, with some of them covering several decades of measurements. The extremely large north-south gradient of δ¹⁸O in CO₂ and its seasonal and interannual dynamics brings unique information on the seasonality and inter-annual variability of the northern hemisphere CO₂ sink, which is the strongest land carbon sink at the global scale and with the largest long-term trend (Ciais et al. 2019). Currently, this information is obscured by the lack of understanding of how soil dwelling organisms (and their carbonic anhydrase activity) affect this signal (Wingate et al. 2009). This study presents the largest soil dataset ever gathered on soil δ¹⁸O–CO₂ exchange. The in-depth analysis of the drivers of soil carbonic anhydrase activity that this study brings also serves as an important stepping-stone to study other emerging tracers of the carbon cycle including the Δ¹⁷O anomaly in CO₂ (Koren et al. 2019) and COS (Campbell et al., 2017). For all these reasons, it would seem awkward not to mention the implication this study will have in the future development of independent tracers to study the global carbon cycle. We also agree that solar-induced chlorophyll fluorescence (SIF), another independent proxy of photosynthesis, has the

advantage of being detected from space, conferring a global coverage. However the relationship between SIF detected from space and land photosynthesis is still not well understood, notably in disentangling structural and physiological factors. For this reason, SIF is most interesting at very high spatial resolution, which is only possible since the late 2010s, with satellite instruments like TROPOMI launched in 2017 or FLEX that will be launched in 2022. Here we do not pretend that $\delta^{18}\text{O}-\text{CO}_2$ is a more powerful tracer compared to other tracers of global photosynthesis (i.e. SIF or COS), but we are convinced that $\delta^{18}\text{O}-\text{CO}_2$ contains unique independent and historical information, strongly linked to the global water and carbon cycle, that cannot be discarded. This study, with its extensive survey of soil types and biomes, addresses one of the key knowledge gaps that currently prevent the routine use of $\delta^{18}\text{O}-\text{CO}_2$ as a global carbon tracer, and should motivate the community to reconsider this independent tracer in global climate models, thus constraining our understanding of variability in the northern hemisphere land carbon sink. Hopefully, this may also stimulate the development of a denser observational network of $\delta^{18}\text{O}$ in CO_2 , which is now possible with the next generation of laser-based CO_2 isotope analysers.

Campbell, J. E., Berry, J. A., Seibt, U., Smith, S. J., Montzka, S. A., Launois, T., Belviso, S., Bopp, L. and Laine, M.: Large historical growth in global terrestrial gross primary production, *Nature*, 544(7648), 84, <https://doi.org/10.1038/nature22030>, 2017.

Ciais P, Tan J, Wang X et al. (2019) Five decades of northern land carbon uptake revealed by the inter-hemispheric CO_2 gradient. *Nature*, 568, 221–225.

Koren G, Schneider L, van der Velde IR et al. (2019) Global 3-D Simulations of the Triple Oxygen Isotope Signature $\Delta 17\text{O}$ in Atmospheric CO_2 . *Journal of Geophysical Research-Atmospheres*, 127, 73.

Wingate L, Ogee J, Cuntz M et al. (2009) The impact of soil microorganisms on the global budget of $\delta 18\text{O}$ in atmospheric CO_2 . *Proceedings of the National Academy of Sciences of the United States of America*, 106, 22411–22415.

3) The writing needs more clarity and conciseness. As a rule of thumb, try not to make sentences more complicated than the ideas they convey. In a paragraph, stick to one point and avoid switching topics or walking back and forth. For example, much of the discussion had the main points hidden in the middle of a paragraph and could use some restructuring. Break long paragraphs if necessary.

Following this comment and similar comments made in EC1 we have re-written the introduction and edited both abstract and text to improve the readability of the manuscript.

4) The hypotheses need to be accurately framed. Hypothesis 2 is a complicated statement, and the only part testable based on your experiments is that k_{iso} increases with soil pH. The rest of Hypothesis 2 describes possible mechanisms and they cannot be answered by your experiments. In Hypothesis 3, you can only test whether k_{iso} increases with $[\text{NO}_3^-]$, but not whether $[\text{NO}_3^-]$ binds carbonic anhydrases or how it inhibits carbonic anhydrase. These two hypotheses should be precisely worded as testable hypotheses. The hypotheses you actually tested were stated in P14L382–383, so why not simplify them just like that?

We have simplified our hypotheses to following this advice.

L112 – L 115: “Based on the potential controls on k_{iso} presented above we tested three specific, non-exclusive, hypotheses; 1) k_{iso} increases as microbial biomass increases (H1), 2) k_{iso} increases as soil pH increases (H2), and 3) k_{iso} decreases as the presence of NO_3^- increases (H3).”

5) Finally, I encourage the authors to make the data sets publicly available in a data repository. This would make the study more easily discoverable and facilitate data reuse in future studies, for example, comparison across sites and parameterization of related soil processes in a land biosphere model.

We have archived the data with PANGAEA.

L662 – L664: “The data produced in this study have been archived with PANGAEA (<https://doi.org/10.1594/PANGAEA.928394>). The data may also be requested from the corresponding author by email.”

Specific comments

6) P1L13: “The expression and activity of carbonic anhydrase [. . .]” - You may need to tell the reader that carbonic anhydrase regulates the hydration of CO₂ in soil pore-space water before you mention that it drives k_{iso}.

We have rephrased this sentence as suggested.

L17 – L 19: “As the carbonic anhydrases (CAs) group of enzymes enhances the rate of CO₂ hydration within the water-filled pore spaces of soils it is important to develop understanding of how environmental drivers can impact k_{iso} through changes in their activity. “

7) P1L19–20: “[. . .] potentially reflecting the direct or indirect inhibition of carbonic anhydrases” - Is there a way to tell which mechanism is more likely?

To distinguish whether the impact of nitrate is direct or indirect an integrated study looking into changes in the concentration of carbonic anhydrase protein and the abundance of carbonic anhydrase transcripts alongside measurements of k_{iso} would be required. Additionally it would also be important to do some detailed protein studies that show the physical interaction of nitrate with the carbonic anhydrase protein and develop a method that could quantify the binding efficiency of nitrate to carbonic anhydrase for a few of the dominant soil carbonic anhydrases e.g. the beta-CA class. Collectively these different experiments would help us tease apart the direct and indirect effects of nitrate on carbonic anhydrase in soils.

8) P2L31: “because the δ18O of leaf–atmosphere CO₂ exchange tends to be enriched [. . .]” - More precisely, this is because leaf preferentially uses lighter isotopologues of CO₂, which diffuse faster than heavier ones. See Farquhar et al. (1993) Nature (<https://doi.org/10.1038/363439a0>).

This section has now been removed as part of the re-write of the introduction. However, diffusion is not the only reason. We agree that the oxygen isotope composition of leaf-atmosphere CO₂ exchange is partly explained by fractionation during diffusion, but not only by this. The isotopic exchange between CO₂ and water is also very important (Farquhar et al. 1993). In contrast the influence of oxygen isotope fractionation during other steps of fixation (e.g. carboxylation) is limited because carbonic anhydrase concentrations are sufficiently high enough for the isotopic equilibration between CO₂ and water to be extremely rapid (Ogée et al. 2018). By analogy to ¹³C fractionation during photosynthesis, Farquhar et al. (1993) described the leaf as consuming isotopically lighter CO₂ in terms of ¹⁸O, thereby leaving behind CO₂ enriched in ¹⁸O in the intercellular air space to diffuse back to the atmosphere. However, the analogy works because the CO₂ inside the leaf equilibrates its oxygen isotopes with evaporatively enriched leaf water. Thus, the mechanism is very different than for ¹³C, and primarily driven by leaf water isotopic composition and secondarily by diffusion.

Farquhar, G. D., Lloyd, J., Taylor, J. A., Flanagan, L. B., Syvertsen, J. P., Hubick, K. T., Wong, S. C. and Ehleringer, J. R. (1993) Vegetation effects on the isotope composition of oxygen in atmospheric CO₂, Nature, 363(6428), 439–443, doi:10.1038/363439a0.

Ogée J, Wingate L, Genty B (2018) Estimating mesophyll conductance from measurements of C18OO photosynthetic discrimination and carbonic anhydrase activity. Plant Physiol., 178, 728–752.

9) P2L44: “Comprising at least six distinct families, [. . .]” - There are seven now, with the newly discovered ι -CA in phytoplanktons. See Jensen et al. (2019) ISME J (<https://doi.org/10.1038/s41396-019-0426-8>).

The introduction and references now include this information.

L60 – L62: “Currently, at least seven distinct CA gene families have been identified, with each catalysing the reversible hydration of CO₂ to bicarbonate (Jensen et al., 2019).”

L540 – L545: “Jensen, E. L., Clement, R., Kosta, A., Maberly, S. C. and Gontero, B.: A new widespread subclass of carbonic anhydrase in marine phytoplankton, The ISME Journal, 13(8), 2094–2106, doi:[10.1038/s41396-019-0426-8](https://doi.org/10.1038/s41396-019-0426-8), 2019.”

10) P3L81–82: “Whilst the sensitivity of soil k_{iso} to the presence of specific functional groups, like phototrophs which employ carbonic anhydrases in their carbon concentration mechanisms [. . .]” - Are phototrophs abundant in soil microbial communities?

In a review of the literature Wingate et al., 2009 estimated that soil algal populations of between 10³ - 10⁶ per gram of soil are typically present in most soils. If cyanobacteria are further included, phototrophs can indeed form an important

part of the soil microbial community under many conditions (Muriel Bristol Roach, 1927; Seppey et al., 2017). This may be either as superficial crusts or within the near surface. Whilst they are likely to be less ubiquitous than fungi and bacteria, the possibility of specialised, carbonic anhydrase dependent, carbon concentration mechanisms might suggest their presence could have a disproportionately strong influence on k_{iso} . In a previous study looking at the role of phototrophs on carbonic anhydrase activity (Sauze et al., 2017) we developed a qPCR approach that helped us show that the putative natural abundance of soil phototrophs derived from the number of 23S reads were relatively small under darkened conditions compared to the bacterial (16S) and fungal (18S) abundances but their relative abundances increased significantly when incubated in the light. This probably and unsurprisingly suggests that such an influence might be somewhat dependent on the canopy cover and light conditions of the system in question.

Muriel Bristol Roach, B. (1927). On the algae of some normal English soils. *The Journal of Agricultural Science*, 17(4), 563-588. doi:10.1017/S0021859600018839

Seppey, C. V. W., Singer, D., Dumack, K., Fournier, B., Belbahri, L., Mitchell, E. A. D. and Lara, E.: Distribution patterns of soil microbial eukaryotes suggests widespread algivory by phagotrophic protists as an alternative pathway for nutrient cycling, *Soil Biology and Biochemistry*, 112, 68–76, doi:10.1016/j.soilbio.2017.05.002, 2017.

Sauze, J., Ogée, J., Maron, P.-A., Crouzet, O., Nowak, V., Wohl, S., Kaisermann, A., Jones, S. P. and Wingate, L.: The interaction of soil phototrophs and fungi with pH and their impact on soil CO₂, CO₁₈O and OCS exchange, *Soil Biology and Biochemistry*, 115(Supplement C), 371–382, doi:10.1016/j.soilbio.2017.09.009, 2017.

Wingate L., Ogée J., Cuntz M., B. Genty, I. Reiter, U. Seibt, D. Yakir, K. Maseyk , E.G. Pendall, M.M. Barbour, B. Mortazavi, R. Burrett, P. Peylin, J. Miller, M. Mencuccini, J.H. Shim, J. Hunt, J. Grace (2009) The impact of soil microorganisms on the global budget of $\delta^{18}\text{O}$ in atmospheric CO₂. *Proceedings of the National Academy of Sciences of America*, 106, 22411–22415.

11) P4L99: Be specific about “the inorganic nitrogen chemistry of soil solutions.”

We have re-written the introduction in response to the comments about writing style. This section has changed accordingly.

L100 – L107: “Various anions may also play a role in controlling the activity of CAs (Tibell et al., 1984). In particular, nitrate (NO₃⁻) has been shown to inhibit different CAs in a range of microbes and plants (Amoroso et al., 2005; Innocenti et al., 2004; Peltier et al., 1995). This suggests that variations in soil nutrient availability between ecosystems could give rise to differences in k_{iso} . Furthermore, the addition of common fertilisers such as ammonium nitrate (NH₄NO₃) to agricultural soils could have an inhibitory role on CA activity in addition to causing shifts in the size and composition of microbial communities present. Indeed, this hypothesis is supported by recent NH₄NO₃ fertilising experiments that demonstrated decreases in the CA catalysed hydrolysis of carbonyl sulphide (Kaisermann et al., 2018b). So far, the impact of nitrates on k_{iso} has not been investigated in soils.”

12) P5L133–134: Does sieving affect carbonic anhydrase activity in soils?

Our experiments did not test for the impact of sieving on soil carbonic anhydrase activity and as far as we are aware this has not been reported in the literature, thus the nature of these effects is not well understood and is discussed in L366 to L388.

13) P7L195–198: What was the precision of the IRIS for CO₂ and $\delta^{18}\text{O}$ -CO₂ measurements when averaged in 40 intervals?

We have added this information..

L200 - L201: “The associated precision for the total concentration and $\delta^{18}\text{O}$ of CO₂ was 0.02 ppm and 0.06 ‰ VPDBg respectively.”

14) P7L210: Eq. (1) requires a steady-state condition. What is the turnover time for gas exchange in the cuvette? Could you show that the measurement period (12, P1L191) is much longer than this turnover time?

The turnover time was less than 10 minutes. We have added this information to the text. Each jar was flushed for 20 or 22 minutes before the measurement period (L193 - L194) and 22 or 24 minutes before the first measurement of the

chamber line was made. These timings reflect the need to balance the trade-off between approximate steady-state conditions and changes in the isotopic composition of the soil water pool (Jones et al. 2017).

L194: “*The turnover time of air in the jar was less than 10 minutes.*”

Jones, S. P., Ogée, J., Sauze, J., Wohl, S., Saavedra, N., Fernández-Prado, N., Maire, J., Launois, T., Bosc, A. and Wingate, L. (2017) Non-destructive estimates of soil carbonic anhydrase activity and associated soil water oxygen isotope composition, *Hydrology and Earth System Sciences*, 21(12), 6363–6377, doi:<https://doi.org/10.5194/hess-21-6363-2017>.

15) P8L238–239: Please considering providing a table of site information and soil characteristics, either as a supplementary table or a metadata file in the online data set associated with this study. Although such information is available for European sites in Kaisermann et al. (2018) ACP, it would not be convenient for C4 the reader to reference across multiple publications. For the Australian sites, I do not see any such data.

We have archived the data with PANGAEA.

L662 – L664: “*The data produced in this study have been archived with PANGAEA (<https://doi.org/10.1594/PANGAEA.928394>). The data may also be requested from the corresponding author by email.*”

16) P10L281: What does the “two-term model” mean? What are the predictors?

Two-term models are those limited to 2 or less predictive terms. We have rephrased this to make it clearer.

L286 – L 289: “*The same approach was also applied to the 27 soils from the EUR sampling campaign and extended to consider the relationships with soil texture and carbon and nitrogen contents to investigate their utility in upscaling efforts. To prevent over-fitting, these models were limited to a maximum of two of predictive terms. The predictive terms considered were soil sand, silt, clay, carbon and nitrogen content, the ratio of carbon to nitrogen content and soil pH.*”

17) P10L282: Have soil texture, carbon content, and nitrogen content been considered in the aforementioned model selection procedures?

The same model selection procedures were used and this now explicitly stated in the text (L286).

18) P11L305: “*Correlations between all other variable pairings were weaker and non-significant ($p > 0.05$).*” - I find this observation in apparent conflict with the interpretation of NH_4NO_3 treatment results. If NO_3^- concentration does not control k_{iso} in natural soils, why would adding NH_4NO_3 cause k_{iso} to decrease through carbonic anhydrase inhibition? One possible scenario could be that the variation in k_{iso} that is attributable to soil pH is so large that any influence from NO_3^- concentration is obscured. To test whether this would be the case, Spearman’s rank correlation would be insufficient. You would need to control for the variation due to pH before testing the effect of $[\text{NO}_3^-]$.

Spearman’s rank correlation is used to identify the strongest patterns between pairs of variables without making a priori assumptions about the data. This is particularly useful as it helps us identify potential co-correlations such as that between pH and ammonium availability that may confound the subsequent analyses discussed in the paragraph following that referred to in this comment.

Subsequent use of multiple generalised linear models lets us test these relationships in a more satisfactory fashion. This analysis bears out the main result of the Spearman’s rank correlation i.e. that most of the variability in k_{iso} is explained by soil pH. However, after controlling for the effect of pH the inclusion of nitrate availability and biomass both significantly increase the degree of variability explained (see also Table S1). This indicates that nitrate concentration does indeed control k_{iso} in natural soils. Figure 3 b shows the nature of this relationship with nitrate concentration, particularly under acidic conditions, causing k_{iso} to decrease.

19) P13L357: While the fraction of explained deviance is high, this is a small sample with $n=14$ and uncertainty associated with # the model could be large. What is the confidence interval of the coefficient of $\ln \text{NO}_3^-$?

We agree that the sample size is small and report this model simply as the best fit to the data out of the variables considered in order to understand the influence of the treatment on the rate of exchange. Indeed, the uncertainty is large particularly at higher values of change. Please see the confidence interval provided in Figure 5.

20) P13L376–380: “*Whether the potential [. . .] remains an unresolved but key question.*” - Not sure what you are trying to mean with this sentence. Please clarify it.

This section has been edited in response to comments about the writing style.

L380 – 388: *“Understanding why k_{iso} can be orders of magnitude greater in the field compared to values observed in laboratory incubations is a key question for further studies. Potentially, the abundance and activity of CAs may be reduced during the process of sieving soils and incubating them for prolonged periods in the dark. For example, the exclusion of intact roots and mycorrhizal fungi interacting within the rhizosphere might reduce k_{iso} (Li et al., 2005). Equally the suppression of phototrophic community members by incubating soils in the dark (Sauze et al., 2017) may also contribute to differences in k_{iso} between the field and such experiments. Furthermore, we cannot exclude the possibility that determining k_{iso} accurately under field conditions is less reliable. For example, the calculation of k_{iso} relies on determining the $\delta^{18}O$ of the soil water pool in equilibrium with CO_2 . Given the potential for increased heterogeneity in the soil water pool in natural conditions this may make it more challenging to determine k_{iso} robustly in the field (Jones et al., 2017).”*

21) P15L425: *“The absence of strong patterns with climate or land-cover in this study may well reflect the fact that the temperature and moisture conditions used are unrepresentative of field conditions especially for colder and drier sites.” - Or, it could also be that soil texture and composition are the main controls.*

It is true that the conditions experienced by the microbes in their natural environments can be very different from those experienced in our experiment. This would definitely be interesting to look at in the future with a different experimental and mechanistic modelling approach. However, the aim of the present study was to standardise moisture and temperature conditions to the best of our abilities and investigate how the gas exchange rates and enzyme activity of these different communities compared. Opting for this experimental design meant we were not able to attribute statistically whether differences in activity were underpinned by land-use or climate class in a way that would facilitate a simple scaling up approach. Our study indicates other soil traits such as pH have the potential to provide more reliable spatial predictions of k_{iso} . With larger databases perhaps land-use or climate patterns will begin to emerge as important large-scale drivers of soil function and predictors of soil-atmosphere gas exchange but for the moment it remains unclear as these datasets are rare in the community.

22) P15L435: *What are the “pedotransfer functions?”*

Pedotransfer functions are predictive functions used to estimate certain soil properties from more readily available data. We have altered this section in response to comments about the writing style and EC1 comment 26)

L445 – L460: *“A significant challenge to using this statistical relationship to predict k_{iso} is underpinned by our capacity to describe the spatial and temporal variations in the important drivers of k_{iso} , namely soil pH, microbial biomass and exchangeable NO_3^- . For this reason we also considered whether more readily available parameters such as soil texture, carbon content and nitrogen content might provide an alternative basis for empirical predictions of k_{iso} (Van Looy et al., 2017). However, relationships between these variables and k_{iso} were relatively weak and could only explain a marginal amount of the observed variability. Fortunately, a number of promising spatial databases are evolving for soil characteristics such as pH and microbial biomass (Serna-Chavez et al., 2013; Slesserev et al., 2016). Likewise a number of land surface models can now estimate the spatial and temporal dynamics of the biosphere nitrogen cycle convincingly (Zaehle, 2013). Predictions of soil nutrient dynamics will likely depend on the use of such advanced soil nitrogen cycle models. Given the interaction between soil pH and exchangeable NO_3^- (Fig. 3 a & b), the absence of such data may not seriously compromise predictions for fertilised agricultural soils as typically they are not strongly acidic. However, accurately predicting natural spatial and seasonal variability and the influence of future changes in atmospheric NO_3^- deposition (DeForest et al., 2004) may be more problematic. Nonetheless, the data reported in this study now lay the foundations for an empirical approach to predicting k_{iso} for a wide range of soils using readily available maps of key soil traits. This represents an important breakthrough in predicting how variations in soil community CA activity impacts the $\delta^{18}O$ of atmospheric CO_2 .”*

Technical comments

23) P1L10: *“gross primary production”* vs. P1L25 *“gross primary productivity”* (emphases mine), pick one.

“gross primary production” is used through-out the text.

24) P1L11: *“ecosystem-scale”* → *“ecosystem scale”*

This was removed as part of the improvements to the writing requested in other comments.

25) P1L15: Add a comma before *“indicating [. . .]”*

A comma was added.

L21 – L 22: “Observed values for k_{iso} always exceeded theoretically-derived uncatalysed rates, indicating a significant influence of CAs on the variability of k_{iso} across the soils studied.”

26) P1L33: “the leaves of plants” → “leaves”. Pleonasm.

This was removed as part of the improvements to the writing requested in other comments.

27) P2L35: “causing CO₂ that interacts with a leaf but is not fixed to inherit the isotopic composition of the leaf water pool” - A difficult sentence. Please clarify.

This was removed as part of the improvements to the writing requested in other comments.

28) P2L44–P3L73: This paragraph has a lot to unpack. In my opinion, to bring clarity to this paragraph, you may consider splitting it into two. Describe the abiotic reaction of oxygen isotope exchange first, and then introduce the role of carbonic anhydrases in accelerating the reaction towards equilibrium. I would consider splitting the paragraph at line 62 and rearranging sentences for a clean separation.

The introduction has been re-written in response to this and other comments about the writing style (L35 – L100).

29) P3L83: “it’s” → “its”

Removed as part of the re-write of the introduction in response to comments about the writing style.

30) P3L87–89: “Such an observation may result from changes in size or composition of the microbial communities involved as discussed (Sauze et al., 2017, 2018).” - This is a reiteration of P3L79–81.

Removed as part of the re-write of the introduction in response to comments about the writing style.

31) P4L95: “non-carbon” → “non-carbonate”

This was changed as part of the response to comments about the writing style.

L93: “Various anions may also play a role in controlling the activity of CAs (Tibell et al., 1984).”

32) P5L123: “principle” → “principal”

Corrected (L118)

33) P5L124: “indicted” → “indicated”

Corrected (L126)

34) P6L171: This should be section 2.2, not 2.1.

Corrected (L176).

35) P11L312–316 and P12L330–337: It is inconvenient to track which model is which. Please consider listing model diagnostics in supplementary tables.

We have added three tables to the supplement listing the relevant models discussed in the text.

Table S1: Ranking and included terms for a subset of the generalised linear models tested to predict variations in the rate of oxygen isotope exchange, k_{iso} , for the entire dataset ($n = 44$). Model selection was limited to a maximum of four predictive terms and the intercept. The terms MB, NO_3^- and NH_4^+ are the natural logarithms of microbial biomass and nitrate and ammonium availability. Selected terms or interactions within each model are indicated by + symbols whilst - symbols indicate their omission. The interactions Campaign:pH and Campaign:MB are omitted from the table for

brevity as they were not selected in any of the models shown. Model ranking was based on comparison of sample size corrected Aikake's Information Criterion (AICc) with Δ AICc indicating the difference in AICc from the best model. Δ AICc of 2 or more indicates real differences in model performance.

Rank	Intercept	Campaign	pH	MB	NO ₃ ⁻	NH ₄ ⁺	Campaign: NO ₃ ⁻	pH: MB	pH: NO ₃ ⁻	MB: NO ₃ ⁻	NO ₃ ⁻ : NH ₄ ⁺	Δ AICc
1	+	-	+	+	+	-	-	-	+	-	-	0.00
2	+	-	+	+	+	-	-	-	-	-	-	6.10
3	+	+	+	+	+	-	-	-	-	-	-	7.06
4	+	-	+	+	+	-	-	+	-	-	-	7.07
5	+	+	+	-	+	-	+	-	-	-	-	7.09
6	+	-	+	+	+	-	-	-	-	+	-	8.79
7	+	+	+	-	+	-	-	-	-	-	-	12.43
8	+	-	-	+	+	+	-	-	-	-	+	13.27
16	+	-	+	-	-	-	-	-	-	-	-	21.56
19	+	-	-	-	-	+	-	-	-	-	-	26.48
21	+	-	-	+	-	-	-	-	-	-	-	43.64
28	+	-	-	-	-	-	-	-	-	-	-	47.91
33	+	+	-	-	-	-	-	-	-	-	-	50.15
34	+	-	-	-	+	-	-	-	-	-	-	50.21

Table S2: Ranking and included terms for a subset of the generalised linear models tested to predict variations in the rate of oxygen isotope exchange, k_{iso} , for the relatively invariant soil properties of the EUR campaign dataset (n = 27). Model selection was limited to a maximum of two predictive terms and the intercept. The terms C, N and CN are soil carbon and nitrogen content and their ratio. Selected terms within each model are indicated by + symbols whilst - symbols indicate their omission. Model ranking was based on comparison of sample size corrected Aikake's Information Criterion (AICc) with Δ AICc indicating the difference in AICc from the best model. Δ AICc of 2 or more indicates real differences in model performance.

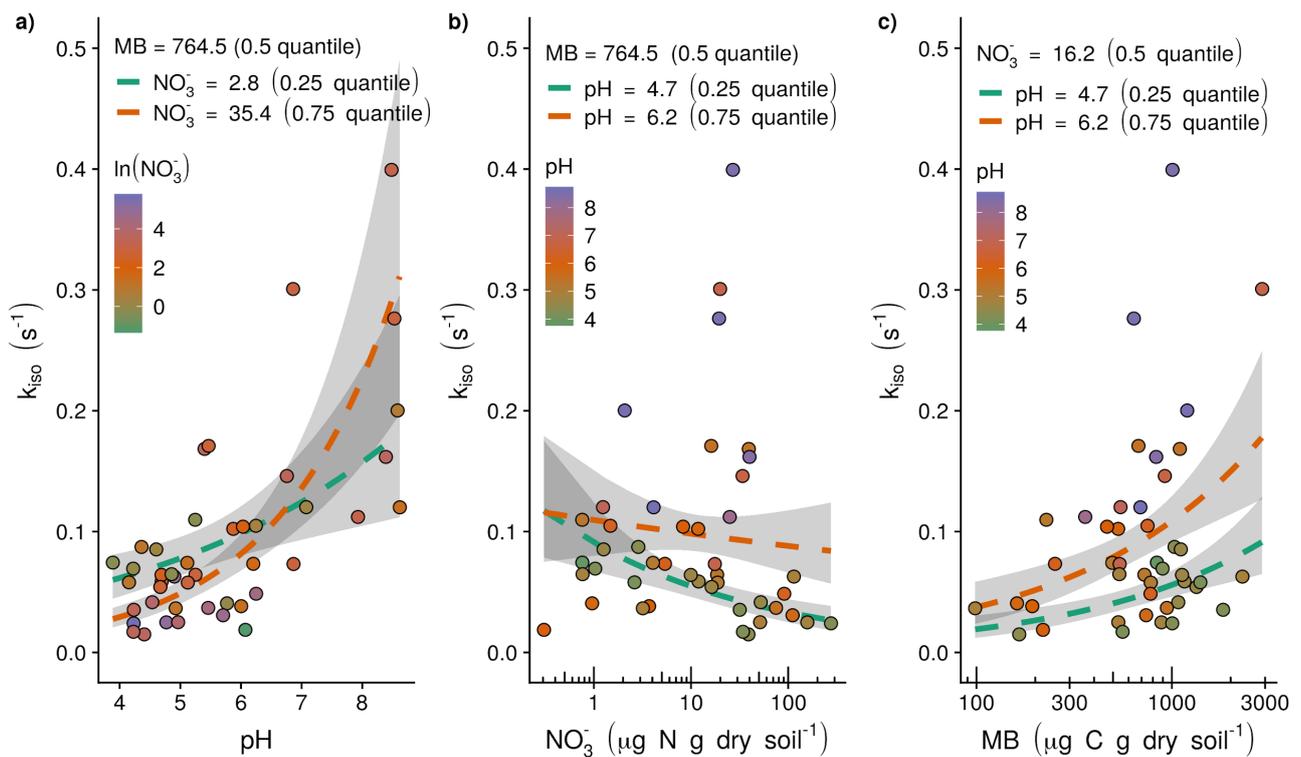
Rank	Intercept	pH	Sand	Silt	Clay	C	N	CN	Δ AICc
1	+	+	-	-	+	-	-	-	0.00
2	+	+	+	-	-	-	-	-	0.57
3	+	+	-	-	-	+	-	-	1.32
4	+	+	-	-	-	-	-	-	1.85
5	+	+	-	-	-	-	-	+	1.92
6	+	+	-	+	-	-	-	-	2.46
7	+	+	-	-	-	-	+	-	4.57
8	+	-	-	-	-	-	+	-	21.26
9	+	-	-	-	-	-	-	-	22.07

Table S3: Ranking and included terms for a subset of the generalised linear models tested to predict variations in the change in rate of oxygen isotope exchange, k_{iso} , following ammonium nitrate addition (n = 15). Model selection was limited to a maximum of one predictive term and the intercept. The terms MB, NO₃⁻ and NH₄⁺ are differences in microbial biomass and nitrate and ammonium availability following ammonium nitrate addition whilst the prefix ln indicates the natural logarithm of these differences. Selected terms within each model are indicated by + symbols whilst - symbols indicate their omission. Model ranking was based on comparison of sample size corrected Aikake's Information Criterion (AICc) with Δ AICc indicating the difference in AICc from the best model. Δ AICc of 2 or more indicates real differences in model performance.

Rank	Intercept	Campaign	pH	MB	NO ₃ ⁻	NH ₄ ⁺	lnMB	lnNO ₃ ⁻	lnNH ₄ ⁺	ΔAICc
1	+	-	-	-	-	-	-	+	-	0.00
2	+	-	-	-	+	-	-	-	-	8.65
3	+	-	+	-	-	-	-	-	-	13.20
4	+	-	-	-	-	-	-	-	-	15.95
5	+	+	-	-	-	-	-	-	-	17.38
6	+	-	-	-	-	-	+	-	-	18.34
7	+	-	-	+	-	-	-	-	-	18.80
8	+	-	-	-	-	-	-	-	+	19.10
9	+	-	-	-	-	+	-	-	-	19.21

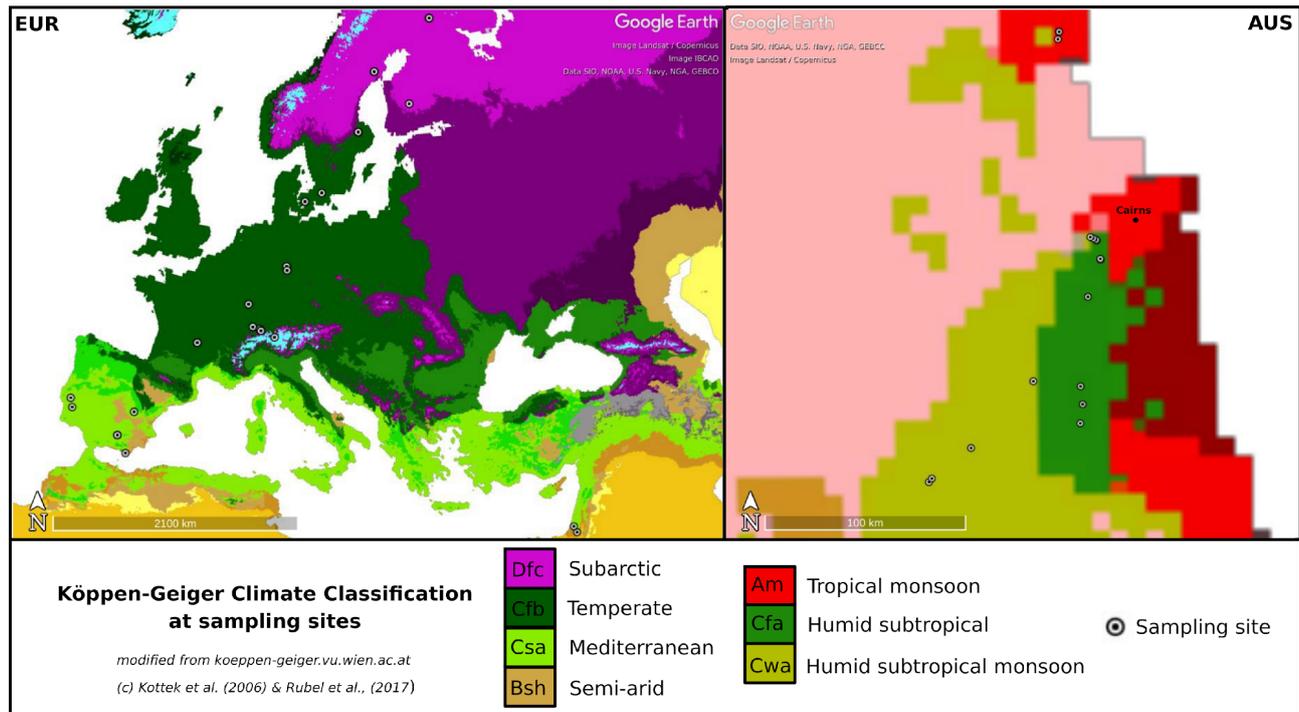
36) Figure 3: It is difficult to distinguish high values from low values indicated by the color bars. Try to increase the contrast.

Figure 3 has been revised to hopefully increase the contrast of the plot gradients and use a more accessible colour palette.



37) Figure S1: Remove the ocean background and other unnecessary information. Please simplify this figure to make the ecoclimatic classification more evident. Consider putting the legend outside of the figure canvas to avoid interference.

Figure S1 has been revised to mask the ocean, move the legend outside of the map area and reduce the classes to only reflect those covered by the samples obtained.



SC1

1) I have read your work with great interest. The exchange of oxygen isotopes between CO_2 and soil water is an important process for $\delta^{18}\text{O}$, and this work contributes to a better understanding of that exchange. However, this exchange is also of great importance for the budget of $\Delta^{17}\text{O}$ in CO_2 , a different tracer for GPP. $\Delta^{17}\text{O}$ in CO_2 was first proposed as a tracer of GPP by Hoag et al. (2005). More recently, laboratory studies confirmed the effect of photosynthesis on $\Delta^{17}\text{O}$ in CO_2 (Adnew et al., 2020), and we simulated large-scale variations of $\Delta^{17}\text{O}$ in atmospheric CO_2 (Koren et al., 2019). We struggled with representing the soil exchange in that model, and for follow-up studies we can possibly improve our representation of soil exchange using Eq. 6 from your manuscript. I think you can reach a greater audience if you also explicitly address the $\Delta^{17}\text{O}$ community in your work.

We have now added a couple of sentences in the introduction to clarify this point.

L35 – L45: “Quantifying the carbon storage potential of terrestrial ecosystems and its sensitivity to climate change relies on our ability to obtain observational constraints of photosynthesis and respiration at large scales (Beer et al., 2010). Over recent decades there has been increasing interest in using the oxygen isotope composition ($\delta^{18}\text{O}$ and $\delta^{17}\text{O}$) of atmospheric carbon dioxide (CO_2) to trace these large and opposing CO_2 fluxes. This is possible because the $\delta^{18}\text{O}$ of leaf-atmosphere CO_2 exchange is relatively enriched in ^{18}O compared to that of atmospheric CO_2 and the $\delta^{18}\text{O}$ of soil-atmosphere CO_2 exchange (Francey & Tans, 1987; Wingate et al., 2009; Welp et al., 2011). Similarly, photochemical processes in the stratosphere cause anomalies between the $\delta^{17}\text{O}$ and $\delta^{18}\text{O}$ of atmospheric CO_2 that are subsequently reset during leaf-atmosphere CO_2 exchange (Hoag et al., 2005; Koren et al., 2019; Adnew et al., 2020). However, the routine use of these tracers to constrain the photosynthetic term of the atmospheric mass budget for the $\delta^{18}\text{O}$ and $\delta^{17}\text{O}$ of CO_2 has been hampered by an incomplete understanding of how the influence of soil-atmosphere CO_2 exchange varies across different soil types and environmental conditions. Here we focus on $\delta^{18}\text{O}$ but the key challenges to understanding these variations are also relevant to considerations of $\delta^{17}\text{O}$.”

2) In the first line and last line of the abstract I would replace " $\delta^{18}\text{O}$ " with " $\delta^{18}\text{O}$ and $\Delta^{17}\text{O}$ ".

We have rephrased the abstract to remove the emphasis on only ^{18}O and replaced this with a more general reference to the oxygen isotope composition of atmospheric CO_2 (L11 - L31).

3) Sec 2. Are you sure that the sampling and transporting of soil samples does not affect the CA or microbes in the sample?

We do indeed expect there to be a disturbance effect on the microbial community when transporting soils and sieving them, thus it is important to be mindful of this when comparing results from soils measured under field conditions and

those measured in laboratory experiments as well as extrapolating results from mesocosms to the large scale. This study however was designed to characterize a set of homogenized climate-controlled soils to make a link between the measured CA activity, the mesocosm soil characteristics and their response to changes in inorganic N concentrations. However the quantitative influence of transport and sieving on carbonic anhydrase activity is so far not well understood but is discussed. Please see L366- L388 in the Discussion.

4) There are two sections with number 2.1.

Corrected (L176).

5) L139: "Tillburg". This should be the lovely city "Tilburg".

Corrected (L142).

6) L147: Why did you choose to report on the VPDBg scale, instead of e.g. VSMOW?

We preferentially report our CO₂ in air measurements on the VPDBg scale (also known as VPDB-CO₂ scale) reflecting the fact that values assigned to our working standards are ultimately tied to the acid digestion of RM NBS-19 calcite. Please see:

Werner, R. A., Rothe, M. and Brand, W. A.: Extraction of CO₂ from air samples for isotopic analysis and limits to ultra high precision $\delta^{18}\text{O}$ determination in CO₂ gas, Rapid Communications in Mass Spectrometry, 15(22), 2152–2167, doi:<https://doi.org/10.1002/rcm.487>, 2001.

Werner, R. A. and Brand, W. A.: Referencing strategies and techniques in stable isotope ratio analysis, Rapid Communications in Mass Spectrometry, 15(7), 501–519, doi:<https://doi.org/10.1002/rcm.258>, 2001.

7) L210: The units provided in the text do not agree with Eq. 1.

Corrected.

L215: "where u is the flow rate (mol s^{-1}) through the chamber line".

8) L423: I would briefly mention $\Delta^{17}\text{O}$ here.

We now reference this in the discussion.

L435: "Improvements in our ability to predict soil k iso and its influence on the $\delta^{18}\text{O}$ of atmospheric CO₂ are important in refining the use of this tracer and others such as ^{17}O to constrain photosynthesis and respiration at large scales (Wingate et al., 2009; Welp et al., 2011; Koren et al., 2020)."

9) Caption Fig. 1: The authors mention twice: "dissolved organic carbon (DIC)". Should this be DIC or DOC?

Corrected to DIC.