1 **Title: Comparing three approaches of spatial disaggregation of legacy soil maps based on**
2 **DSMART algorithm**

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33 **Abstract:**

34 Enhancing the spatial resolution of pedological information is a great challenge in the field of Digital Soil

35 Mapping (DSM). Several techniques have emerged to disaggregate conventional soil maps initially

36 available at coarser spatial resolution than required for solving environmental and agricultural issues. At the

37 regional level, polygon maps represent soil cover as a tessellation of polygons defining Soil Map Units

38 (SMU), where each SMU can include one or several Soil Type Units (STU) with given proportions derived

39 from expert knowledge. Such polygon maps can be disaggregated at finer spatial resolution by machine

40 learning algorithms using the Disaggregation and Harmonisation of Soil Map Units Through Resampled

41 Classification Trees (DSMART) algorithm. This study aimed to compare three approaches of spatial

42 disaggregation of legacy soil maps based on DSMART decision trees to test the hypothesis that the

43 disaggregation of soil landscape distribution rules may improve the accuracy of the resulting soil maps.

44 Overall, two modified DSMART algorithm (DSMART with extra soil profiles, DSMART with soil

45 landscape relationships) and the original DSMART algorithm were tested. The quality of disaggregated soil

46 maps at 50 m resolution was assessed over a large study area (6,775 km²) using an external validation based

47 on independent 135 soil profiles selected by probability sampling, 755 legacy soil profiles and existing

48 detailed 1:25,000 soil maps. Pairwise comparisons were also performed, using Shannon entropy measure,

49 to spatially locate differences between disaggregated maps. The main results show that adding soil landscape

50 relationships in the disaggregation process enhances the performance of prediction of soil type distribution.

51 Considering the three most probable STU and using 135 independent soil profiles, the overall accuracy

52 measures are: 19.8 % for DSMART with expert rules against 18.1 % for the original DSMART and 16.9 %

53 for DSMART with extra soil profiles. These measures were almost twofold higher when validated using

54 3x3 windows. They achieved 28.5% for DSMART with soil landscape relationships, 25.3% and 21% for

55 original DSMART and DSMART with extra soil observations, respectively. In general, adding soil

56 landscape relationships as well as extra soil observations constraints the model to predict a specific STU

57 that can occur in specific environmental conditions. Thus, including global soil landscape expert rules in

58 the DSMART algorithm is crucial to obtain consistent soil maps with clear internal disaggregation of SMU

59 across the landscape.

60 **Key words:** digital soil mapping, soil landscape relationships, spatial disaggregation, DSMART

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65   1) Introduction

66 Characterizing soil variability especially over large areas, remains a crucial challenge to foster

67 sustainable management of agronomic and environmental issues and help stakeholders to design

68 regional projects (Chaney et al., 2016). At the regional as well as country level, soil maps are often

69 available at coarse spatial resolution (Bui and Moran, 2001) which limits their ability to depict

70 accurate soil information. For instance, the finest soils maps covering France were elaborated by

71 administrative region at 1:250,000 scale, via a set of polygons, called Soil Map Units (SMU) with

72 crisp boundaries. The delineation of SMU is based on soil survey programmes involving

73 pedologists' expertise. In a coarse scale map, each polygon includes one or several Soil Type Unit

74 (STU), which are not explicitly mapped, but their proportions and their environmental conditions,

75 as well as soil characteristics, are provided in a detailed database (Le Bris et al., 2013).

76 To improve soil variability knowledge and overcome the limitation of a coarse mapping scale,

77 several methods have emerged in the field of Digital Soil Mapping (DSM). These methods offer

78 useful tools to predict soil spatial pattern from scarce or limited soil datasets by exploiting the

79 availability of model based methods and an extensive array of spatialise (and more often than not

80 gridded) environmental variables. In recent decades, DSM techniques have been increasingly used

81 to downscale soil information and improve their spatial resolution. Depending on the quality of

82 data and the complexity of soil cover, Minasny and McBratney (2010) supply a workflow that

83 outlines different models that can be explored. In general, two main pathways can be distinguished:

84 point based DSM approaches and map disaggregation approaches (Odgers et al., 2014; Holmes et

85 al., 2015). Point DSM approaches used legacy soil profiles, which are irregularly distributed and

86 collected according to specific objectives rather than to optimise a statistical criterion (Holmes et

87 al., 2015). The spatial distribution of soil properties can be estimated by fitting geostatistical

88 models such as ordinary kriging (Odgers et al., 2014; Holmes et al., 2015; Chaney et al., 2016;

89 Vincent et al., 2018; Chen et al., 2018) or cokriging, which takes into account the spatial

90 interrelations among several soil properties (Webster and Oliver, 2007). Additionally, McBratney

91 et al. (2003) developed the SCORPAN soil landscape model. It is an empirical quantitative function

92 of environmental covariates, allowing predicting soil attributes (soil type or soil property) based

93 on correlative and statistical relationships with predictor variables.

94   The second approach, known as spatial disaggregation, attempts to downscale the soil map unit
95   information to delineate unmapped STUs (Bui and Moran, 2001; Odgers et al., 2014; Holmes et
96   al., 2015). Alternatively, it can be defined as the process that allows estimating soil properties at a
97   finer scale than the initial soil map. Several techniques have been demonstrated through soil science
98   literature and tested in different case studies around the world. For instance, Kempen et al. (2009)
99   have explored the use of multinomial logistic regression (MLR) for digital soil mapping. Other
100   techniques have also been applied as decision trees using rule based induction (Bui and Moran,
101   2001), Bayesian techniques (Bui et al., 1999) and an area to point kriging method (Kerry et al.,
102   2012).

103   In the DSM field, machine learning techniques are increasingly used to elucidate the spatial
104   distribution of both soil type and soil properties across a large range of scale (Bui and Moran.,
105   2001; Scull et al., 2005; Lacoste et al., 2011; Lemercier et al., 2012; Nauman and Thompson, 2014;
106   Holmes et al., 2015; Vaysse and Lagacherie, 2015; Ellili et al., 2019). They were also applied to
107   disaggregate superficial geology maps available at 1: 250 000 scale in Australia (Bui and Moran,
108   2001). The main advantage of these approaches is they allow handling both quantitative and
109   categorical (ordinal or nominal) soil and environmental variables, as explanatory covariates (Bui
110   and Moran, 2001).

111   Odgers et al. (2014) have developed a machine learning algorithm entitled Disaggregation and
112   Harmonisation of Soil Map Units Through Resampled Classification Trees (DSMART) to predict
113   STU as a function of the high resolution environmental data supplied over different study areas in
114   Australia. The DSMART algorithm is based on a calibration dataset derived from a random
115   selection of a fixed number of sampling points within each soil polygon. Each sampling point is
116   then assigned to one soil type following a weighted random allocation procedure based on the
117   proportions informed in the soil map database. The same procedure was applied by Chaney et al.
118   (2016) to spatially disaggregate the soil map of the contiguous United States at a 30m spatial
119   resolution using petascale High Performance Computer (HPC). Because integration of pedological
120   knowledge has been recognized as an effective way to improve digital soil mapping approaches
121   (Cook et al., 1996; Walter et al., 2006; Stoorvogel et al., 2017; Machado et al., 2018; Møller et al.,
122   2019), Vincent et al. (2018) have applied the DSMART algorithm with additional expert soil
123   landscape rules describing soil distribution in the local context of the Brittany region (France). By

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124   adding supplement sampling points to the calibration dataset selected according to soil parent
125   material, soil redoximorphic conditions and topographic features, and by integrating soil landscape
126   relationships in the DSMART sample allocation scheme, the authors obtained a coherent soil
127   spatial distribution observing soil organisation along hillslopes and occurrence of intensely
128   waterlogged soils in the stream neighbourhood, as observed in Brittany.

129   This study aimed to test the hypothesis that adding soil landscape relationships in the disaggregation
130   procedure improved the accuracy of produced disaggregated soil maps. This involves assessing the
131   contribution of soil landscape relationships implemented in the DSMART algorithm by Vincent et
132   al. (2018). To achieve this objective, we compared disaggregated soil maps either derived from the
133   original DSMART algorithm, the DSMART algorithm with extra soil observations and the
134   DSMART algorithm fed by soil landscape relationships over an area of 6,775 km² in the eastern part
135   of Brittany, France.

2) Materials and methods

136

137   2.1) Study area

138   The Ille et Vilaine department covers an area of 6,775 km² and is located at the eastern part of
139   Brittany, France (48°N, 2° W) (Fig 1). It is drained by the rivers Ille and Vilaine and their
140   tributaries. Its climate is oceanic, with a mean annual rainfall of 669 mm and mean annual
141   temperature of 11.3° (Source: Climate Data EU). Main land uses comprise arable land, temporary
142   and permanent grasslands, woodland, and urban areas. In the present study, anthropogenic areas
143   were not considered. Elevation ranges between 0_20 m in the coastal zone and 20_150 m almost
144   everywhere expect in the western part of the department where it tills 256 m. The topography is
145   generally gentle with maximum slopes not exceeding 16%. The Ille et Vilaine department is part
146   of the Armorican Massif with complex geology (BRGM, 2009): intrusive rocks (granite, gneiss
147   and micaschist) in northern and north western zones, sedimentary rocks (sandstone) and
148   metamorphic rocks (Brioverian schist) in the central and southern zones, and superficial deposits
149   (Aeolian loam with decreasing thickness from north to south overlaying bedrock, alluvial and
150   colluvium deposits). According to the World Reference Base of Soil Resources, soils occurring in
151   Ille et Vilaine include Cambisols, Luvisols Stagnic Fluvisols, Histosols, Podzols, and Leptosols
152   (IUSS Working Group WRB, 2014).

153   2.2) Soil data

154   2.2.1) Regional soil database at 1:250 000 scale

155 In Brittany, soils are represented through a regional geographic database called "Référentiel

156 Régional Pédologique (RRP)" available at 1:250,000 scale (INRA Infosol, 2014).This regional

157 database identifies soils within Soil Map Units (SMUs), each containing one to several soil types

158 called Soil Type Units (STUs). STUs are defined as areas with homogeneous soil forming factors,

159 such as morphology, geology, and climate. In the study area, 96 SMU and 171 STU have been

160 distinguished and represented by a spatial coverage of 479 polygons.

161 In the regional database, SMUs were spatially delimited with crisp boundaries, while STUs were

162 not explicitly mapped, but their proportion in each SMU as well as associated environmental and

163 soil characteristics were accurately described in a semantic database (Le Bris et al., 2013; INRA

164 Infosol, 2014).

165   2.2.2) Soil validation data

166 To assess the quality of disaggregated soil maps, three validation datasets were used (Fig. 1):

167 • 135 soil profiles chosen following a stratified random sampling design and specifically

168   described and sampled from March to May 2017 for independent validation purposes in the

169   framework of the Soilserv research project (Ellili et al., 2019, submitted).

170 • 755 legacy soil profiles collected between 2005 and 2008 during the "Sols de Bretagne"

171   programme (INRA Infosol, 2014).These profiles were sampled to characterize

172   hydromorphic soil conditions and soil landscape heterogeneity.

173 Existing detailed soil maps (1:25,000) covering 87,150 ha, surveyed according to Rivière et al.

174 (1992) and revised later to adapt to the STU typologies developed in the RRP (Le Bris et al., 2013).

175
176 All soil profiles were allocated after description and analysis by an expert to a suitable STU. Both

177 legacy soil profiles and detailed maps were converted to raster format to perfectly meet the

178 prediction raster at 50m spatial resolution.

179   2.3) Environmental covariates

180 The SCORPAN concept (McBratney et al., 2013) allows one to predict STU as a function of a set

181 of covariates describing seven soil forming factors, namely soil properties (s), climate (c),

182 organisms (o), relief (r), parent material (p), age (a) and geographic position (n). In this study, ten

183   environmental variables (Table 1) were considered as covariates in the disaggregation process at a

184   50m spatial resolution. Terrain attributes included elevation, slope, Compound Topographic Index

185   (CTI) (Beven and Kirkby, 1979, Merot et al., 1995) and Topographic Position Index (TPI) (Vincent

186   et al., 2018) that together were derived from a 50m resolution Digital Elevation Model (IGN, 2008).

187   These attributes were computed using ArcGIS 10.1 (ESRI, 2002) and MNT surf software

188   (Squividant, 1994).

189   Environmental attributes describing soil parent material (Lacoste et al., 2011) and hydromorphic

190   soil conditions via waterlogging index (Lemercier et al., 2013) were obtained using decision tree

191   methods. Waterlogging index derives from a natural soil drainage prediction. Four classes were

192   distinguished: well drained, moderately drained, poorly drained and very poorly drained. Aeolian

193   silt deposits and Soil Map Units boundaries are environmental covariates also obtained via expert

194   knowledge from soil scientists.

195   Landscape units reflecting vegetation, land use, and relief attributes were derived from a MODIS

196   imagery by supervised classification (Le Du Bayo et al., 2008). The Airborne gamma ray

197   spectrometry variable (K:Th ratio) (Messner, 2008), characterizing the degree of weathering of the

198   geological material, was also taken into account.

199   All soil environmental covariates were converted to raster format at 50 m spatial resolution.

200   2.4) Disaggregation procedure: DSMART algorithm

201   2.4.1) Original DSMART algorithm (Method 1)

202   The open source DSMART algorithm (Odgers et al., 2014) was applied to spatially disaggregate

203   the existing legacy soil map at 1:250,000 scale. DSMART algorithm uses machine learning

204   classification trees implemented in C5.0 (Quinlan, 1993) to build a decision tree from a target

205   variable (STU) and the environmental covariates supplied. The DSMART algorithm was written

206   in the Python programming language by Odgers et al. (2014) and was recently translated in the R

207   programming language.

208   Running DSMART algorithm requires four main steps (Fig. 2):

209   1) Polygon sampling by a random selection of a fixed number of sampling points (n=30)
210   within each polygon. This procedure allowed to select a total of 14,370 sampling points,
211   per iteration, covering the study area and ensured that all polygons were sampled.

212   2) Soil Type Unit (STU) assignment to each sampling point following a weighted random
213   allocation method. This step was based on the proportion of each STU informed in the RRP
214   database.

215   3) Decision tree generation: the full set of sampling points were spatially intersected with the
216   selected environmental covariates. This georeferenced dataset was then used as a
217   calibration dataset to build the decision tree allowing the prediction of an STU as a function
218   of environmental covariates. C5.0 created explicit models, which were applied to the
219   covariates rasters to generate a realisation of STU distribution over the study area at 50 m
220   resolution.

221   These three steps were repeated 100 times to generate 100 realisations of the potential soil type
222   distribution over the study area at 50 m of resolution.

223   4) Computing the probabilities of occurrence: the 100 realisations were stacked to calculate
224   the probability of occurrence of each predicted STU by counting the frequency of each STU
225   at each pixel. This procedure led to a set of 171 rasters depicting the probability of
226   occurrence of 171 STU.

227

228   2.4.2) Original DSMART algorithm + soil observations (method 2)

229

230   This disaggregation approach is similar to the original DSMART algorithm. However, the main
231   difference is that 755 additional soil profiles, spatially collocated, were added to the calibration
232   dataset to build decision trees. These soil profiles make it possible to incorporate real field
233   observations with established soil landscape relationships. For each realisation, a calibration
234   dataset (15, 125 samples) including virtual samples randomly selected from polygon units, as well
235   as soil observations were used to model soil type with environmental covariates. The model was
236   then extrapolated over the study area.

237

238   2.4.3) Original DSMART algorithm + expert rules (Method 3)

239   Including soil landscape relationships in the disaggregation process was explored by Vincent et al.

240   (2018) in a specific regional pedoclimatic context in Brittany (France). Expert soil landscape

241   relationships were used to assign STU to sampling points. These relationships were based on expert

242   pedological knowledge, which takes into account soil parental material as well as topography and

243   waterlogging in the UTS allocation procedure. This approach combines two sources of the dataset

244   to calibrate the model. The first one was derived from semantic information for each SMU/STU

245   combination. It consists in attributing a barcode to each SMU/STU combination, derived from a

246   concatenation of four features contained in the RRP database (parent material, SMU identifier, TPI

247   and waterlogging index), and to compare these barcodes to a stack of regional covariates

248   representing the same four features, to assign each pixel of the study area to a suitable STU. This

249   procedure allowed matching soil exhibiting specific features with their potential spatial

250   distribution. For instance, hydromophic soils occur with slope sequences and valley positions,

251   while well drained soils occur in upslope or middle slope positions. Using a random sampling

252   stratified by SMU's area, a set of sampling points was selected with a proportion of one sample for

253   every 5 hectares and a minimum of five samples per polygon unit.

254   The second dataset was derived from a random sampling of a fixed number of sampling points in

255   each polygon unit. This procedure ensured that all polygons had been sampled. STU allocation was

256   based on the soil map unit proportions. The full set of each realisation (18, 320 samples) combining

257   expert calibration dataset as well as dataset derived from random sampling procedure was spatially

258   intersected with existing environmental covariates and used as a unique calibration dataset to build

259   decision trees.

260

261   2.4.4) Prediction of the most probable STUs

262   From all soil type probability rasters obtained, only the three most probable STUs (with the highest

263   probability of occurrence) were considered: for each pixel, the final prediction was the combination

264   of the three most probable predicted STUs (1st STU, 2^{sd} STU, and 3rd STU) and their associated

265   probability of occurrence.

266   The classification confusion index (CI) between the first most probable STU and the second most

267   probable STU was calculated following Eq.1:

268    $CI = 1 - (P_{1^{st} STU} - P_{2^{sd} STU})$ [1]

269   Where $P_1^{st}{}_{STU}$ and $P_2^{sd}{}_{STU}$ denote respectively the highest probability of occurrence for 1st STU

270   and the second highest probability of occurrence for 2^{sd} STU, calculated at each pixel (Burrough et

271   al., 1997; Odgers et al., 2014).

272   This index was considered as an indicator of certainty assessment about the most probable

273   predicted soil class and is ranging between 0 and 1. It tends to 1. When the 1st STU and 2^{sd} STU

274   are predicted with similar probability of occurrence and zero when the probability of occurrence

275   of the 2^{sd} STU is close to zero.

276

277   2.5) Validation of disaggregated soil maps

278   The quality of soil maps resulting from the three DSMART algorithm based approaches was

279   assessed by combining both spatial and semantical validation methods. Spatial validation is divided

280   into 2 sub approaches ("pixel to pixel" and "window of 3x3 pixels"). For detailed soil maps and

281   accurate soil profiles, "pixel to pixel" validation consists in checking, at each pixel, if the predicted

282   STU respects the observed STU value (Heung et al., 2014; Nauman et al., 2014; Chaney et al.,

283   2016; Møller et al., 2019). The "window of 3x3 pixels" validation assumes that, for each pixel, the

284   predicted STU respects the observed STU value if it matches at least one of its 9 surrounding

285   neighbours (Heung et al., 2014; Chaney et al., 2016). This method provides some flexibility by

286   compensating spatial referencing error of soil maps and avoids the impact of fine scale spatial

287   noise.

288   The semantical validation was also performed considering either each STU or a group of STUs

289   sorted by expert on the basis of similar pedogenesis factors and similar diagnostic horizons

290   (Vincent et al., 2018; Møller et al., 2019). From the initial 171 STUs described in the soil database,

291   the sorting procedure led to 78 groups and 11 STU remained single.

292

293   2.6) Pair wise comparisons of disaggregated soil maps

294   To compare the soil type rasters derived from the three DSMART based approaches, pairwise

295   comparisons were performed using *Vmeasure* method implemented as open source software in an

296   R package called Spatial Association Between REgionalisations (SABRE) (Rosenberg and

297   Hirschberg, 2007). This is a spatial method developed to compare maps in the form of vector

298   objects and it was commonly used in computer science to compare (non spatial) clustering.

299 We divide the entire study area into 2 different sets of regions, referred to as regionalizations R and

300 Z. The first regionalization R divides the domain into n regions r_i (i=1 to n) and the second

301 regionalization Z divides the domain into m zones z_j (j=1 to m). Superposition of the 2

302 regionalization R and Z divides the domain into n x m segments having a_{ij} area. The total area of a

303 region r_i is $A_i = \sum_{j,1}^{m} a_{ij}$, the total area of a zone z_j is $Aj = \sum_{i,1}^{n} a_{ij}$ and the total of the domain is

304 $A = \sum_{j=1}^{m} \sum_{i=1}^{n} a_{ij}$.

305 The SABRE package calculates a degree of spatial agreement between two regionalizations using

306 an information theoretical measure called the *V measure*. *V measure* provides two intermediate

307 metrics: *homogeneity* and *completeness*. *Homogeneity* is a measure of how well regions from the

308 first map fit inside zones from the second map (Eq 2). *Completeness* measures how well zones

309 from the second map fit inside regions from the first map (Eq 5). The final value of *V measure* is

310 calculated as the weighted harmonic mean of homogeneity and completeness (Eq 8). All metrics

311 range between 0 and 1, where larger values indicate better spatial agreement. *V measure*,

312 homogeneity, and completeness are global measures of association between the two

313 regionalizations.

314 Additional indicators of disaggregation quality were calculated using Shannon entropy index of

315 regions and zones (Shannon 1948; Nowosad and Stepinskie, 2018). These indicators qualify local

316 associations by highlighting the region's inhomogeneities (Eq 3, Eq 4), or zone's inhomogeneities

317 (Eq 6, Eq 7). Two normalized Shannon entropy was also computed using the ratios (S_j^R/S^R) and

318 (S_i^Z/S^Z) to derive maps of local spatial agreement between the two regionalizations R and Z. These

319 measures have a range between 0 and 1.

320 When S_j^R (Eq3) is close to zero, this denotes that the zone j is homogenous in terms of regions

321 (each zone is within a single region). However, when S_j^R value increases the zone is increasingly

322 inhomogeneous in terms of regions (it overlays an increasing number of regions). Therefore, S_j^R

323 (Eq 3) assesses the degree of this inhomogeneity or a variance of region in zone *j*. A global indicator

324 that measures a homogeneity of a given zone in terms of regions is given via Eq 2.

325 Analogous to homogeneity but with the roles of regions and zones reversed, the dispersion of zones

326 over the entire area is also computed using Shannon entropy (Eq 4 and Eq 7), and a global indicator

327 *C* (Eq 5) measures a homogeneity of a given region in terms of zones.

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328    $h = 1 - \sum_{j=1}^{m}(\frac{A_j}{A})\left(\frac{Variance\ of\ regions\ in\ zone_j = S_j^R}{Variance\ of\ regions\ in\ the\ domain = S^R}\right)$ [2]

329    $S_j^R = -\sum_{i=1}^{n}(\frac{a_{i,j}}{A_j}) \log(\frac{a_{i,j}}{A_j})$ [3]

330    $S^R = -\sum_{i=1}^{n}(\frac{A_i}{A}) \log(\frac{A_i}{A})$ [4]

331    $c = 1 - \sum_{i=1}^{n}(\frac{A_i}{A})\left(\frac{Variance\ of\ zones\ in\ region_i = S_i^Z}{Variance\ of\ zones\ in\ the\ domain = S^Z}\right)$ [5]

332    $S_i^Z = -\sum_{j=1}^{m}(\frac{a_{i,j}}{A_i}) \log(\frac{a_{i,j}}{A_i})$ [6]

333    $S^Z = -\sum_{j=1}^{m}(\frac{A_j}{A}) \log(\frac{A_j}{A})$ [7]

334    $V_\beta = \frac{(1+\beta)hc}{(\beta h)+c}$ [8]

335   β is a coefficient that allows promoting the first or the second regionalization, and by default, β

336   equals 1. V_β has a range between 0 and 1. It equals 0 in case of no spatial association and 1 in case

337   of perfect association.

338   The *V measure* method was applied in two main situations (DSMART+expert rules, Original

339   DSMART) and (DSMART + expert rules, DSMART+extra soil observations). The reference map

340   is always the map derived from DSMART algorithm with expert soil landscape relationships.

341   3)   Results

342   3.1) Disaggregated soil maps

343   Applying DSMART based approaches yielded a set of soil maps and associated probability of

344   occurrence rasters. The original DSMART approach allowed to disaggregate the 96 SMUs into

345   108 STUs while DSMART with expert rules approach yielded 158 STUs and DSMART with extra

346   soil observations approach yielded 172 STUs with respect to the first most probable STU map. A

347   total of 171 STUs were identified in the Ille et Vilaine department within the RRP database.

348   Unpredicted STUs correspond mainly to rare STUs with low proportions ranging between 2 and

349   10% within the SMUs containing them.

350   Figure 3 shows the three maps of the 1st most probable STU derived from each approach as well

351   as the original soil map. Overall, the three most probable STUs maps captured the main pattern of

352   soil distribution of the coarse soil map. As one could expect according to the geological parent

353   material map (Lacoste et al., 2011), extensive areas of deep silty soils are developed in Aeolian

354   loam deposits encountered in the north east as well as in the north central parts of the study area.

355   Colluvial and alluvial soils were mainly predicted in the north coast part and large valleys zones.

356   The visual comparison of disaggregated soil maps highlighted global similarities in the soil spatial

357   distribution markedly affected by SMU boundaries. The three approaches distinguished very well

358   soils developed in marsh parent material in the coastal part (north) of the study area. However,

359   DSMART with soil landscape expert rules map as well as DSMART with extra soil observations

360   map remained more detailed and underlined a clear internal disaggregation of SMUs especially in

361   the north and the central parts of the Ille et Vilaine department. Visual inspection of the obtained

362   DSMART with extra soil observations map as well as DSMART with expert rules map showed an

363   increase in soil heterogeneity when compared to Original DSMART map. More importantly,

364   legacy soil profiles made it possible to take into account some rare soil types with low probability

365   to be predicted. Therefore, adding supplement sampling points via the expert calibration dataset

366   and the 755 extra soil profiles allowed to predict STUs characterized in the soil database with a

367   low spatial extent. Nevertheless, the three DSMART based approaches spatially disaggregated the

368   most frequent components disregarding the less frequent ones.

369   Figure 4 shows maps of the global probability of redoximorphic soils across the study area. STU

370   probability rasters, depicting hydromorphic soils, were added together to produce continuous maps

371   of hydromorphic soil probability. Visual inspection of three maps highlighted global similarities,

372   but local differences were recorded along the hydrographic network and in the southern part of the

373   study area. As could be expected, DSMART with expert rules well predicted hydromorphic soils

374   in valleys and coastal areas, with a probability of occurrence exceeding 80%. Adding soil landscape

375   relationships in the allocation process constrained hydromorphic soil predictions in specific

376   landscape positions. The same trend characterized DSMART with extra soil observations map,

377   particularly in the central part of the study area. Therefore, including 755 soil profiles had an

378   important role in the disaggregation process in the northern and the central parts where these

379   profiles were located.

380   The quality of maps resulting from DSMART based approaches was quantified via the probabilities

381   of occurrence of each STU predicted and the confusion index maps (Fig. 5). The latter measure

382   indicated areas where the probability of occurrence of the two most probable soil types was close.

383   Over the study area, the average probability of occurrence of the most probable soil type achieved

384   respectively 0.41 for DMSART map, 0.68 for DMSART with expert rules and 0.28 for DSMART

385   with extra soil observations maps. Meanwhile, the average confusion index reached 0.8 for the

386   original DSMART approach while DSMART with extra soil observations and DSMART with

387   expert rules achieved 0.9 and 0.43, respectively. Although the most probable soil classes provide

388   plausible maps of soil distribution, there is a significant prediction uncertainty as depicted by these

389   measures.

390   In regions where disaggregated soil maps showed low confusion index, particularly in northwest

391   and north coast areas of Ille et Vilaine department, high confidence in predictions was recorded.

392   These areas were predominantly deep loamy soils or developed in alluvial and colluvium deposits.

393   Figure 6 compares the cumulative area of the STUs estimated from the three disaggregated maps

394   and that derived from the regional soil database. For each STU, its relative predicted area was

395   estimated by counting the number of pixels where it was predicted. For the regional soil database,

396   each STU area was computed from total SMU area multiplied by the proportion of the STU. This

397   comparison shows that some STUs were overestimated by the disaggregation approaches when

398   comparing to the soil database. DSMART with extra soil observations and original approaches

399   showed similar cumulative STU areas under the curve whereas DSMART with expert rules had a

400   shape similar to the regional soil database.

401   The most abundant STU in the database (431: Stagnic Fluvisol developed from alluvial and

402   colluvium deposits) was predicted as the most frequent STU by DSMART with extra soil

403   observations and DSMART with expert rules, and it was predicted as the second most abundant

404   STU by the original DSMART algorithm. The 10 most abundant STUs in the soil database covers

405   almost 43% of the study area. Of them, 7 belong to the 10 STU most predicted by the three

406   disaggregation approaches (Table 2).

407   3.2) Covariates importance in the decision trees

408   Figure 7 gives the relative importance of the covariates used in DSMART based approaches. Soil

409   parent material and SMU boundaries were used systematically in condition rules regardless of the

410   disaggregation method. This was consistent with the contrasting pattern of geology and the

411   dependence relationship between SMU and its soil components. Considering the original

412   DSMART approach (Fig. 7.a), distribution functions of Aeolian silt deposits, airborne gamma ray

413   spectrometry variable (K:Th ratio) and elevation contributions were more dispersed according to

414   the STU considered than those of other covariates. For instance, Aeolian silt deposits contribution

415   varied between 20 and 80% with a median value of 42%, whereas slope contribution ranged

416   between 20 and 40 % with a median value of 28%. Aeolian silt deposits have an important weight

417   in STU predictions, due to its ability to represent soils inherited from this superficial parent

418   material, which is poorly represented in lithological maps.

419   DSMART with soil landscape relationships (Fig. 7.b) showed almost the same distribution function

420   of all covariates except for elevation where its distribution function was more dispersed. Since a

421   part of training samples was chosen with expert knowledge based on three environmental

422   covariates: TPI, a waterlogging index and soil parent material, we would expect the prominent role

423   of waterlogging index and TPI to constrain hydromorphic soils predictions and to achieve STU

424   distribution in the appropriate order along the toposequence. This most likely explains the

425   dominance of Fluvisol Stagnic in valleys areas followed by a transition to Cambisols commonly

426   found at upslope and midslope positions along the toposequences.

427   Analogous to the original DSMART algorithm, DSMART with extra soil observations (Fig. 7.c)

428   highlighted almost the same distribution of use of soil environmental covariates in the decision

429   trees, except for aeolian silt deposits, K:Th ratio and elevation. The latter covariates contributions

430   remained less dispersed compared to the original DSMART approach.

431   3.3) Validation of disaggregated soil maps

432   The validation procedure was performed for each DSMART based approach applied, considering

433   the three most probable soil types and using both semantic objects (STU or soil group) and spatial

434   neighbourhood (per pixel or 3x3 window of pixels).

435   Considering 755 legacy soil profiles prospected in the framework of "Sols de Bretagne" project,

436   per pixel validation accuracy reached 27%, for original DSMART maps and 34 % for DSMART

437   with expert rules (Table 3). A similar comparison using 135 validation sites derived from Soilserv

438   project showed that 18.1 % of soil profiles match DSMART maps, 19.8 % match DSMART with

439   expert rules maps and only 16.9 % match DSMART with extra soil observations maps (Table 3).

440 Using a 3 x 3 window of pixels markedly improves the global accuracies, which increased for the

441 two validation datasets (Table 3). DSMART with soil landscape relationships remained the best

442 performing method.

443 When compared to accurate soil maps (1:25,000), the validation procedure showed that DSMART

444 with extra soil observations as well as DSMART with soil landscape expert rules had almost the

445 same performance (37% and 38%) while best accuracy (44%) was observed for Original DSMART

446 maps (44%) (Table 3). These scores were clearly improved by considering soil groups and 3x3

447 pixels neighbourhood. For instance, the accuracy of DSMART with expert rules maps using soil

448 group reached 45.9% and increased to 62.1 % when considering 3x3 pixels windows (Table 3).

449 3.4) Comparing disaggregated maps

450 Figure 8 shows inhomogeneity maps measured by Shannon entropy. The map derived from

451 DSMART with soil landscape relationships was chosen as a reference map. This map deeply

452 disaggregates the initial SMUs into 120,653 regions with irregular shapes. By contrast, Original

453 DSMART map remained very similar to the original map and delineated the study into 40,459

454 regions. Both disaggregated maps reflect the main pattern of soil distribution over the study area

455 despite the difference in the disaggregation process. Visual inspection of maps DSMART with soil

456 landscape rules map and Original DSMART map revealed an overall similarity between

457 disaggregated maps, but local differences between them were depicted.

458 We calculated $h_1 = 0.49$, $c_1 = 0.58$ and $V_1 = 0.53$ as global measures of spatial agreement between

459 the two maps (DSMART+expert rules and Original DSMART). The average homogeneity of the

460 DSMART with soil landscape rules map with respect to the Original DSMART map was qualified

461 via h homogeneity index. Similarly, the average homogeneity of the Original DSMART map with

462 respect to the DSMART with soil landscape rules map was qualified via c completeness index.

463 Visually, the Fig. 8.b map seemed to be more homogeneous than the map Fig. 8.a in agreement

464 with the statistical assessment $c > h$. The large number of DSMART with soil landscape rules map

465 regions, which was three times higher than Original DSMART map zones, might explain this

466 difference. It is more likely that DSMART with soil landscape rules map regions cross through

467 multiple Original DSMART map zones than vice versa. However, two disaggregated maps

468 remained spatially associated according to the high V_1 score. The two inhomogeneity maps (Figs.

469   8a and 8b) highlighted the locations of greatest differences between two maps, mainly along the

470   hydrographic network.

471

472   When comparing disaggregated soil maps derived from modified DSMART algorithm (DSMART

473   with soil landscape rules and DSMART with supplement soil observations), we note that the

474   DSMART with extra soil observations map delineated the study area into 132,942 regions. For

475   both maps, internal disaggregation was well pronounced expect for DSMART with extra soil

476   observations map in the southern part of the study area. Visual inspection of selected maps showed

477   high spatial agreement and highlighted some locations of greatest differences, particularly in the

478   southern part of the Ille et Vilaine department. Even if the hydrographic network was well detailed

479   in both maps, it appeared more developed in DSMART with extra soil observations soil map.

480   Applying *V measure* method for assessing the spatial similarity between DSMART with soil

481   landscape rules map and DSMART with supplement soil observations map provided similar

482   information theoretical measures $h_2 = 0.47$, $c_2 = 0.48$, and $V_2 = 0.47$. Visual comparison of soil

483   inhomogeneity maps revealed constant variance measured by normalized Shannon entropy. This

484   was in agreement with the quantitative assessment $c = h$. Overall, the two disaggregated maps were

485   spatially correlated, as indicated by the global spatial agreement measure V_2.

486

487   4) Discussion

488

489   4.1) Performance of the disaggregation procedures

490

491   Produced disaggregated soil maps closely resemble the abundant soils in the original soil map

492   (Holmes et al., 2015; Fig.3). The 1st most probable STU map derived from DSMART based

493   approaches captured the main spatial pattern of soil distribution across the study area. More internal

494   variation within SMUs was found when using DSMART with added point observations and

495   DSMART with soil landscape relationships. Local soil heterogeneity reflecting inherent

496   pedological complexity was depicted by the 1st STU maps which deliver a deterministic soil

497   landscape distribution, continuously varying with landscape features.

498   External validation was performed to assess the quality of disaggregated soil maps. Using 135

499   independent soil profiles and a per pixel validation approach, the overall accuracy reached 18.1%

500   for DSMART algorithm 1st STU map, 19.8% for DSMART with expert rules 1st STU map and

501   16.9% for DSMART with extra soil profiles 1st STU map. In the DSM literature, researchers who

502   applied classification tree decision methods founded similar validation results. For instance, by

503   applying DSMART algorithm in eastern Australia and using 285 legacy soil profiles, Odgers et al.

504   (2014) achieved an overall accuracy of 23%. Similarly, Nauman and Thompson (2014) explored

505   the use of expert rules for soil landscape relationships in the United States and achieved global

506   accuracy ranged between 22% and 24%. Similar disaggregation performance was recorded by

507   Holmes et al. (2015) in Western Australia (20%), Chaney et al. (2016) in the United States (17%)

508   and Møller et al. (2019) in Denmark (18%) using DSMART algorithm (Table 4). In contrast to the

509   latter studies, a large number of STU (171 STU) compose our soil dataset. This could certainly

510   decrease the chance of predicting the right STU, even through mobilizing relevant geographic

511   dataset to implement soil landscape relationships.

512

513   When considering a window of 3x3 pixels, the overall accuracy increased considerably for the

514   three DSMART based approaches maps, but DSMART with expert soil landscape relationships

515   achieved the highest accuracy scores. Chaney et al. (2016) highlighted a high degree of spatial

516   noise in the predictions by including pixel validation neighbours. Overall, prediction accuracy

517   increased twofold with a 3x3 pixel validation window and when grouping soils to a coarser level

518   of soil classification (171 versus 89 soil group). This was recorded for all disaggregated maps

519   regardless of the disaggregation procedure and suggests that fine soil taxonomic dissimilarities can

520   not be accurately mapped by disaggregation processes.

521

522   4.2) Legacy soil data

523

524   Legacy soil data used in this study provide an overall representation of soil over large areas (1:

525   250,000 scale). This database was derived from several soil surveys and pedological expert

526   knowledge. SMUs were spatially delineated, and their spatial organisation, as well as STUs

527   features, were described according to available soil data and pedological expertise. STUs and their

528   associated landscape characteristics were identified as accurately as possible using legacy soil

529   profiles collected according to a not probabilistic sampling design between 1968 and 2012. Hence,

530   differences in survey methods covering a large area over a long sampling period could lead to

531   errors in the STU definition or uncertainties in the estimation of their area in a given SMU.

532   Moreover, soil survey intensity was not uniform within SMUs. Thus, SMU components may be

533   derived from the unequal representation of soil samples across SMUs.

534   Harmonising soil data to reduce the number of STU is a great challenge by itself. Grouping some

535   STUs regarding their pedological similarities such as sharing comparable morphological criteria,

536   having similar pedogenic horizons and occurring in analogous environmental conditions is

537   worthwhile to be investigated. More importantly, unifying soil data according to more functional

538   aspects such as soil agricultural potential allows also to generate a relevant regional soil database

539   easily handled by soil users to satisfy their needs. Many countries around the world have already

540   harmonized their soil databases such as Denmark and Australia, where high pedological

541   complexity was captured with a reasonable STU number, with not exceeding 23 soil groups in

542   Denmark (Møller et al., 2019) and 73 soil groups in Australia (Holmes et al., 2015).

543

544   4.3) Taxonomic similarities

545

546   In the recent DSM literature, DSMART approach is considered as an efficient tool to disaggregate

547   existing coarse soil maps. In this study, we compared variants of the DSMART based approach,

548   which differed by the training dataset used to calibrate the C5.0 model and the allocation procedure.

549   Modified DSMART algorithms used additional calibration datasets derived from supplement soil

550   observations and expert sampling of polygons. Hence, taxonomic similarities were not taken into

551   account neither in the calibration process nor in the current component assignment scheme. Even

552   if there is a large number of STUs addressing inherent soil landscape heterogeneity, there is most

553   likely a short taxonomic distance between many of them. As a result, these STUs may have similar

554   forming conditions, making it a challenge to suitably constrain the prediction probabilities using

555   DSMART algorithm. This likely explains the high confusion index scores recorded in the present

556   study, particularly for original DSMART and DSMART with extra soil profiles approaches. As

557   demonstrated by Minasny and McBratney (2007), including taxonomic distance in decision trees

558   using pedological knowledge is a relevant way to decrease the misclassification error. Therefore,

559   future effort and improvements of the DSMART algorithm should take into account the taxonomic

560   distance between STU in the disaggregation procedure.

561

562   4.4) Mapping comparison

563

564   A quantitative comparison between disaggregated soil maps was performed using a novel approach

565   called *V measure* method. This method was commonly used to assess the spatial agreement

566   between land cover maps and thematic biotic and abiotic factors maps, as done by Nowosad and

567   Stepinski (2018) in the United States, but never before for soil maps.

568   In the present study, V_1 (0.53) was larger than V_2 (0.47) suggesting that DSMART with expert soil

569   landscape relationships map is much more similar to Original DSMART map than DSMART with

570   extra soil observations map. This might be explained by the allocation procedure for training

571   samples. The original DSMART algorithm tends to promote most abundant STUs with high

572   proportions of occurrence within polygons and penalized STUs with low proportions (comprise

573   between 2 and 10%). Therefore, frequent STUs are more likely to be predicted rather than rare

574   STUs. Meanwhile, by adding supplement soil profiles, preliminarily assigned to a suitable STU to

575   the training dataset, we constrain STUs with low proportions of occurrence predictions.

576   Major differences between DSMART with expert rules map and DMSART with soil observations

577   were mainly observed in the southern part of the study area and valleys areas. In general, Fluvisol

578   Stagnic soils were overestimated by DSMART with extra soil observations. This was likely due to

579   the purposive sampling design followed to supplement soil observations. The 755 legacy soil

580   profiles were selected to characterize hydromorphic soil conditions and to characterize inherent

581   soil landscape variability supposed to be organized along the hillslope.

582

583   4.5) Improvements and future work

584

585   Even though this work emphasizes the contribution of pedological knowledge in the disaggregation

586   process, other pathways can also be explored to improve map's accuracy. As recommended by

587   Mulder et al. (2016), compensating the temporal changes and differences in laboratory analytics is

588   a good option to improve the quality of legacy soil data. This suggests harmonising local soil

589   database and regrouping some STUs with similar soil forming factors through statistical modelling.

590   Moreover, additional environmental covariates with high spatial resolution should be used to

591   capture micro landscape variability (Lacoste et al., 2014; Odgers et al., 2014; Chaney et al., 2016;

592   Møller et al., 2019). For example, adding a more detailed Digital Elevation Model allowed to

593   capture small terrain features, where may be particular, STUs occurs. Improving both polygon

594 sampling procedure and current components assignment scheme turned out to be important to
595 reduce uncertainty prediction. This suggests drawing virtual soil samples proportionally to
596 polygons areas and using supplement STU characteristics based on surveyor observations (slope
597 shape, hillslope position, soil texture …) to guide STU allocation procedure (Møller et al., 2019).
598 Assuming that the decision tree can be built to relate STU descriptors to legacy soil data, this
599 method can replace weighted random allocation procedure and should help minor STU prediction
600 by constraining raster probabilities.

601 5) Conclusion

602

603 We applied three DSMART based approaches, including original DSMART algorithm, DSMART
604 with extra soil observations and DSMART with soil landscape relationships, to disaggregate legacy
605 soil polygons over a large area in Brittany (France). Regardless of the disaggregation approach, the
606 produced soil maps at 50 m spatial resolution successfully address the main soil spatial pattern
607 regarding prior pedological knowledge of our study area. Performance assessed against 135
608 independent soil profiles, 755 legacy soil profiles, and accurate 1:25,000 soil maps highlighted that
609 DSMART with expert rules maps achieved highest validation measures. Overall, modified
610 DSMART algorithms allowed minor STUs prediction, whereas original DSMART algorithm
611 promoted abundant STUs prediction with poor spatial structure improvement. Adding pedological
612 knowledge as well as extra soil observations in the prediction process constrained STU
613 probabilities, even STUs with low proportions. However, some particular STUs reflecting
614 hydromorphic soils or loamy soils were greatly overestimated for all the three DSMART based
615 approaches.
616 Soil maps produced using the original DSMART and DSMART with expert rules have a high
617 spatial agreement, but the latter map appeared more detailed and provided a spatially continuous
618 and consistent STU's prediction. Therefore, generalizing soil landscape relationships taken to
619 account several STU descriptors and landscape features should be implemented in the future
620 version of DSMART algorithm to capture soil landscape heterogeneity and consequently guarantee
621 coherent variability of soil properties.

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Figure captions

Figure 1: Location of the study area and the validation datasets

Figure 2: Schematic of the DSMART based approaches algorithm. The steps in DSMART are: 1) construct the calibration dataset; 2) train C5.0 model;   3) estimate STU maps and their associated probabilities of occurrence

Figure 3: Digital soil map of the most probable STU and their associated probability of occurrence for the whole study area and for a focus zone, a) Legacy soil map: most probable STU for each SMU, b) original DSMART approach; c) DSMART with expert rules; d) DSMART with extra soil observations

Figure 4: Global probability of hydromorphic soils over the study area derived from a) original DSMART, b) DSMART with soil landscape relationships and c) DSMART with extra soil observations. The probabilities of the three STU with highest prediction occurrence are summed if they are hydromorphic

Figure 5: Confusion index maps for a) Original DSMART approach; b) DSMART with expert rules; c) DSMART with extra soil observations

Figure 6: Cumulative area of the 171 STUs estimated from the regional soil database and predicted by different DSMART based approaches

Figure 7: Violin plots of the relative importance of each environmental covariate used in a) Original DSMART approach; b) DSMART with expert rules; c) DSMART with extra soil observations

Figure 8: Spatial association between disaggregated maps of Ille et Vilaine department. a) map of inhomogeneity of DSMART with soil landscape relationships map in terms of original DSMART map b) map of inhomogeneity of original DSMART map in terms of DSMART with soil landscape relationships map c) map of inhomogeneity of DSMART with soil landscape relationships map in terms of DSMART with extra soil observations map d) map of inhomogeneity of DSMART with extra soil observations map in terms of DSMART with soil landscape relationships map. Inhomogeneity (variance) is measured by normalised Shannon entropy

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684 **Table headings**

685

686 Table 1. Description of the environmental covariates selected. Summary of environmental
687 covariates. P: parent material; S: soil properties; R: relief; O: Organisms; C: categorical; Q:
688 quantitative.

689 Table 2. Ten most extended STUs according to the regional soil database and their respective rank
690 by area using three DSMART based disaggregation procedures

691 Table 3. Overall accuracies (%) obtained using various external validation approaches for the three
692 most probable STU

693 Table 4: Comparison between the size areas covered, number of soil map units, soil type units of
694 the original legacy soil maps and the accuracy achieved in other studies using DSMART algorithm

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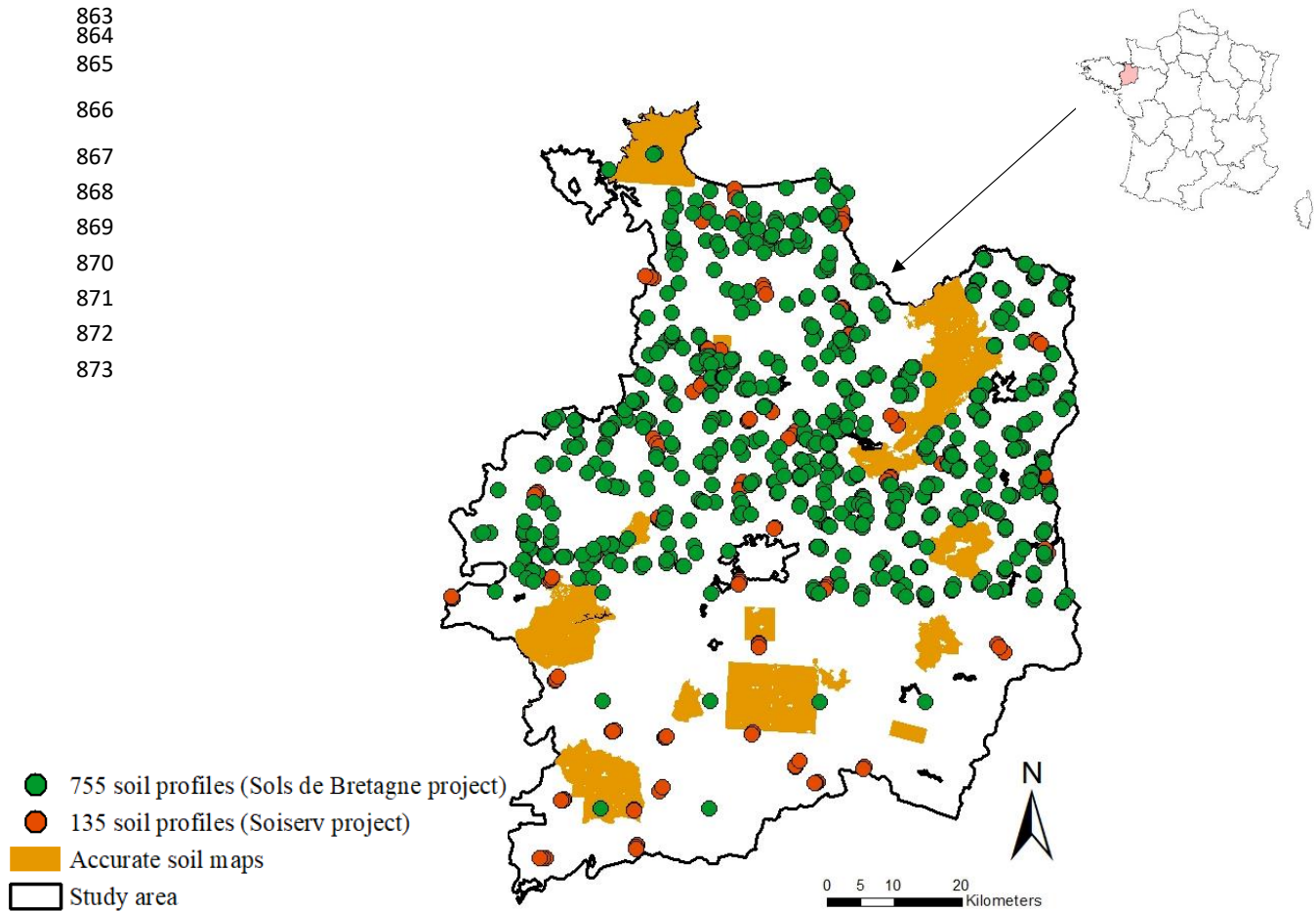862

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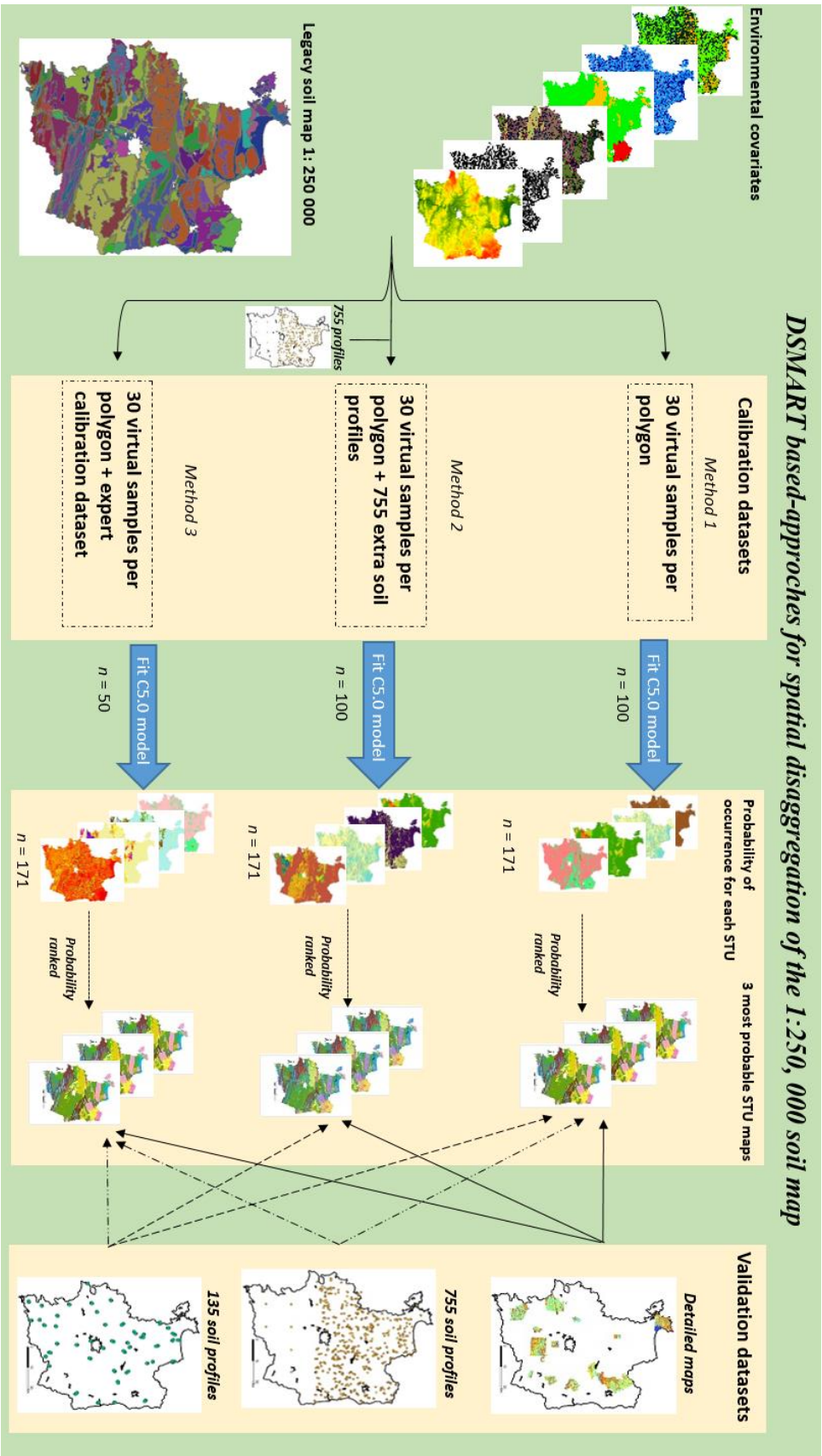*Figure 1: Location of the study area and the validation datasets*

Figure 2: Schematic of the DSMART based approaches algorithm. The steps in DSMART are: 1) construct the calibration dataset; 2) train C5.0 model; 3) estimate STU maps and their associated probabilities of occurrence
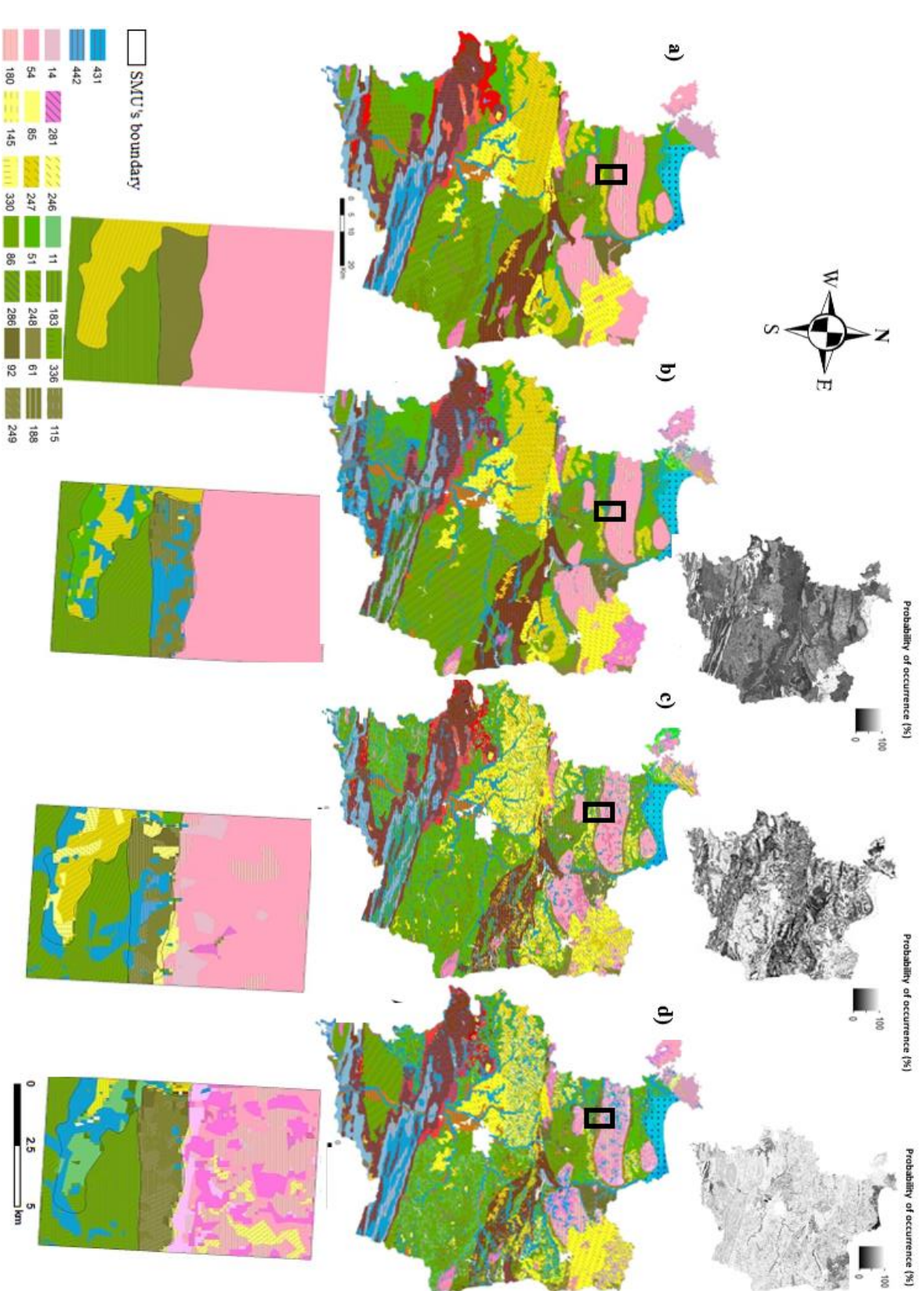
Figure 3: Digital soil map of the most probable STU and their associated probability of occurrence for the whole study area and for a focus zone, a) Legacy soil map: most probable STU for each SMU, b) original DSMART approach; c) DSMART with expert rules; d) DSMART with extra soil observations
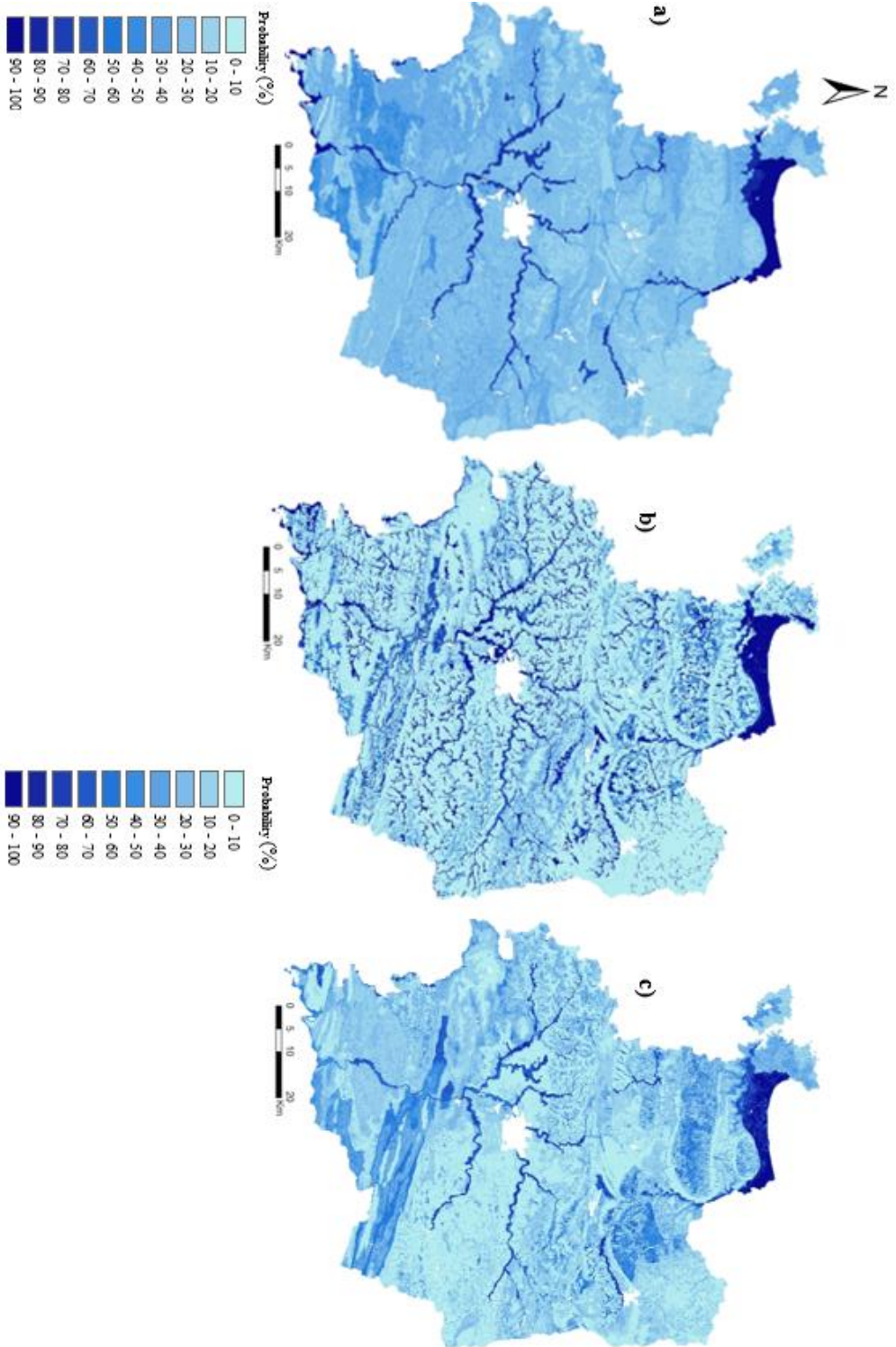
Figure 4: Global probability of hydromorphic soils over the study area derived from a) original DSMART, b) DSMART with soil landscape relationships and c) DSMART with extra soil observations. The probabilities of the three STU with highest prediction occurrence are summed if they are hydromorphic.
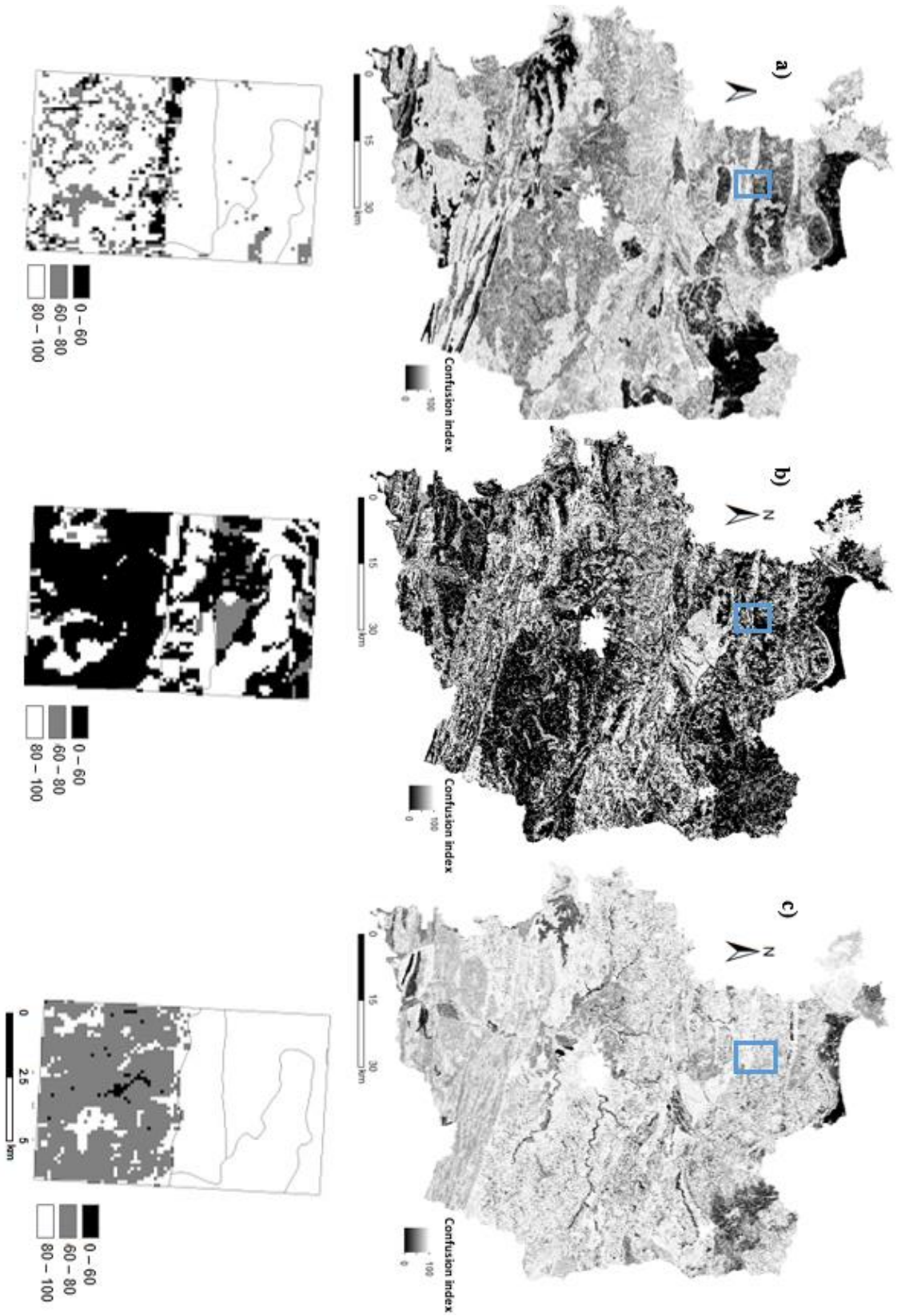
Figure 5 Confusion index maps for a) Classic DSMART approach; b) DSMART with expert rules; c) DSMART with extra soil observations
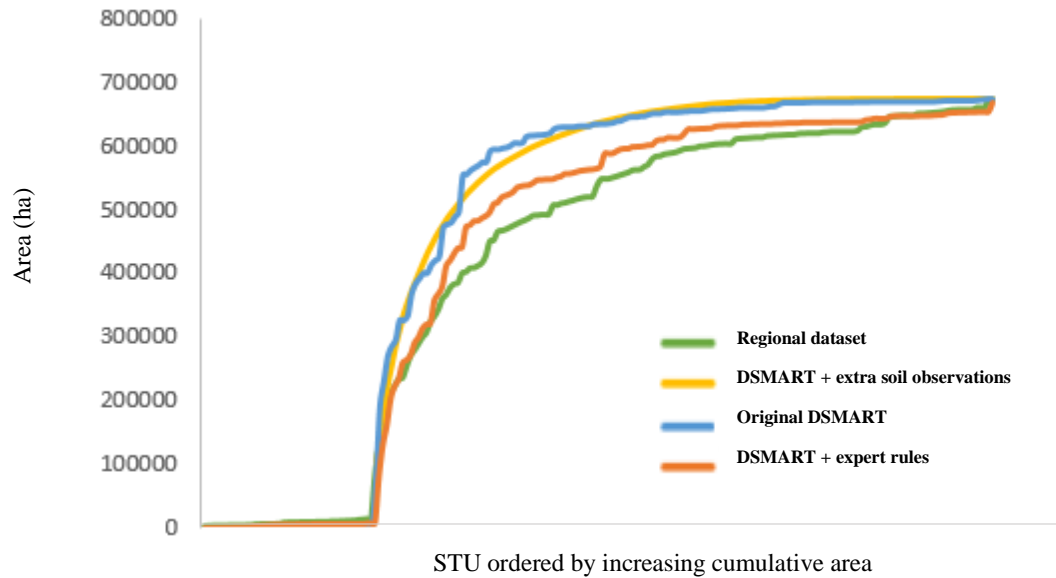
Figure 6: Cumulative area of the 171 STUs estimated from the regional soil database and predicted by different DSMART based approaches
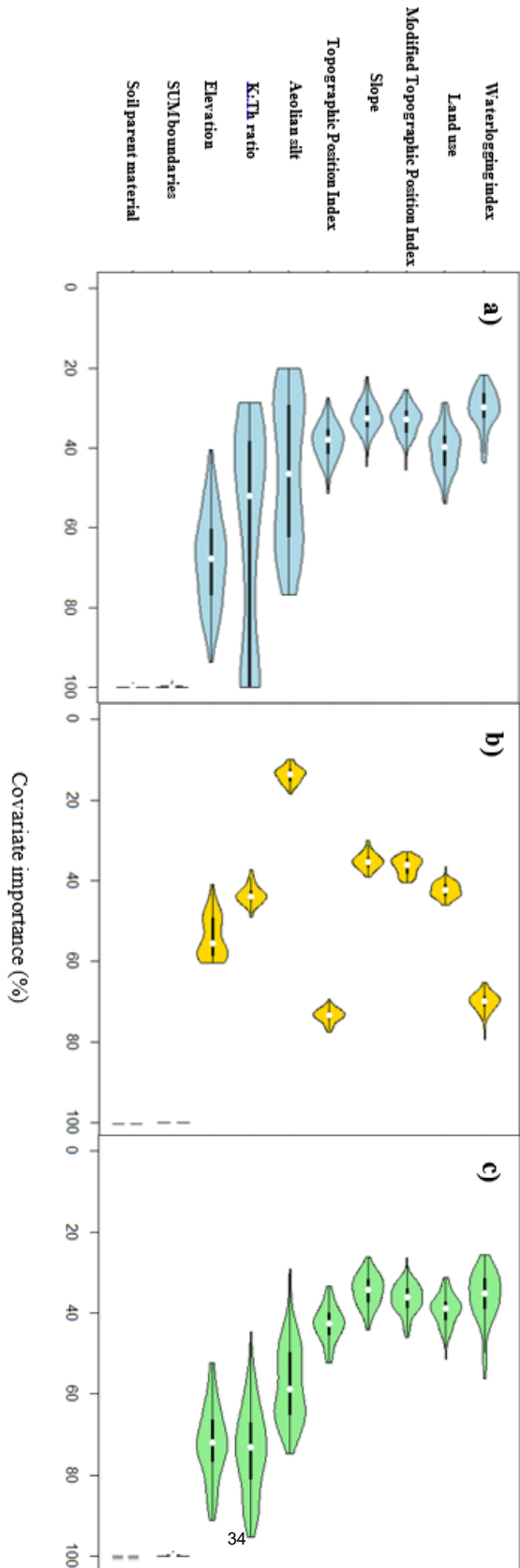
Figure 7: Violin plots of the relative importance of each environmental covariate used in a) Original DSMART approach; b) DSMART with expert rules; c) DSMART with extra soil observations

Figure 8: Spatial association between disaggregated maps of Ille et Vilaine department. a) map of inhomogeneity of DSMART with soil landscape relationships map in terms of original DSMART map b) map of inhomogeneity of original DSMART map in terms of DSMART with soil landscape relationships map c) map of inhomogeneity of DSMART with soil landscape relationships map in terms of DSMART with extra soil observations map d) map of inhomogeneity of DSMART with extra soil observations map in terms of DSMART with soil landscape relationships map. Inhomogeneity (variance) is measured by normalised Shannon entropy

Table 2 Ten most extended STUs according to the regional soil database and their respective rank by area using three DSMART based disaggregation procedures

STU Label	WRB classification	Parent material	1:250,000 dataset		Original DSMART approach		DSMART with extra soil profiles		DSMART with expert rules	
			Rank	Estimated area (km²)	Rank	Predicted area (km²)	Rank	Predicted area (km²)	Rank	Predicted area (km²)
431	Fluvisol Stagnic	Alluvial and colluvial deposits	1	688	2	757	1	983	1	740
248	Cambisol	Brioverian schists	2	480	1	1154	2	461	2	492
51	Cambisol	Brioverian schists	3	402	5	397	4	395	3	424
61	Cambisol	Gritty schists	4	227	9	177	30	53	14	128
183	Cambisol Stagnic	Sandstone	5	216	11	162	5	308	10	192
256	Cambisol	Aeolian loam	6	200	6	385	3	418	6	314
286	Cambisol Stagnic	Brioverian schists	7	179	23	62	9	187	24	80
86	Cambisol	Brioverian schists	8	169	12	126	15	124	4	358
340	Albeluvisol Stagnic	Granite and gneiss	9	168	7	347	10	177	11	189
54	Cambisol	Brioverian schists	10	167	4	451	18	98	5	324

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Table 1. Description of the environmental covariates selected

Summary of environmental covariates. P: parent material; S: soil properties; R: relief; O: Organisms; C: categorical; Q: quantitative.

Environmental covariate	SCORPAN factor	Type	Unit or number of classes
Terrain attributes derived from the digital elevation model			
Elevation	R	Q	m
Slope	R	Q	%
Compound Topographic Index (TPI)	R	Q	Log (m^3)
Topographic Position Index	R	C	5 classes
Pedology and geology			
Soil parent material	P	C	22 classes
Soil Map Units	R	C	96 classes
Aeolian silt deposits	P	C	2 classes
Waterlogging index	S	C	4 classes
Organism			
Landscape units	O	C	19 classes
Gamma ray spectrometry from 250 m airborne geophysical survey interpolations			
K:Th ratio	P	Q	

Table 3. Overall accuracies (%) obtained using various external validation approaches for the three most probable STU

Pixel to pixel validation of STU					
	DSMART approach	Most probable STU	Second most probable STU	Third most probable STU	Total
	Original DSMART	23	13	8	44
	DSMART with expert rules	19	11	7	37
Soil maps (87 150 ha)	DSMART with extra soil observations	22	9	7	38
	Original DSMART	11	5	3.8	18.1
	DSMART with expert rules	10	4.4	3.7	19.8
Independent soil profiles (n=135)	DSMART with extra soil observations	8.2	6	2.7	16.9
	Original DSMART	14	7	6	27
	DSMART with expert rules	18	9	7	34
Legacy soil profiles (n=755)	DSMART with extra soil observations				

Pixel to pixel validation of STU group					
	DSMART approach	Most probable STU	Second most probable STU	Third most probable STU	Total
Soil maps (87 150 ha)	Original DSMART	26	13	9	48
	DSMART with expert rules	22.5	13.7	9.7	45.9
	DSMART with extra soil observations	25	10	7	42
Independent soil profiles (n=135)	Original DSMART	16	7	4.6	27.6
	DSMART with expert rules	18	8.4	5.2	31.6
	DSMART with extra soil observations	15	8	3.8	26.8
Legacy soil profiles (n=755)	Original DSMART	19	12	9	40
	DSMART with expert rules	23.4	15	11.8	50.2
	DSMART with extra soil observations				

Neighbourhood of 3 x 3 validation of STU					
	DSMART approach	Most probable STU	Second most probable STU	Third most probable STU	Total
Soil maps (87 150 ha)	Original DSMART	31	16	14	61
	DSMART with expert rules	29.6	19.4	13.1	62.1
	DSMART with extra soil observations	28	11	9	48
Independent soil profiles (n=135)	Original DSMART	15	6	4.3	25.3
	DSMART with expert rules	17	6.7	4.8	28.5
	DSMART with extra soil observations	11	7	3	21
Legacy soil profiles (n=755)	Original DSMART	19	10	7	36
	DSMART with expert rules	27.9	15	11.9	54.8
	DSMART with extra soil observations				

Table 4: Comparison between the size areas covered, number of soil map units, soil type units of the original legacy soil maps and the accuracy achieved in other studies using DSMART algorithm

Study	Area (km²)	Map units	Soil type unit	Accuracy
Odgers et al (2014)	68,000	1,110	72	23
Holmes et al. (2015)	2,500,000	5,069	73	20-22
Chaney et al. (2016)	-	-	-	17
Møller et al. (2019)	43,000	11-14	18-23	12-18