

1 **Opportunities and limitations related to the application of** 2 **plant-derived lipid molecular proxies in soil science**

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4 **Boris Jansen¹ and Guido L. B. Wiesenberg²**

5 [1]{Institute for Biodiversity and Ecosystem Dynamics, University of Amsterdam,
6 Amsterdam, P.O. Box 94240, NL-1090GE, The Netherlands}

7 [2]{Department of Geography, University of Zürich, Winterthurerstrasse 190, CH-8057
8 Zürich, Switzerland}

9 *Correspondence to:* B. Jansen (B.Jansen@uva.nl)

10 11 **Abstract**

12 The application of lipids in soils as molecular proxies, also often referred to as biomarkers,
13 has dramatically increased in the last decades. Applications range from inferring changes in
14 past vegetation composition, climate and/or human presence to unraveling input and turnover
15 of soil organic matter (SOM). Molecules used include extractable and non-extractable lipids,
16 including ester-bound lipids. In addition, the carbon or hydrogen isotopic composition of
17 such molecules is used. While holding great promise, the application of soil lipids as
18 molecular proxies comes with several constraining factors the most important of which are: i)
19 variability in the molecular composition of plant-derived organic matter plant-internally and
20 in between plant individuals; ii) variability in (relative contribution of) input pathways into
21 the soil; and iii) transformation and/or (selective) degradation of (some of) the molecules
22 once present in the soil. Unfortunately, the information about such constraining factors and
23 their impact on the applicability of molecular proxies is fragmented and scattered. The
24 purpose of this study is to provide a critical review of the current state of knowledge with
25 respect to the applicability of molecular proxies in soil science, specifically focusing on the
26 factors constraining such applicability. Variability in genetic, ontogenetic and environmental
27 factors influence plant *n*-alkane patterns in the way that no unique compounds or specific
28 molecular proxies pointing to e.g. plant-community differences or environmental influences,
29 exist. Other components such as *n*-alcohols, *n*-fatty acids, cutin- and suberin-derived
30 monomers have received far less attention in this respect. Furthermore, there is a high

1 diversity of input pathways offering both opportunities and limitations for the use of
2 molecular proxies at the same time. New modelling approaches might offer a possibility to
3 unravel such mixed input signals. Finally, transformation and turnover of SOM offer
4 opportunities when tracing such processes is the purpose of applying a molecular proxy,
5 whilst posing limitations when they obliterate molecular proxy signals linked to other
6 phenomena. For *n*-alkanes several modelling approaches have recently been developed to
7 compensate for (selective) degradation. Still such techniques are in their infancy and
8 information about their applicability to other classes of components than *n*-alkanes is lacking
9 yet. All constraining factors considered can have a significant influence on the applicability
10 of molecular proxies in soil science. The degree of influence strongly depends on the type of
11 molecular proxy as well as the environmental context in which it is applied. However, the
12 potential impact of the constraining factors should always explicitly be addressed whenever
13 molecular proxies are applied in a soil scientific context. More importantly, there is still a
14 serious lack of available information in particular for compound classes other than the *n*-
15 alkanes. Therefore, we urgently call for the consideration of more holistic approaches
16 determining various factors during sampling as well as using as many compound classes as
17 possible.

18

19 **1 Introduction**

20 Since more than a century, various compounds deriving from the substance class of lipids,
21 have been investigated in plant and soil science. Some of the earliest publications in plant
22 science date back to the first half of the 19th century (Liebig et al., 1837; Wöhler F. and
23 Liebig, 1839) and in soil science to the early 20th century as already reviewed by Stevenson
24 (1966). One of the main interests to study lipids was the large heterogeneity of compounds
25 included in this substance class. Some of the individual compounds have been described as
26 ‘biomarkers’ or ‘biogenic markers’, i.e. compounds that “*may be diagnostic of specific*
27 *organisms, classes of organism, or general biota that contribute organic matter to the*
28 *atmosphere, aqueous or sedimentary environment*” (Peters et al., 2005). In addition,, in
29 environmental sciences also anthropogenic and petroleum markers were highlighted that have
30 the ability to be preserved with “no or only minor change” (Tissot and Welte, 1984; Peters et
31 al., 2005). Because *sensu strictu* the term biomarker has been used for the differentiation of
32 biological tissues of different origin in environmental matrices, recently the term ‘molecular
33 proxy’ has come in vogue. This term allows for an inclusion of biomarkers *sensu strictu* as

1 individual compounds characterizing specific biogenic sources, but also individual
2 compounds acting as specific proxy e.g. for anthropogenic impact or thermal alteration.
3 Furthermore, it accommodates the use of groups of compounds used in the before mentioned
4 way. Finally, it implies the use of molecular ratios of compounds like the carbon preference
5 index (CPI) or the average chain length (ACL) that could also be indicative for biogenic
6 sources, alteration or overprint of organic matter. Therefore, in the present work we use the
7 term molecular proxy rather than biomarker.

8 In its broadest sense, molecular proxies allow determination of the presence, absence, or
9 certain characteristics of a (set of) molecule(s) that are indicative for a process in, or state or
10 composition of a system of interest. For instance, in the clinical sciences molecular proxies
11 among other applications are used as indicators of the presence of a disease or response to
12 treatment (Brennan et al., 2013; Van Bon et al., 2014); in toxicology to assess the effect of
13 toxicant exposure on biota (Clemente et al., 2014); in the forensic sciences to link suspects to
14 a crime scene (Concheri et al., 2011); in limnology to examine past lacustrine environmental
15 conditions (Castañeda and Schouten, 2011); and in organic geochemistry to follow oil
16 formation and translocation in source and reservoir rocks (Curiale, 2002).

17 Also in soil science, molecular proxies have been used for decades, and their application has
18 exponentially increased in the last decade as indicated by the number of related articles
19 published in Web of Science indexed journals (Table 1). The types of molecular proxies used
20 are as diverse as the field of soil science itself. They range from the use of phospholipid fatty
21 acids to estimate bacterial and fungal biomass in soils (Frostegard and Bååth, 1996), to the
22 application of preserved retene/caldalene ratios to infer palaeoecological vegetation shifts
23 (Hautevelle et al., 2006). Also the archives of the molecular proxies in soil sciences that are
24 used are diverse and, in addition to soils themselves, include lacustrine and terrestrial
25 sediments, peat deposits, as well as paleosols (Zhang et al., 2006; Bai et al., 2009; Andersson
26 et al., 2011; Berke et al., 2012). However, in spite of this large variety a limited number of
27 scientific topics can be discerned that encompass the great majority of molecular proxy
28 application in the soil sciences. These are:

- 29 • Changes in vegetation composition inferred from extractable and/or ester-bound lipids
30 of plant origin, and/or their carbon isotopic composition (e.g. Huang et al., 1996; Zech
31 et al., 2009; Le Milbeau et al., 2013).

- 1 • Changes in climate, i.e. mean annual temperature and/or precipitation inferred from
2 bacterial membrane lipids and/or the hydrogen isotopic composition of plant-derived
3 lipids (e.g. Weijers et al., 2006; Krull et al., 2006; Rao et al., 2009).
- 4 • Changes in palaeoelevation inferred from bacterial membrane lipids and/or the
5 hydrogen isotopic composition of plant-derived lipids (e.g. Sachse et al., 2006; Bai et
6 al., 2011; Ernst et al., 2013).
- 7 • Changes in human impact or settlement inferred from compound-specific N isotope
8 analysis or transformation products of plant-derived lipids, e.g. through burning, or
9 manure derived lipids (e.g. Bull et al., 1999; Eckmeier and Wiesenberg, 2009;
10 Zocatelli et al., 2012).
- 11 • Contribution of fossil fuel-derived carbon to soil assessed by lipid molecular
12 composition and compound-specific isotopes (e.g. Lichtfouse et al., 1995; Lichtfouse
13 et al., 1997; Rethemeyer et al., 2004).
- 14 • Input, transformation and/or decomposition of soil organic matter inferred from or
15 traced through extractable and/or ester-bound lipids of plant origin and/or bacterial
16 membrane lipids and/or their carbon isotopic composition. (e.g. Nierop et al., 2001;
17 Amelung et al., 2008; Hamer et al., 2012).

18 In Table 1 an overview is given of the classes of molecules frequently used as molecular
19 proxies in soil archives in relation to their application as well as total and recent (last ten
20 years) publications including the respective keywords.

21 When using molecular proxies to answer research questions in any of the areas identified, in
22 particular when soils are used as an archive, several constraining factors have to be taken into
23 account that vary with the type of application and research question to be answered. The most
24 important ones are:

- 25 i) Variability in the source of plant-derived organic matter, i.e. abundance and
26 composition of the molecular proxies in different plant species, plant specimens and
27 plant parts as a result of genetic or life stage variations and/or external factors such as
28 climate, seasonality or exposure to the sun (e.g. Nødskov Giese, 1975; Lockheart et
29 al., 1998; Shepherd and Griffiths, 2006).
- 30 ii) Variability in (relative contribution of) input pathways into the soil, in particular
31 microbial versus vegetation input, and root versus aboveground biomass input (e.g.
32 Jackson et al., 1996; Schefuß et al., 2003; Mambelli et al., 2011).

1 iii) Transformation and/or (selective) degradation of (some of) the compounds once
2 present in the soil, when it is not the aim of the study to use the molecular proxies to
3 study such transformations (e.g. De Leeuw and Baas, 1986; Nguyen Tu et al., 2004;
4 Andreetta et al., 2013).

5 However, the information about such constraining factors and their impact on the
6 applicability of molecular proxies is fragmented and scattered over different publications
7 inside and outside the scientific discipline of soil sciences. For instance, much of the
8 available information about variation of leaf wax lipid composition is presented in the plant
9 physiological literature in studies that were not conducted with the application of such lipids
10 as molecular proxy for past vegetation composition from soil archives in mind (e.g. Tulloch,
11 1973; Avato et al., 1984; Kim et al., 2007). The fragmentation of the information makes it
12 difficult for researchers to assess the potential influence of constraining factors on the
13 application of molecular proxies. It also hinders the identification of hiatuses in the available
14 knowledge about the constraining factors as well as the designation of potential strategies to
15 compensate or correct for such constraints.

16 Therefore, the purpose of the present study is to provide a critical review of the current state
17 of knowledge with respect to the applicability of molecular proxies in soil science,
18 specifically focusing on the factors constraining such applicability. Based on this we will
19 identify areas for future research both with respect to the application of molecular proxies in
20 soil science as well as the constraints thereof.

21 The vastness of the field of molecular proxies forced us to restrict the scope of the present
22 study. With respect to the molecules to consider, a first restriction was to focus on those
23 related to the earlier mentioned main areas of application of molecular proxies in soil science.
24 A second restriction was to focus on the main classes of components as used by several
25 researchers. Finally, in spite of their common application, we explicitly excluded lignin and
26 phospholipid fatty acids (PLFA) as lignin was subject of another recent review article
27 (Thevenot et al., 2010) and PLFAs are considered in such a large set of studies (c.f. Table 1)
28 that they would require a separate review. Finally GDGTs were excluded because their
29 application is predominantly in aquatic sediments rather than soils and they have been
30 recently reviewed (Schouten et al., 2013). This leaves the component classes labeled in bold
31 in Table 1 to be considered in the present study. Our study is relevant to the application of
32 compound-specific isotope analysis inasmuch that such analysis is directly affected by
33 variability and transformation of the underlying molecules. However, we did not explicitly

1 consider sources and effects of variation of the stable $\delta^{13}\text{C}$ and $\delta^2\text{H}$ isotope signature of
2 specific molecules themselves, this being a research area of its own and also subject of other
3 recent reviews by Sachse et al. (2012) and Diefendorf and Freimuth (2017). Furthermore,
4 when considering application and preservation of molecular proxies we restricted ourselves
5 to topsoils (i.e. surface soil horizons = A horizons as defined by the FAO in the Guidelines
6 for soil description (2006)) as archives.

7

8 **2 Source related variability of molecular proxies**

9 **2.1 Definition**

10 Source related variability of molecular proxies pertains to intra-species variation in the
11 abundance of the molecules that are used as proxy. Such variability entails: i) variation in
12 relative abundance of individual compounds that together constitute the proxy, e.g. of *n*-
13 alkanes of different chain length in leaf waxes of a certain species; ii) variation in absolute
14 abundance of the molecules used as proxy either between different specimens or between
15 different parts of a single specimen. Depending on the research question, intra-species
16 variability of molecular proxies may be desirable or not. For instance when preserved leaf
17 wax lipids patterns are used to reconstruct past vegetation composition, the implicit
18 assumption is that the intra-species variability in the source vegetation is small compared to
19 the inter-species variability.

20 There are two main causes of intra-species variability in molecular proxies: internal variation
21 related to genetics and/or ontogeny; and external variation related to the growing
22 environment. Both are related in the sense that differences in response to environmental
23 factors are also often genetically determined (Shepherd and Griffiths, 2006). Here we discuss
24 both causes separately with a third paragraph devoted to studies where combined effects were
25 examined. For a detailed description of the biomolecular mechanisms of wax genesis and all
26 potential sources of change, the reader is referred to the review provided by Shepherd and
27 Griffiths (2006).

2.2 Variation related to genetics and/or ontogeny

2.2.1 Wax lipids

Many studies have indicated that the clear genetic control of leaf wax genesis leads to a significant and meaningful difference in their composition (Shepherd et al., 1995; Shepherd and Griffiths, 2006). For instance, prompted by the early works in this area (e.g. Eglinton et al., 1962; Herbin and Robins, 1968; Herbin and Robins, 1969), Maffei performed an extensive evaluation of the *n*-alkane patterns in several hundreds of plant species belonging to the Poaceae, Apiaceae, Brassicaceae, Fabaceae, Cactaceae, Pinaceae, Lamiaceae, Boraginaceae, Verbenaceae, Solanaceae and Scrophulariaceae (Maffei, 1994; Maffei, 1996a; Maffei, 1996b; Maffei et al., 1997; Maffei et al., 2004). These studies were replenished by those on Styracaceae (Li et al., 2013), Moraceae (Sonibare et al., 2005), and Clusiaceae (Medina et al., 2004; Medina et al., 2006). Further, Dove et al. (1996) described the alkane diversity among a grassland plant community, which enables tracing of the diet of grazing animals due to the different alkane compositions of the plants. Recently, Mueller-Niggemann and Schwark (2015) were able to differentiate rice from alternating crop plants based on their *n*-alkane patterns. The results support the chemotaxonomic discriminatory power of *n*-alkane patterns at family, sub-family and tribal level, which has been further examined by Diefendorf et al. (2017). Examining plant *n*-alkane and *n*-alcohol distribution of 37 C_4 grasses, Rommerskirchen et al. (2006) also found chemotaxonomic differentiation was possible at the sub-family level. Mongrand et al. (2001) examined the fatty acid composition of the leaves of over 137 species of gymnosperms belonging to 14 families and collected from different locations in France. They found a taxonomically meaningful clustering into four main groups, with the highest discriminatory power in the Pinaceae at the genus level (Mongrand et al., 2001). Additionally, Wiesenberg and Schwark (2006) determined differences in the fatty acid composition between temperate C_3 - and C_4 -crops. Within the same *Brassica* species of kale and swede Shepherd et al. (1995) observed a difference in chain length distribution of wax lipids between two genotypes of the same species, indicative of genetic control through variation in the enzyme system. Also for the isoprenoids, a genetically driven discriminatory power related to (groups of) plant species is attributed (Ohsaki et al., 1999; Jansen et al., 2007).

The chemotaxonomic potential of wax lipids as just described has been exploited to reconstruct past vegetation history from wax lipids preserved in soil archives (e.g. Bull et al.,

1 1999; Zech et al., 2009; Jansen et al., 2013). Such reconstructions often focus on the use of
2 shifts in ratios of certain dominant higher chain length *n*-alkanes, fatty acids and *n*-alcohols
3 representative of a shift in dominant vegetation over time (Jansen et al., 2010; Gocke et al.,
4 2016; Wiesenberg et al., 2015). In few instances also the entire suite of higher chain length *n*-
5 alkanes and *n*-alcohols (Jansen et al., 2013) or *n*-alkanes and fatty acids have been used
6 (Wiesenberg et al., 2015).

7 However, in addition to other issues such as discussed in the other sections of this review
8 paper, an important issue when exploiting the chemotaxonomic potential of leaf wax lipids is
9 the phenotypic plasticity of the genetic variability in leaf wax lipid patterns found and the
10 implications thereof for the stability of the patterns observed. Maffei et al. (2004) concluded
11 that phenotypic plasticity may overcome genetic variability, particularly when plant
12 developmental stages are considered along with abiotic and biotic stress conditions. Several
13 plant physiological studies have focussed on wax lipid composition related to plant life stage,
14 and report different results. Avato et al. (1984) found that where the relative contribution of
15 *n*-fatty acids, *n*-alcohols and *n*-alkanes differed between *Sorghum* seedlings and mature
16 leaves, the chain-length distribution within a component class remained the same for the *n*-
17 alkanes and *n*-alcohols. Giese (1975) observed a difference in homologue dominance of *n*-
18 alkanes between leaves of seedlings and mature barley plants. Also Herbin and Robins
19 (1969), Dyson and Herbin (1970), Baker and Hunt (1981), and Zhang et al. (2004) identified
20 increasing chain length dominance of leaf wax alkanes with increasing leaf age. However,
21 averaging of sampling over leaves of different age, position etc. within a stand of trees did
22 allow for distinction from other stands, indicating that inter-species variation was larger than
23 intra-species variation (Dyson and Herbin, 1970). Baker and Hunt (1981) observed
24 differences between adaxial and abaxial parts of leaves for some of the plant species. Also
25 Tulloch (1973) observed a variation of leaf waxes of several *Triticum* species with age. In
26 particular the whole plant *n*-alkane predominance shifted from C₃₁ at 24 days after
27 germination to C₂₉ at 100 days after germination (Tulloch, 1973). Furthermore, Wiesenberg
28 et al. (2004; 2012) and Wiesenberg and Schwark (2006) observed changes in *n*-alkane and *n*-
29 fatty acid compositions of a variety of temperate crop species with plant age. Other
30 publications reported seasonal variations in the *n*-alkane composition for variety of pasture
31 and crop plants by Dove et al. (1996), Hellgren and Sandelius (2001), Moseley (1983),
32 Shelvey and Koziol (1986) and various trees especially by Gülz and collaborators (Prasad
33 and Gülz, 1990; Gülz et al., 1991; Gülz and Muller, 1992; Gülz and Boor, 1992). Variations

1 in the *n*-alkane composition could be observed during the growing season among all
2 investigated plants, but general trends of increasing or decreasing chain length and *n*-alkane
3 contents have not consistently been determined. The *n*-alcohol predominance also varied but
4 to a much smaller extent, not affecting the predominance of a specific *n*-alcohol (Tulloch,
5 1973). Esters gradually showed an increase in esters of trans 2,3-unsaturated C₂₃ and C₂₄
6 acids with plant age (Tulloch, 1973). The variation was related to the development of the
7 plant, in particular that of flag leaf and sheets between 55 and 66 days (Tulloch, 1973).

8 In contrast to the previous, Li et al. (1997) studied the influence of ontogeny on leaf wax
9 lipids (*n*-alkanes, *n*-aldehydes, *n*-alcohols, esters, β-diketones, flavonoids and triterpenoids)
10 in several *Eucalyptus* species of the subgenus *Symphyomyrtus* on Tasmania, and found no
11 significant effect of ontogeny on leaf wax composition, which they found to clearly and
12 consistently differ between species (Li et al., 1997). Also Eglinton et al. (1962) observed that
13 the *n*-alkane composition of leaf waxes of 74 species of Crassulaceae from the Canary Islands
14 showed no appreciable variation with respect to leaf position, age, size or specimen. Further,
15 Bush and McInerney (2013) found no influence of canopy position or sampling time on the
16 *n*-alkane patterns of mature leaves from 24 tree species.

17 **2.2.2 Cutin and suberin monomers**

18 Cutin and suberin monomers are mainly used as proxies to distinguish leaf from root input in
19 soils (Schreiber et al., 1999; Bull et al., 2000; Mendez-Millan et al., 2011) or as proxy for
20 related phenomena such as the degree of bioturbation in the topsoil (Nierop and Verstraten,
21 2004). Therefore, the possible (onto)genetic effects on cutin and suberin composition are a
22 concern if they were to alter the composition of the polyesters to such an extent that the
23 separation between cutin and suberin is compromised.

24 Some general observations in literature are that long-chain even numbered C₂₀-C₃₀ ω-hydroxy
25 fatty acids and α,ω-alkanedioic acids mainly originate from suberin, whereas shorter chained
26 C₁₆ and C₁₈ ω-hydroxy fatty acids mainly derive from cutin (Schreiber et al., 1999; Otto et
27 al., 2005; Mendez-Millan et al., 2011). However, several publications challenge the universal
28 applicability of such general observations, indicating instead that genetic variability results in
29 many exceptions to such general rules. For instance, Hamer et al. (2012) found that ωC_{22:0},
30 ωC_{24:0} and ωC_{26:0} hydroxy fatty acids were not exclusively associated to roots, but also
31 occurred in the shoots of several species. In addition, ωC_{16:0} and ωC_{18:0} fatty acids were not
32 exclusive to the leaves, but also occurred in the roots of several species.

1 **2.3 Variation related to environmental factors**

2 **2.3.1 Effects of temperature**

3 Increased solar radiation levels are generally reported to lead to higher absolute amounts of
4 waxes produced (Sanchez et al., 2001; Shepherd and Griffiths, 2006). In addition, the
5 composition of the various component classes of wax lipids, i.e. the relative contribution of
6 *n*-fatty acids, *n*-alkanes, *n*-alcohols etc., has been reported to change. A shift towards lower
7 chain lengths within different component classes was sometimes found (Shepherd and
8 Griffiths, 2006). Thus, a positive correlation of long-chain odd *n*-alkanes with temperature
9 was observed (Maffei et al., 1993; Zhang et al., 2004). Also, the abundance of membrane
10 fatty acids with 16 and 18 carbons can change as a result of temperature (Maffei et al., 1993;
11 Williams et al., 1995; Matteucci et al., 2011). Often, under heat stress the relative abundance
12 of C_{16:0} fatty acid was found to increase and vice versa the abundance of polyunsaturated
13 C_{18:3} fatty acid to decrease (Larkindale and Huang, 2004; Bakht et al., 2006). Furthermore,
14 effects of temperature were observed for mono- and sesquiterpenes, with compounds like
15 limonene and myrcene having a close correlation with temperature, whereas others like 1,8-
16 cineol were not affected by temperature (Maffei et al., 1993). As a cause, a different
17 sensitivity of individual steps in the genesis of the wax lipid components is assumed
18 (Shepherd and Griffiths, 2006). However, results were found to vary between different
19 species and genotypes, indicating a species or genotype related sensitivity to changes in
20 irradiation (Shepherd and Griffiths, 2006), whereas cold- or heat-acclimated plants respond
21 differently than those that are not acclimated (Larkindale and Huang, 2004). Thus, a
22 dependency of temperature and lipid metabolism is widely observed, but especially in plants
23 other factors such as humidity or greenhouse gas composition might coincide with a larger
24 effect on the overall lipid composition.

25 In addition to the effect of temperature on lipid synthesis, temperature can also influence
26 lipids after production specifically as a result of fire. This topic is addressed in section 4.2.

27 **2.3.2 Effects of humidity**

28 With respect to the effects of water stress and/or high humidity, in their review Shepherd and
29 Griffith (2006) reported mixed results, with respect to absolute amounts as well as chain
30 length distribution. Bondada et al. (1996) reported an increase in absolute amounts of
31 epicuticular wax production by 69 % in the leaves of cotton (*Gossypium hirsutum*) under
32 water stress, which was confirmed by Hamrouni et al. (2001), Koch et al. (2006), Kim et al.

1 (2007), and Bettaieb et al. (2010) for neutral lipids of other plant species. However, Kim et
2 al. (2007) found that water stress had only a minor effects on chain length distribution. The
3 relative contribution of different component classes to the wax composition remained
4 unchanged except for *Brassica oleracea* var. *gongylodes* at the highest relative humidity,
5 which showed an increased contribution of ketones and primary alcohols and a reduction of
6 secondary alcohols and aldehydes (Koch et al., 2006). Recently, Srivastava et al. (2017)
7 determined that sustainable effects of drought on plant lipid composition are commonly
8 missing with few exceptions for perennial plants. Thus, several months after exposure to
9 drought the lipid biosynthesis and composition of leaves is resilient. The existing data shows
10 that general effects of drought on plant lipid composition are difficult to draw.

11 **2.3.3 Effects of increased CO₂**

12 Changes in greenhouse gases such as CO₂ have also been discussed to influence the lipid
13 biosynthesis and thus the lipid composition of plants. Short-term exposure of several hours to
14 elevated CO₂ concentrations e.g. during ¹³CO₂ or ¹⁴CO₂ labelling experiments has no or little
15 effect on the lipid composition, especially if sampling occurs several days after labelling
16 (Wiesenberg et al., 2009). In contrast a long-term rise in atmospheric CO₂ concentration has
17 been investigated in laboratory or free air carbon dioxide enrichment (FACE) experiments
18 (Ainsworth and Long, 2005). Although numerous such experiments have been maintained in
19 the meantime, implication of investigations of lipid composition is limited. Greenhouse
20 experiments showed that elevated CO₂ concentration affects the relative composition of
21 saturated and unsaturated fatty acids in wheat plants (Williams et al., 1994; Williams et al.,
22 1995; Williams et al., 1998). However, rising nitrogen fertilization and rising temperature can
23 lead to competing trends so that with elevated temperature and nitrogen fertilization
24 (Williams et al., 1995; Griepentrog et al., 2016). Although specific abundances of individual
25 long-chain alkanes and alcohols changed under elevated CO₂ concentration, the overall lipid
26 composition expressed as ACL and CPI did not change (Huang et al., 1999). Nevertheless,
27 concentration changes like an increase in *n*-alkane and *n*-alcohol abundances and a decrease
28 in *n*-fatty acid abundance was determined under rising CO₂ concentration, whereas nitrogen
29 fertilization led to a decrease in the effect (Huang et al., 1999), which was confirmed by
30 Wiesenberg et al. (2008a) for *n*-alkanes, *n*-fatty acids and *n*-alcohols. In some forest FACE
31 and open top chamber experiments, the effect of elevated CO₂ on plant lipid concentration
32 were not identified (Feng et al., 2010; Griepentrog et al., 2015), but the ¹³CO₂ labelling

1 associated with the CO₂ enrichment was used for tracing turnover of lipids in soils as
2 introduced by Wiesenberg et al. (2008b) for lipids.

3 **2.4 Other or combined genetic, ontogenetic and/or environmental effects**

4 Many studies considered the effects of e.g. geographical location on wax amounts and/or
5 composition without differentiating between individual genetic or environmental causes.
6 Again the exact parameters investigated vary greatly between studies, as do the conclusions
7 drawn. Cowlshaw et al. (1983) examined the *n*-alkane, *n*-alcohol, *n*-aldehydes and ester
8 composition of composite samples of four species of the Poaceae *Chionochloa*, one of which
9 was sampled at three different environmental locations to investigate environmental effects.
10 They found distinct chain length patterns that allowed for chemotaxonomic identification,
11 where variation between the three sampling sites did not alter dominant chain length patterns
12 for any of the component classes (Cowlshaw et al., 1983). Similar observations were made
13 by Herbin and Sharma (1969) for ω -hydroxy fatty acid composition of *Pinus* species from
14 Asia, Europe, North-America, Central America and the Caribbean. Kreyling et al. (2012)
15 described differences in the *n*-fatty acid and *n*-alkane composition of the same plant species
16 originating from different regions across Europe with different climatic conditions most
17 likely due to biosynthetic adaptation to the specific conditions. Piervittori et al. (1996) found
18 that the distribution of C₂₅, C₂₇, C₂₉ and C₃₁ *n*-alkanes in the lichen *Xanthoria parietina*
19 varied significantly between two different Piedmont valleys in Italy, and within those with
20 altitude, reflecting a combined influence of elevation, water availability, radiation and
21 temperature. For plaggen ecosystems Kirkels et al. (2013) also observed a significant
22 variability in reported ratios of the dominant *n*-alkanes with chain lengths C₂₇, C₂₉, C₃₁, C₃₃
23 most likely attributable to the causes examined here. However, in spite of this they found
24 meaningful clustering of the three different plant groups grasses, shrubs and trees indicating
25 that the variability did not obliterate the power of distinction (Kirkels et al., 2013). In a larger
26 study based on 2093 observations from 86 sources of plant material, Bush and McInerney
27 (2013) concluded that the general observation that C₂₇ and C₂₉ *n*-alkanes are dominant
28 markers for woody vegetation and C₃₁ for graminoids does not rigorously hold true. At the
29 same time C₂₃ and C₂₅ *n*-alkanes do seem to be robust indicators of *Sphagnum* (Bush and
30 McInerney, 2013) as already observed by Baas et al. (2000) and Pancost et al. (2002). Bush
31 and McInerney (2013) indicated that the lack of rigour of the mentioned proxies is likely caused

1 by environmental conditions as indicated by a shift in patterns across African savannah and
2 rainforest environments.

3 The distinction between African savannah and rainforest environments in general and C₃
4 versus C₄ vegetation in particular have been the subject of more detailed research. Vogts et
5 al. (2009) studied the leaves and sometimes whole plants of 24 African rain forest and 45
6 savannah species. They found that as a result of environmental influence, including
7 temperature and aridity, chain length distributions of the *n*-alkanes and *n*-alcohols of some
8 species shifted to different chain length predominance. The environmental influences
9 overshadowed a taxonomic distinction at the order, family or sub-family level (Vogts et al.,
10 2009). Patterns in grasses were more consistent and thus less dependent on environmental
11 factors (Vogts et al., 2009). As a result, in spite of the environmental variability observed,
12 Vogts et al. (2009) found that by averaging lipid patterns within a given environment a clear
13 distinction between rain forest and savannah plants can be made, with a dominance of C₂₉ *n*-
14 alkane representative of the average rain forest plant signal and a dominance of C₃₁ *n*-alkane
15 of the savannah plants and C₄ savannah grasses. For the *n*-alcohols, C₂₈ dominated on
16 average for savannah plants, C₃₀ for rain forest plants and C₃₂ for C₄ savannah grasses (Vogts
17 et al., 2009).

18 Rommerskirchen et al. (2006) observed a generally higher content of C₃₁ and C₃₃ *n*-alkanes
19 and therefore higher ACL value in African C₄ grasses with respect to C₃ grasses from the
20 same area as a result of the genetic adaptation of C₄ grasses to warm, arid habitats. In
21 addition, *n*-fatty acid patterns have also been shown to vary with C₃ and C₄ metabolism, with
22 C₃ crops having relatively large proportions of C₂₄ *n*-fatty acid in leaves, stem and roots as
23 compared to C₂₂ and C₂₆ *n*-fatty acids in C₄ crops (Wiesenberg and Schwark, 2006).

24 **2.5 Conclusions and implications regarding source related variability**

25 Already Herbin and Robins (1969) concluded that there is a basic genetic control on the
26 composition of the wax components, including the *n*-alkanes, of plant leaves. However,
27 variable factors associated with age and environment can be superimposed upon the specific
28 pattern in some cases, while in others the genetically controlled pattern appears to be stable
29 and unaffected by external influences (Herbin and Robins, 1969). Now, 48 years later, a
30 much more extensive database has been accrued, albeit with a large emphasis on leaf wax
31 lipids in general and *n*-alkanes in particular. Nevertheless, the results are still equivocal. On
32 the one hand, there is ample evidence that genetically driven variability of leaf wax lipid

1 composition in principle leads to chemotaxonomically meaningful clustering that can form
2 the basis of the application of leaf wax lipids as molecular proxies. On the other hand, it is
3 clear that both ontogeny and environmental factors can have a significant and sometimes
4 dominant influence on lipid composition like e.g. chain length distribution. Matters are
5 complicated by the fact that much data with respect to the effects of environmental stress
6 originates from studies where plants were studied for a limited period of time (typically one
7 growing season), where extreme conditions were artificially imposed. In contrast, the lipid
8 signal from soil or sediment archives as used in reconstructions typically represents a mixture
9 of input of decades or longer from plants in various life stages of perennial plants, the
10 induced diversity of plants by frequent changes of annual plants in managed ecosystems and
11 the average of natural fluctuations in stress conditions during that time period.

12 In general from what is known to date, the conclusion seems justified that on the one hand
13 because of genetic and environmental influences there are no unique compounds nor ‘golden
14 ratios’ of different chain lengths of compounds that can always be linked to certain plants
15 under all circumstances. On the other hand, there are many situations where the influence of
16 genetic and environmental effects are small enough that they do not prevent the use of plant
17 lipids as molecular proxies. The currently available data does not allow for objective,
18 quantitative rules to be formulated in this respect. From the plant wax components, the *n*-
19 alkanes are the dominant class studied. In addition, research attention has focussed to a lesser
20 extent on *n*-alcohols and *n*-fatty acids. The other wax components such as isoprenoids and
21 ester bound lipids received hardly any research attention to date with respect to source related
22 variability in the context of their use as molecular proxies. Yet even for the *n*-alkane patterns
23 in leaf waxes, only a tiny portion of dominant plant species on the planet have been examined
24 in detail for the effects of genetics and environment on their amounts and patterns. It is clear
25 that much more research is needed in this respect.

26 Based on the current insights it seems prudent to explicitly take the possibility of genetically
27 and environmentally driven variability of lipid patterns into account when considering the use
28 of lipids as molecular proxies. For instance by considering plant species from the same
29 climatic zone as where the reconstruction takes place, and by mixing plant material from
30 different life stages to obtain the average molecular fingerprint to look for.

31

1 **3 Input pathway related variability of molecular proxies**

2 **3.1 Definition**

3 Here we discuss differences in the amount and composition of molecules used as proxies,
4 which is possible due to different input pathways of such molecules to the soil. A schematic
5 representation of the different input routes of molecular proxies into the soil is provided in
6 Fig. 1. The emphasis lies on potential effects for their use as molecular proxies. For a general
7 description of the different molecular origins of organic matter in soil, the reader is referred
8 to a dedicated review on this topic by Kögel-Knabner (2002).

9 **3.2 Leaf versus root input**

10 Conservative estimates calculate roots to represent 33 % of global annual net primary
11 productivity (Jackson et al., 1997), whereas more recent studies highlight that the
12 contribution of root-derived organic matter in soils can account for >70 % of total plant-
13 derived carbon (Rasse et al., 2005). As a result, roots form a considerable input of organic
14 matter in soils and are proposed to improve carbon storage in soils (Kell, 2012). In addition,
15 root input occurs to considerable depth in soils, ranging from an average depth of 0.5 m in
16 tundra biomes to 15.0 m in tropical grassland/savannah (Canadell et al., 1996). But also in the
17 temperate zone under certain circumstances such as the presence of nutrient rich fossil A
18 horizons at depth, deep rooting can be very significant (Gocke et al., 2015). However, on
19 average the majority of root biomass appears to be incorporated in the top 0.3 m of the soil in
20 most biomes, i.e. in the topsoil (Jackson et al., 1996). The ratio of root/shoot biomass input is
21 also very variable across biomes, ranging from an average of 0.10 in cropland to 4.5 in
22 deserts (Jackson et al., 1996). Table 2 represents an overview of the average maximum
23 rooting depth, root biomass input in the first 0.3 m of the soil and root/shoot biomass input
24 for different biomes (see also Fig. 1).

25 Therefore, if the molecules to be used as proxy are present in both leaves and roots of plants,
26 the possibility of root input is a factor that has to be considered depending also on the
27 purpose of the proxy. In the case of cutin and suberin monomers root input does not cause
28 interference as discerning root from leaf input is the specific purpose of this molecular proxy
29 (Mendez-Millan et al., 2011). However, this may be different for the wax lipids, i.e. *n*-
30 alkanes, *n*-alcohols, *n*-fatty acids and isoprenoids, that have been found to occur in leaves as
31 well as roots of species at varying concentrations (Jansen et al., 2007; Huang et al., 2011).

1 Particularly when such wax derived lipids are applied as molecular proxies for vegetation
2 cover in soil, root input can be an issue for two reasons: i) roots may contain a different wax
3 lipid composition than leaves qualitatively and quantitatively, thereby clouding the leaf signal
4 (Jansen et al., 2006; Martelanc et al., 2007); ii) young root input at depth may disrupt the
5 chronology of a reconstruction in time by overprinting the originally present signal (Lavrieux
6 et al., 2012; Gocke et al., 2014).

7 The main discussion with respect to the influence of root input in wax lipid based
8 environmental reconstructions from soils therefore revolves around assessing the relative
9 importance of root versus aboveground biomass input. Since plant wax lipids reside on the
10 outer parts of leaves and roots, relative surface area and bioproductivity are important. On a
11 global scale root surface area is almost always calculated to be higher than leaf surface area,
12 more than an order of magnitude so in grasslands (Jackson et al., 1997). However, in many
13 cases the absolute amount of lipids present per mass unit of root material is an order of
14 magnitude or more lower than on leaf material (Marseille et al., 1999; Zech et al., 2011). In
15 addition, also with respect to the degree in which wax lipid chain length patterns vary
16 between the leaves and roots of plant species, there appears to be quite some variability. In
17 general, the observed differences between roots and leaves of the same species are reported to
18 be of the same order of magnitude as the differences between leaves of different species (e.g.
19 Jansen et al., 2006; Kirkels et al., 2013; Gocke et al., 2014).

20 The concurrent influence of such various quantitative and qualitative factors makes the
21 impact of root input a complex issue that still is subject of scientific debate (Wiesenberg and
22 Gocke, 2013). Given that different factors will have a highly variable influence in different
23 situations, no general conclusion can be drawn. In some situations, the influence of roots as
24 input pathway of extractable lipids to be used as molecular proxy may be limited (Quenea et
25 al., 2006). In others, root input may be dominant (Van Mourik and Jansen, 2013). In addition,
26 the relative degree of influence may vary greatly with depth leading to the concurrent
27 presence of leaf lipid dominated and root lipid dominated zones at different depths in the
28 same profile (Angst et al., 2016).

29 **3.3 Microbial input**

30 In general, microbial biomass can be a significant source of soil organic matter, with up to
31 40 % transformed to non-living soil organic matter, but is turned over much faster than plant
32 residues (Miltner et al., 2012). Focussing specifically on lipids, isotopic studies show that

1 90 % of fatty acids of microbial origin are turned over rapidly after cell death, whereas the
2 majority of biomass derived residual bulk C was stabilized in the non-living soil organic
3 matter pool (Kindler et al., 2009). In spite of the potentially shorter residence time, a
4 concurrent faster production makes that microorganism derived molecules are a factor to
5 consider when applying molecular proxies in soils except when such proxies are used to
6 study microbial input.

7 For wax lipids generally *n*-alkanes, *n*-alcohols and *n*-fatty acids with longer chain lengths
8 (>C₂₀) and a distinct odd-over-even (*n*-alkanes) or even-over-odd (*n*-alcohols and *n*-fatty
9 acids) chain length predominance are considered to be higher plant derived, whereas shorter
10 chain length homologues are considered to be predominantly of microbial origin (Eglinton et
11 al., 1962; Diné et al., 1990). Moreover, with the exception of an abundance of C₁₆ and C₁₈ *n*-
12 alcohol and *n*-fatty acid, such microbial lipids are described to lack a specific chain length
13 predominance (Stevenson, 1994; Lichtfouse et al., 1995). However, several researchers
14 challenge the observation that higher chain length lipids in soils are exclusively of higher
15 plant origin. Microorganisms have been shown capable of synthesizing higher chain length
16 straight-chain lipids, albeit usually to a limited extent (Ladygina et al., 2006; Nguyen Tu et
17 al., 2011). Jambu et al. (1978) indicated that while chain lengths >C₂₀ in soils are
18 predominantly plant derived, particularly in acidic soils fungi may contribute such lipids as
19 well. Furthermore, Marseille et al. (1999) observed an abundance of C₂₅ and C₂₇ *n*-alkanes
20 that they also attribute to *in-situ* production by fungi. This was confirmed for an agricultural
21 soil by Quenea et al. (2006), who observed old forest and fungi derived odd long-chain
22 alkanes based on compound-specific isotope analysis and lipid distribution patterns. Possible
23 pathways of *in-situ* genesis of *n*-alkanes in soils are reduction of *n*-alkenes and *n*-alcohols,
24 decarboxylation of bacterial *n*-fatty acids as well as degradation of biopolymers containing
25 aliphatic side chains (Lichtfouse et al., 1998). Nevertheless, based on the large number of
26 studies where typical higher plant derived patterns of lipids are reported and used in soils
27 (Table 1), including indicative ACL and CPI values, microbial input of longer chain length
28 straight-chain lipids generally does not seem to be a major factor compared to direct plant
29 derived input in the topsoil (Jansen and Nierop, 2009; Bai et al., 2009). In contrast, for
30 steroids and triterpenoids such as camosterol, stigmasterol and lupeol, microbial input in
31 soils can be considerable (Naafs et al., 2004). As another example, arbuscular mycorrhizal
32 fungi derived β -sitosterol is by far the most abundant sterol identified in soils (Grandmougin-
33 Ferjani et al., 1999).

1 With respect to cutin and suberin monomers, *in-situ* genesis in soils through microbial
2 transformation of other precursor molecules can be an issue. For instance, oxidation of free
3 fatty acids could be a source of ω -hydroxy fatty acids, whereas microbial β -oxidation of
4 unsaturated fatty acids and/or mid-chain hydroxy fatty acids may be a source of α,ω -
5 alkanedioic acids, thus clouding the cutin/suberin signal (Naafs et al., 2004)

6 **3.4 Airborne input**

7 In addition to *in-situ* production and incorporation of soil lipids, airborne input must be
8 considered. The distance of airborne transport of larger constituents such as leaves can be
9 expected to be limited. However, smaller physical forms containing lipids, such as aerosols
10 and dust particles, can travel substantial distances (Conte and Weber, 2002) thus causing
11 input of alien molecules that may influence the local signal. This is of special importance
12 where airborne sediments with low content of organic matter are investigated as in these
13 environments already low inputs of foreign organic matter can significantly influence the
14 molecular proxies. Liu et al. (2007) showed that the $\delta^{13}\text{C}$ signature of sediment organic
15 carbon in loess deposits of the western Chinese Loess Plateau corresponds to that of dust
16 sources instead of the local vegetation. While in a study of marine sediment cores along the
17 Southwest African continental margin, Rommerskirchen et al. (2003) revealed that aerosol
18 derived input of higher chain-length *n*-alkanes and *n*-alcohols provides a significant signal,
19 the $\delta^{13}\text{C}$ signal of which corresponded well with continental C_3/C_4 plant distribution and
20 fossil pollen input when prevailing wind patterns were taken into account. However, in this
21 case, in contrast to vegetated soils, there was no *in-situ* input from higher plant vegetation.

22 Aerosol studies above plant canopies revealed a certain relationship of the plant wax
23 composition of the present plants, but significant differences from the biomass were observed
24 for *n*-alkanols and *n*-alkanes (Conte et al., 2003). While the wax molecular composition was
25 not directly linked between biomass and aerosol, especially the compound-specific isotope
26 composition ($\delta^{13}\text{C}$) revealed a closer link of both. For Bermuda aerosols it could be shown
27 that the aerosol compound-specific isotope composition of *n*-alcohols and *n*-fatty acids
28 reflects the plant wax compound-specific isotope composition as well as the course of the
29 bioproductivity during the different seasons of the years (Conte and Weber, 2002).

30 In a study of PM_{10} aerosols collected during a winter season in Baoij, China, Xie et al. (2009)
31 found concentrations of $\text{C}_{21}\text{-C}_{33}$ *n*-alkanes in the $10\text{-}600\text{ ng/m}^3$ range as a result of intensive
32 coal burning in the region. In a two year study of PM_{10} and $\text{PM}_{2.5}$ aerosols in urban sites in

1 Nanjing, Wang et al. (2006) observed C₂₁-C₃₃ *n*-alkanes present in the 10-100 ng/m³ range.
2 Concentrations of C₂₁-C₃₅ *n*-alkanes in PM₁₀ aerosols in urban sites in Beijing sampled in all
3 seasons were even lower (Zhou et al., 2009). In this study also *n*-fatty acids and hopanes were
4 considered, but were found in small concentrations that, together with the *n*-alkanes,
5 constituted ca. 3 % of the total organic matter in the aerosols (Zhou et al., 2009). In all
6 studies, the straight chain lipid patterns lacked the odd-over-even chain length predominance
7 typical of higher plants (Wang et al., 2006; Xie et al., 2009; Zhou et al., 2009). Nevertheless,
8 in a large survey a clear odd-over-even chain length predominance was found in spite of such
9 potentially intense aerosol derived input (Rao et al., 2011). This indicates that even in areas
10 under large aerosol deposition, as in the case of intensive anthropogenic pollution associated
11 with fossil fuel burning, the effect of aerosol deposition on *n*-alkane patterns in the soil is
12 limited as a result of the large *in-situ* input via roots and leaves of the local vegetation.

13 **3.5 Conclusions and implications regarding input pathway related variability**

14 The diversity of input pathways offers both opportunities and limitations for the use of
15 molecular proxies. Opportunities arise when different sources can be elucidated using
16 molecular proxies. Examples are the differences in molecular composition of leaf and root
17 waxes as used to differentiate between their respective influences, or when aerosol associated
18 lipids are used for source apportionment of terrestrial plant input in terrestrial or marine
19 sediments. This can help budgeting organic matter input of different sources and thus
20 improve (paleo-)environmental interpretations and reconstructions. Limitations are posed
21 when input through multiple pathways clouds the linkage of a (set of) molecule(s) to a certain
22 source for which it is to serve as proxy. For instance when linking a suite of straight-chain
23 lipids to a particular group of plants at a certain site. When looking at the application of
24 molecular proxies in soils, in particular the assessment of the influence of root derived input
25 is a challenge that is not always acknowledged. The significance of root derived organic
26 matter in soils and terrestrial sediments has been neglected for decades and has only been
27 recently highlighted (Rasse et al., 2005; Rumpel and Kögel-Knabner, 2011). More research
28 attention is needed to pinpoint how large possible interferences are and how the potential can
29 be to compensate for them, e.g. through modelling approaches. For instance, the VERHIB
30 model was designed to unravel the mixed *n*-alkane, *n*-alcohol and/or *n*-fatty acid signal
31 observed in soils into the most likely combination of plant groups responsible for the original

1 lipid input, treating leaves and roots explicitly as separate entities (Jansen et al., 2010). This
2 might form a starting point to disentangle leave and root derived lipid input.

3 Although the aerosol studies so far provide useful information that plant wax components are
4 transported via aerosols to remote places, other factors like degradation during transport and
5 integration of regional vegetation patterns may hamper direct source-to-sink relationship of
6 airborne molecular markers. Nevertheless the overall impact of aerosol borne molecules on
7 molecular proxy based reconstructions seems to be limited whenever the total abundance in
8 the soil is high.

9

10 **4 Transformations and turnover in soil**

11 Transformations and turnover of soil organic matter are an important study area in their own
12 right (Kögel-Knabner, 2002; Von Lützow et al., 2008). Important in the context of the
13 application of molecular proxies is the recent paradigm shift to the attribution of external
14 factors as drivers of SOM turnover rates as opposed to inherent recalcitrance related to
15 molecular structure (Schmidt et al., 2011; Lehmann and Kleber, 2015). Coupled to this are
16 indications that microbial recycling of organic matter upon entering the soil decouples the
17 molecules from their biological sources (Miltner et al., 2012; Gleixner, 2013). Here, we focus
18 on the effects of (differences in) transformations/degradation of molecules in soils for their
19 use as molecular proxies. This includes transformations during the stages of senescence or
20 litter and covers attempts to estimate successive degradation processes of organic matter
21 occurring after burial until stages of long-term preservation (see also Fig. 1). Transformation
22 processes can also include processes that affect the detectability of a molecule used as proxy,
23 for instance a transformation from the extractable to the non-extractable lipid fraction as a
24 result of chemical alterations or interactions with the mineral phase (e.g. Almendros et al.,
25 2001). A special case is the influence of fire on SOM, including molecular proxies, as
26 reviewed by González-Pérez et al. (2004).

27 All of the attempts dealing with incorporation and preservation of organic matter deal with
28 different assumptions and entail different problems in terms of uncertainties. Thus, in
29 dependency of the environmental conditions, assumptions that are relevant for incorporation
30 and burial of organic matter play a major role, as should the different aspects of degradation
31 and preservation. However, currently much uncertainty exists regarding the influences of
32 individual environmental and genetic factors concerning degradation and preservation.

1 Therefore, the following paragraphs only provide the first insights tackling these issues,
2 which need further attention in future research projects.

3 Molecular transformations and variations thereof of molecular proxies mostly offer
4 complicate application of molecular proxies. However, in some instances they may also offer
5 opportunities. For instance, *n*-alkanes can be degraded to *n*-methyl ketones through β -
6 oxidation (Chaffee et al., 1986; Amblès et al., 1993), which can be used to assess and trace *n*-
7 alkane degradation in soils (Jansen and Nierop, 2009). Similarly, the presence of certain *seco*-
8 acids formed through A-ring opening of 3-oxytriterpenoids under anaerobic conditions, may
9 be used as proxy for the occurrence of such anaerobic episodes (Jaffe et al., 1996), e.g. under
10 stagnant water conditions.

11 **4.1 Differences related to incorporation pathway**

12 The incorporation pathway (Fig. 1) may influence subsequent turnover of molecular proxies.
13 This includes (differences in) degradation during senescence and/or litter degradation stages,
14 e.g. due to different input shapes (like root vs. leaf) offer a different degree of physical
15 protection. It also includes alterations induced by fire prior to or upon incorporation of
16 organic matter into the soil.

17 In a study of *Gingko biloba* leaf wax lipids during the senescence and litter stages, Nguyen
18 Tu et al. (2003) found limited degradation that did not affect the dominant chain lengths of
19 alkyl molecular proxies. When comparing different classes of wax lipids they found the *n*-
20 alkanes to be the most resistant to degradation, followed by the *n*-fatty acids and then the *n*-
21 alcohols (Nguyen Tu et al., 2003). Also, more in general, in a study of grassland and forest
22 soils, Otto and Simpson (2005) determined that characteristic patterns of wax lipids and
23 isoprenoids were preserved throughout the stages between fresh plant material and soil
24 organic matter. They also determined preferential enrichment of suberin with respect to cutin
25 monomers in particular in one of the grassland soils (Simpson et al., 2008). This indicated for
26 example the fact that the former is embedded in woody tissue while the latter is exposed on
27 leaf surfaces (Simpson et al., 2008) (see also 4.3.3).

28 When looking at bulk organic matter in soils, Rasse et al. (2005) estimated that the main
29 residence time of root derived organic matter is on average 2.4 times that of shoot derived
30 organic matter. When comparing cutin and suberin monomers, Andreetta et al. (2013)
31 described selective preservation of leaf derived monomers in the more acidic and dryer soil,

1 while in the more fertile soil root derived monomers were preferentially preserved. They
2 attributed the former to inhibited microbial degradation due to drought and acidity, and the
3 latter to protection within aggregates. In another study still small differences in degradation
4 of the same *n*-alkanes that derived from different plants were found, with a slower
5 degradation of *n*-alkanes derived from more woody roots (Nierop and Jansen, 2009),
6 although lipids were generally well preserved. Killops and Frewin (1994) reported that
7 persistency of plant cuticles protected their composite isoprenoids from degradation in
8 mangrove sediments. Similar preservation in soils is also perceivable.

9 More in general, Mambelli et al. (2011) observed root litter, including biomarkers, to be
10 selectively preserved with respect to needle litter, which was confirmed by Mendez-Millan et
11 al. (2010) for maize and wheat roots versus shoots. Using isotopic signatures, Mendez-Millan
12 et al. (2011) were able to quantify and subsequently compensate for such differences in
13 turnover rate. This further emphasizes the significance of root derived organic matter for
14 turnover determinations as already discussed by Wiesenberg et al. (2004). In other words, the
15 relative abundance of roots and the uncertainties in terms of root related overprint in the
16 rhizosphere and rhizosphere extension entail large uncertainties and strong differences
17 between different plant species and environmental settings, especially at a molecular level.
18 Further research is required to enable extrapolations to or across ecosystem scales.

19 With respect to the effects of fire, burning of litter or biomass can release additional
20 extractable lipids (González-Pérez et al., 2004). In addition, fire has been reported to alter the
21 chain length distribution of *n*-alkanes and *n*-fatty acids, shifting it towards shorter chain
22 lengths (González-Pérez et al., 2004; Wiesenberg et al., 2009; Knicker et al., 2013). Also the
23 composition of terpenoids can be influenced, resulting in preferential degradation of those
24 with the lowest thermal stability (González-Pérez et al., 2004). All such processes potentially
25 adversely affect the application of molecular proxies to an extent that depends on the
26 frequency and intensity of fires. At the same time, fire events may also offer opportunities.
27 For instance thermal alteration of animal fats in fireplaces may produce specific *n*-alkane/*n*-
28 alkene doublets preserved in the soil, the presence of which can be used to reconstruct human
29 fire usage in an archaeological context (Lejay et al., 2016).

30 **4.2 Differences between different soil compartments**

31 When soils are used as archives of molecular proxies, mostly bulk samples are used and
32 replication per horizon or stratigraphic layer is often limited or absent. However, several

1 studies indicate that preservation of molecules used as proxies can differ between different
2 soil compartments (Flessa et al., 2008; Clemente et al., 2011; Griepentrog et al., 2014).
3 Depending on the research question this may pose a problem, for instance it might obscure
4 chronology when molecules are used as proxies to reconstruct changes over time.

5 Already Lichtfouse et al. (1998) showed that straight-chain lipids can become encapsulated in
6 larger organic macromolecules, thus being protected against degradation. In addition,
7 physical protection in (micropores of) aggregates and/or through association with clay
8 minerals have been identified as important pathways for stabilization of soil organic matter in
9 general, including molecules used as molecular proxies (Tonneijck et al., 2010). Using bulk
10 and compound-specific $\delta^{13}\text{C}$ analysis, Cayet and Lichtfouse (2001) showed that plant-
11 derived *n*-alkanes in a soil under maize cultivation varied in average age per particle size
12 fraction, with the C_{31} *n*-alkane from the 200-2000 μm fraction being significantly younger
13 than that from the 50-200 μm and 0-50 μm fractions. A general trend of preferential
14 preservation in smaller size fractions, in particular the clay fraction, is also reported in other
15 studies. For instance, Quenea et al. (2004) and Flessa et al. (2008) observed longer turnover
16 rates of SOM in smaller size fractions. Clemente et al. (2011) studied the preservation of long
17 chain aliphatic compounds in three soils with similar clay mineralogy but different carbon
18 contents and standing vegetation. Irrespective of these differences, they too found the
19 aliphatic compounds to be preferentially preserved in the silt and clay fractions, and again
20 linked this to strong interactions with the present clay minerals. In a recent study,
21 Griepentrog et al. (2015, 2016) confirmed the higher residence time of organic matter in
22 small sized density fractions when compared to macro-aggregates as a result of interaction
23 with the mineral phase. This implies an improved preservation of organic matter associated
24 with higher density and thus mineral association when compared to organic matter associated
25 to lower density. However, physical fractionation techniques such as particle and density
26 fractionation have a potential of creating analytical artifacts, especially when molecular
27 proxies are investigated. In addition, occlusion or strong adsorption in the smallest mineral
28 fractions might hamper extraction and analysis of the proxy in question.

29 Furthermore, the effects of size or density fractions of soil on preservation of organic matter,
30 including molecular proxies, are not uniform. For instance, Höfle et al. (2013) found size and
31 density fraction related organic matter stabilization to be much less pronounced in the active
32 upper layer than in the deeper soil horizons. This points to selective preservation of organic
33 matter in the deeper soil because of more extensive aggregation and organo-mineral

1 association. In a study of volcanic ash soils, Stewart et al. (2011) did not find differences in
2 preservation of bulk SOM in general or lipids in particular between different size fractions.
3 They attributed this lack of differentiation to the presence of a large proportion of the SOM
4 that was not associated with mineral components as these were already saturated with
5 previously incorporated soil organic matter (Stewart et al., 2011).

6 In general a combination of physical protection and sorptive preservation seems to be
7 responsible for the observed differences (or lack thereof) in preservation of organic molecules
8 in soils between different size or density fractions. This is corroborated amongst others by a
9 study by Guggenberger et al. (1995), where they observed differences in the preservation of
10 SOM derived from tropical pastures compared to the preceding native savannah vegetation.
11 They attributed this effect to a difference in interactions with the mineral phase, leading to
12 physical protection of SOM and molecular proxies contained therein. Similarly, differences
13 in turnover rates of ca. one order of magnitude between forest and grass derived molecules
14 after land use change have been observed as a result of saturation of the adsorption sites on
15 the mineral phase (Hamer et al., 2012).

16 In addition to heterogeneity in the effects of interactions with the mineral phase on
17 preservation of molecular proxies, analytical artifacts cannot be completely excluded when
18 physical and chemical fractionation techniques are applied to separate particle size or density
19 fractions. To date systematic investigations addressing these issues are lacking, which
20 hampers the drawing of general conclusions with respect to processes that are relevant e.g.
21 under different climates and for different soil mineralogical composition.

22 **4.3 Selective preservation within or between classes of molecules**

23 Turnover rates of molecular proxies do not only vary between different compartments, but
24 may also vary within the same compartment; between and even within different (classes of)
25 molecules (Dinel et al., 1990; Bull et al., 2000; Amelung et al., 2008). For instance, Feng and
26 Simpson (2007) found preferential enrichment of straight-chain lipids as well as cutin and
27 suberin monomers with increasing depth with respect to bulk SOM. In contrast, in a study of
28 grain-maize and silage-maize cropped soils Wiesenberg et al. (2004) found turnover times in
29 the sequence bulk SOM > *n*-alkanes > *n*-fatty acids, with rate differences that varied
30 substantially between the two cultivations. The differences could be related to the different
31 biomass input on the one hand and large amount of lignite dust and the low biomass input on
32 the other hand, thus hampering degradation at this site. The faster turnover of *n*-fatty acids

1 than *n*-alkanes as also confirmed by Wiesenberg et al. (2008a) and Griepentrog et al. (2015;
2 2016). In contrast, it may also offer opportunities to apply such differences between
3 molecular classes and their response to external factors to trace transformations and input of
4 organic matter in soils (Feng and Simpson, 2007).

5 An important issue with respect to the application of straight-chain lipids as molecular
6 proxies is also preferential degradation of certain chain lengths within a certain class of
7 molecules, as molecular ratios of various (higher) chain lengths are often used as proxies for
8 certain vegetation types (see paragraph 2). This issue is addressed in the following
9 paragraphs.

10 **4.3.1 Straight-chain lipids**

11 Already Moucawi et al. (1981) reported decreasing degradation rates with larger chain-length
12 for *n*-alkanes in soils, which was confirmed by Lichtfouse et al. (1998) who determined a
13 higher resistance of long straight-chain lipids in soil compared to their shorter chain
14 counterparts. However, such preferential degradation was found in agricultural and acidic
15 soils and in the absence of Fe(OH)₃ (Moucawi et al., 1981; Lichtfouse et al., 1998). Similar
16 results were found for other lipid classes as well (Moucawi et al., 1981). More recently,
17 several authors also indicate that such preferential degradation can occur in other soils
18 (Jansen and Nierop, 2009; Cui Jingwei et al., 2010). However, the extent of the effect
19 questions the suitability of the compounds in question as molecular proxies. For instance,
20 Jansen and Nierop (2009) found the overall effect of preferential degradation on higher plant
21 derived *n*-alkane patterns in soils to be small and not of influence for their use as vegetation
22 proxy. Similarly, Lei et al. (2010a, 2010b) determined that in spite of strong evidence of
23 microbial degradation, relative abundance of long-chain *n*-alkanes could still be used to
24 distinguish coniferous from broadleaf tree input in soils.

25 Within the group of straight-chain lipids, overall degradation rates of subclasses have been
26 found to vary depending on soil physicochemical properties. For instance, *n*-alkanes have
27 been reported to be better preserved in alkaline soils, whereas *n*-fatty acids accumulate in
28 more acidic soils (Simpson et al., 2008).

29 **4.3.2 Isoprenoids**

30 Isoprenoids are reported to have varying turnover rates both under oxic and anoxic conditions
31 in soils (Jaffe et al., 1996; Amelung et al., 2008). Generally, sterols, diterpenes and

1 pentacyclic triterpenes are reported to be turned over rapidly as compared to straight-chain
2 lipids in grassland as well as forest soils, hindering their application as molecular proxies for
3 their sources (Bull et al., 2000; Naafs et al., 2004; Jansen et al., 2007). However, Otto and
4 Simpson (2005) observed the exact opposite trend, indicating a strong environmental control
5 on the relative transformation rate of different classes of components. In an incubation study
6 of derived triterpenols, Koch et al. (2005) highlighted marked differences between
7 degradation rates of individual triterpenols, leading to a sharp relative increase in the
8 proportion of taraxerol with respect to the other triterpenols.

9 In addition, Δ^5 sterols are transferred both aerobically and anaerobically to 5α - and 5β -stanols
10 (De Leeuw and Baas, 1986), which are reported to persist much longer in soils than their
11 precursors (Bull et al., 2000). Simpson et al. (2008) suggest to use the ratio of precursor
12 sterols to their stanol and stanone degradation products as measure for their degree of
13 degradation.

14 **4.3.3 Cutin and suberin monomers**

15 Bull et al (2000) observed different degradation rates for different components within the
16 classes of free and ester bound lipids, depending on soil chemical and physical composition.
17 However, Otto and Simpson (2006) found degradation of cutin and suberin to take place
18 without preference for specific constituents. In general, Quenea et al. (2004) described cutin
19 and suberin to be more resistant to degradation than free lipids residing in the same particle
20 size fraction.

21 In a study of hydrolysable lipids using compound-specific ^{13}C analysis, Feng et al. (2010)
22 described mean turnover times for cutin and suberin derived ester-bound lipids of 32-34
23 years. While slower than for bulk soil organic matter in this system, it was much shorter than
24 anticipated, leading them to conclude that a large portion of cutin and suberin derived
25 compounds reside in the non-hydrolysable fraction (Feng et al., 2010).

26 As mentioned earlier (section 4.1), Simpson et al. (2008) observed preferential enrichment of
27 suberin monomers with respect to cutin monomers, which was confirmed by Mendez-Millan
28 et al. (2010). In addition to the physical location of suberin versus cutin as potential cause,
29 Simpson et al. (2008) suggested a higher resistance of suberin to degradation than cutin
30 owing to a larger content of phenolic units in the former. Mendez-Millan et al. (2010) argued
31 that microbial degradation, potentially influenced by the access to degradation sites are other
32 factors influencing the slower turnover of suberin vs. cutin monomers. Regardless of the

1 mechanism, the general difference in root vs. aboveground biomass derived suberin and
2 cutinin monomers and their individual turnover would clearly influence the application of the
3 cutin/suberin monomer ratio as proxy for leaf vs. root input.

4 **4.4 Conclusions and implications regarding differences in transformations** 5 **and turnover of molecular proxies in soils**

6 Although available data is limited, it is clear that degradation of organic matter at a molecular
7 level in terrestrial archives such as soils, paleosols and sediments can significantly influence
8 the applicability of molecular proxies. As a result it seems useful to explore the possibility for
9 a correction to improve the determination of paleovegetation and vegetation shifts and other
10 paleoenvironmental information like paleotemperature and pH. The number of published
11 approaches to compensate for the influence of degradation on paleoenvironmental
12 reconstructions is still small. Zech et al. (2009) provided a simple two endmember model
13 approach to improve paleovegetation reconstruction based on molecular ratios of different
14 long-chain *n*-alkanes (C₂₇-C₃₃). Assuming that forest vegetation is dominated by *n*-C₂₇ alkane
15 and grass vegetation by *n*-C₃₁ and *n*-C₃₃ alkanes, high relative contributions of the respective
16 homologues of the assumed source vegetation are used as end-members. At the same time the
17 source vegetation is typically characterized by high odd-over-even predominance of long-
18 chain *n*-alkanes. On the other hand, soils reveal a low odd-over-even predominance and
19 abovementioned molecular ratios with smaller differences between the different vegetation
20 types. In theory, the degradation continuum from plant leaves to soils of the respective
21 vegetation type thus enable the identification of the degradation intensity of an unknown
22 sample, if the sample is mainly influenced by a single vegetation. If the unknown sample
23 does not plot on the degradation continuum, but between the different lines of different
24 vegetation types, the relative contribution of grass vs. tree derived vegetation might be
25 estimated and also corrected for the vegetation.

26 A slightly different approach was established by Buggle et al. (2010) who also used long-
27 chain *n*-alkane ratios and the odd-over-even predominance of alkanes for their correction.
28 While Zech et al. (2009) used correlations and then graphical-based reconstructions, Buggle
29 et al. (2010) used a calculation based approach. The degradation in the continuum from
30 recent soils is taken as an analogy and the slope of the regression line is multiplied with the
31 odd-over-even predominance and the addition of the intercept of a long-chain *n*-alkane ratio
32 in the crossplot of the ratio with the odd-over-even predominance. By moving the regression

1 line to an ancient sample set, the end of the regression line yields the former topsoil value of
2 the molecular ratio and odd-over-even predominance. Variation in the corrected long-chain *n*-
3 alkane ratio enable the assessment of fluctuations in palaeovegetation.

4 Both mentioned approaches rely on the general differentiation of grass vs. forest vegetation
5 based on long-chain *n*-alkane composition. As mentioned above such clear distinction of
6 vegetation types exclusively based on compounds deriving from one compound fraction such
7 as *n*-alkanes might be hampered by various factors such as variability within and between
8 plant species, thus leading to similar composition of e.g. *n*-alkanes from coniferous trees and
9 grass plants (Maffei, 1996b; Maffei et al., 2004). Thus, such simple approaches might be
10 appropriate only in very well defined settings, where independent records such as pollen data
11 confirm the composition of specific plant assemblages determined by molecular proxies.

12 The expansion of approaches like the ones mentioned here to a broader range of molecular
13 proxies is required to receive more complete pictures and to acknowledge the different
14 turnover and degradation of different substance classes. However, the availability of datasets
15 on plant and soil chemical composition for substance classes other than the *n*-alkanes are
16 quite limited, hindering such expanding approaches. Thus, further surveys are required for
17 other molecular proxies than *n*-alkanes for a high diversity of plants and soils from different
18 climates. Afterwards, combined studies of more than one substance class enable improved
19 paleoenvironmental reconstructions, whereas cross-checking with other non-molecular
20 proxies, e.g. fossil pollen data, might be essential, especially if the paleorecord is targeted.
21 Also the extrapolation of such approaches to different environmental and climatic settings
22 might be limited as the effects of temperature, moisture, oxygen availability and others
23 influence the degradation of organic matter as discussed above. Consequently, proper
24 modelling approaches are required to assess not only palaeoenvironmental changes, but also
25 to acknowledge and identify degradation of organic matter at a molecular scale.

26

27 **5 General conclusions**

28 In this review we considered the three most important constraining factors for the application
29 of molecular proxies in soil science: i) variability in the molecular composition of plant
30 derived organic matter as a result of genetic or life stage variations or external environmental
31 factors; ii) variability in (relative contribution of) input pathways into the soil; and iii)
32 transformation and/or (selective) degradation of (some of) the molecules once present in the

1 soil. From the various studies done within and outside of soil science over the last decades
2 the following general picture emerges. All constraining factors considered can have a
3 significant influence on the applicability of molecular proxies in soil science. The degree of
4 influence of the constraining factors strongly depends on the type of molecular proxy as well
5 as the environmental context in which it is applied. In addition, the research question to be
6 addressed by application of the molecular proxy has a strong influence. A factor that poses a
7 constraining factor in one study might offer an opportunity in another. For instance fire
8 induced alteration of biomass may release lipids to the soil that potentially confound their
9 chemotaxonomic application, but may offer opportunities for reconstruction of the
10 occurrence of human induced fire in an archaeological context. Recently, the first modelling
11 approaches to potentially compensate for some of the constraining factors, specifically
12 variability in input pathways and degradation of molecular proxies once in the soil, have
13 started to emerge. Based on the previous we strongly recommend that the potential
14 constraining factors are always explicitly considered whenever studies are planned in which
15 molecular proxies in soils play a role. This review may serve as starting point for gathering
16 the necessary information to decide, which constraining factors may play a role and how they
17 can be addressed best. At the same time, it became clear from available literature that much
18 information about the mentioned constraining factors is still lacking. In particular for
19 molecular classes other than *n*-alkanes, systematic information is often very scarce. We
20 therefore strongly appeal to the soil scientific community to address this knowledge gap. Also
21 for this our review may serve as a starting point with future applicability in soil science and
22 furthermore in paleopedology.

23

24 **References**

25 Ainsworth, E. A. and Long, S. P.: What have we learned from 15 years of free-air CO₂
26 enrichment (FACE)? A meta-analytic review of the responses of photosynthesis, canopy,
27 New Phytol., 165, 351–371, 2005.

28 Almendros, G., Tinoco, P., González-Vila, F.J., Lüdemann, H.D., Sanz, J. and Velasco, F.:
29 ¹³C NMR of forest soil lipids, Soil Sci., 166, 186–196, 2001.

30 Amblès, A., Jambu, P., Jacquesy, J. C., Parlanti, E. and Secouet, B.: Changes in the ketone
31 portion of lipidic components during the decomposition of plant debris in a
32 hydromorphic forest Podzol, Soil Sci., 156, 49–56, 1993.

- 1 Amelung, W., Brodowski, S., Sandhage-Hofmann, A. and Bol, R.: Combining biomarker
2 with stable isotope analyses for assessing the transformation and turnover of soil organic
3 matter, *Adv. Agron.*, Vol 100, 155–250, 2008.
- 4 Andersson, R. A., Kuhry, P., Meyers, P., Zebuhr, Y., Crill, P. and Morth, M.: Impacts of
5 paleohydrological changes on *n*-alkane biomarker compositions of a Holocene peat
6 sequence in the eastern European Russian Arctic, *Org. Geochem.*, 42, 1065–1075, 2011.
- 7 Andreetta, A., Dignac, M. and Carnicelli, S.: Biological and physico-chemical processes
8 influence cutin and suberin biomarker distribution in two Mediterranean forest soil
9 profiles, *Biogeochemistry*, 112, 41–58, 2013.
- 10 Angst, G., John, S., Mueller, C.W., Kögel-Knabner, I. and Rethemeyer, J.: Tracing the
11 sources and spatial distribution of organic carbon in subsoils using a multi-biomarker
12 approach, *Sci. Rep. UK*, 29478, 2016.
- 13 Avato, P., Bianchi, G. and Mariani, G.: Epicuticular waxes of *Sorghum* and some
14 compositional changes with plant age, *Phytochemistry*, 23, 2843–2846, 1984.
- 15 Ayari, A., Yang, H., Wiesenberg, G. L. B. and Xie, S.: Distribution of archaeal and bacterial
16 tetraether membrane lipids in rhizosphere-root systems in soils and their implication for
17 paleoclimate assessment, *Geochem. J.*, 47, 337–347, 2013.
- 18 Baas, M., Pancost, R., Van Geel, B. and Damste, J. S. S.: A comparative study of lipids in
19 *Sphagnum* species, *Org. Geochem.*, 31, 535–541, 2000.
- 20 Bai, Y., Fang, X. M., Wang, Y. L., Kenig, F., Miao, Y. F. and Wang, Y. X.: Distribution of
21 aliphatic ketones in Chinese soils: Potential environmental implications, *Org. Geochem.*,
22 37, 860–869, 2006.
- 23 Bai, Y., Fang, X., Gleixner, G. and Mügler, I.: Effect of precipitation regime on δD values of
24 soil *n*-alkanes from elevation gradients – Implications for the study of paleo-elevation,
25 *Org. Geochem.*, 42, 838–845, 2011.
- 26 Bai, Y., Fang, X., Nie, J., Wang, Y. and Wu, F.: A preliminary reconstruction of the
27 paleoecological and paleoclimatic history of the Chinese Loess Plateau from the
28 application of biomarkers, *Palaeogeogr. Palaeoclimatol.*, 271, 161–169, 2009.
- 29 Baker, E. A. and Hunt, G. M.: Developmental changes in leaf epicuticular waxes in relation
30 to foliar penetration, *New Phytol.*, 88, 731–747, 1981.
- 31 Bakht, J., Bano, A. and Dominy, P.: The role of abscisic acid and low temperature in
32 chickpea (*Cicer arietinum*) cold tolerance. II. Effects on plasma membrane structure and
33 function, *J. Exp. Bot.*, 57, 3707–3715, 2006.

- 1 Berke, M. A., Johnson, T. C., Werne, J. P., Grice, K., Schouten, S. and Sinninghe Damsté, J.
2 S.: Molecular records of climate variability and vegetation response since the Late
3 Pleistocene in the Lake Victoria basin, East Africa, *Quat. Sci. Rev.*, 55, 59–74, 2012.
- 4 Bettaieb, I., Bourgou, S., Wannas, W. A., Hamrouni, I., Limam, F. and Marzouk, B.:
5 Essential oils, phenolics, and antioxidant activities of different parts of cumin (*Cuminum*
6 *cuminum* L.), *J. Agric. Food Chem.*, 58, 10410–10418, 2010.
- 7 Birk, J. J., Dippold, M., Wiesenberg, G. L. B. and Glaser, B.: Combined quantification of
8 faecal sterols, stanols, stanones and bile acids in soils and terrestrial sediments by gas
9 chromatography–mass spectrometry, *J. Chrom. A*, 1242, 1–10, 2012.
- 10 Bol, R., Eriksen, J., Smith, P., Garnett, M. H., Coleman, K. and Christensen, B. T.: The
11 natural abundance of ^{13}C , ^{15}N , ^{34}S and ^{14}C in archived (1923–2000) plant and soil
12 samples from the Askov long-term experiments on animal manure and mineral fertilizer,
13 *Rapid Commun. Mass Spectrom.*, 19, 3216–3226, 2005.
- 14 Bondada, B. R., Oosterhuis, D. M., Murphy, J. B. and Kim, K. S.: Effect of water stress on
15 the epicuticular wax composition and ultrastructure of cotton (*Gossypium hirsutum* L.)
16 leaf, bract, and boll, *Environ. Exp. Bot.*, 36, 61–69, 1996.
- 17 Brennan, C. W., Verhaak, R. G. W., McKenna, A., Campos, B., Noushmehr, H., Salama, S.
18 R., Zheng, S., Chakravarty, D., Sanborn, J. Z., Berman, S. H., Beroukhim, R., Bernard,
19 B., Wu, C., Genovese, G., Shmulevich, I., Barnholtz-Sloan, J., Zou, L., Vegesna, R.,
20 Shukla, S. A., Ciriello, G., Yung, W. K., Zhang, W., Sougnez, C., Mikkelsen, T.,
21 Aldape, K., Bigner, D. D., Van Meir, E. G., Prados, M., Sloan, A., Black, K. L.,
22 Eschbacher, J., Finocchiaro, G., Friedman, W., Andrews, D. W., Guha, A., Iacocca, M.,
23 O'Neill, B. P., Foltz, G., Myers, J., Weisenberger, D. J., Penny, R., Kucherlapati, R.,
24 Perou, C. M., Hayes, D. N., Gibbs, R., Marra, M., Mills, G. B., Lander, E., Spellman, P.,
25 Wilson, R., Sander, C., Weinstein, J., Meyerson, M., Gabriel, S., Laird, P. W., Haussler,
26 D., Getz, G., Chin, L. and TCGA Res Network: The somatic genomic landscape of
27 glioblastoma, *Cell*, 155, 462–477, 2013.
- 28 Buggle, B., Wiesenberg, G. L. B. and Glaser, B.: Is there a possibility to correct fossil *n*-
29 alkane data for postsedimentary alteration effects?, *Appl. Geochem.*, 25, 947–957, 2010.
- 30 Bull, I. D., Nott, C. J., Van Bergen, P. F., Poulton, P. R. and Evershed, R. P.: Organic
31 geochemical studies of soils from the Rothamsted Classical Experiments - VI. The
32 occurrence and source of organic acids in an experimental grassland soil, *Soil Biol.*
33 *Biochem.*, 32, 1367–1376, 2000.

- 1 Bull, I. D., Simpson, I. A., Dockrill, S. J. and Evershed, R. P.: Organic geochemical evidence
2 for the origin of ancient anthropogenic soil deposits at Tofts Ness, Sanday, Orkney, Org.
3 *Geochem.*, 30, 535–556, 1999.
- 4 Bush, R. T. and McInerney, F. A.: Leaf wax *n*-alkane distributions in and across modern
5 plants: Implications for paleoecology and chemotaxonomy, *Geochim. Cosmochim. Acta*,
6 117, 161–179, 2013.
- 7 Canadell, J., Jackson, R. B., Ehleringer, J. R., Mooney, H. A., Sala, O. E. and Schulze, E. D.:
8 Maximum rooting depth of vegetation types at the global scale, *Oecologia*, 108, 583–
9 595, 1996.
- 10 Castañeda, I. S. and Schouten, S.: A review of molecular organic proxies for examining
11 modern and ancient lacustrine environments, *Quat. Sci. Rev.*, 30, 2851–2891, 2011.
- 12 Cayet, C. and Lichtfouse, E.: $\delta^{13}\text{C}$ of plant-derived *n*-alkanes in soil particle-size fractions,
13 *Org. Geochem.*, 32, 253–258, 2001.
- 14 Chaffee, A. L., Hoover, D. S., Johns, R. B. and Schweighardt, F. K.: Biological markers
15 extractable from coal, in: *Biological markers in the sedimentary record*, Johns, R. B.
16 (Ed.), Elsevier, Amsterdam, 311–345, 1986.
- 17 Clemente, J. S., Simpson, A. J. and Simpson, M. J.: Association of specific organic matter
18 compounds in size fractions of soils under different environmental controls, *Org.*
19 *Geochem.*, 42, 1169–1180, 2011.
- 20 Clemente, Z., Castro, V. L. S. S., Moura, M. A. M., Jonsson, C. M. and Fraceto, L. F.:
21 Toxicity assessment of TiO₂ nanoparticles in zebrafish embryos under different exposure
22 conditions, *Aquat. Toxicol.*, 147, 129–139, 2014.
- 23 Concheri, G., Bertoldi, D., Polone, E., Otto, S., Larcher, R. and Squartini, A.: Chemical
24 elemental distribution and soil DNA fingerprints provide the critical evidence in murder
25 case investigation, *Plos One*, 6, e20222, 2011.
- 26 Conte, M. H. and Weber, J. C.: Plant biomarkers in aerosols record isotopic discrimination of
27 terrestrial photosynthesis, *Nature*, 417, 639–641, 2002.
- 28 Conte, M. H., Weber, J. C., Carlson, P. J. and Flanagan, L. B.: Molecular and carbon isotopic
29 composition of leaf wax in vegetation and aerosols in a northern prairie ecosystem,
30 *Oecologia*, 135, 67–77, 2003.
- 31 Conte, P., Spaccini, R., Chiarella, M. and Piccolo, A.: Chemical properties of humic
32 substances in soils of an Italian volcanic system, *Geoderma*, 117, 243–250, 2003.
- 33 Cowlshaw, M. G., Bickerstaffe, R. and Young, H.: Epicuticular wax of four species of
34 *Chionochloa*, *Phytochemistry*, 22, 119–124, 1983.

- 1 Cui Jingwei, Huang Junhua, Meyers, P. A., Huang Xianyu, Li Jingjing and Liu Wengui:
2 Variation in solvent-extractable lipids and *n*-alkane compound-specific carbon isotopic
3 compositions with depth in a Southern China karst area Soil, J. Earth Sci., 21, 382–391,
4 2010.
- 5 Curiale, J. A.: A review of the occurrences and causes of migration-contamination in crude
6 oil, Org. Geochem., 33, 1389–1400, 2002.
- 7 D’Anjou, R. M., Bradley, R. S., Balascio, N. L. and Finkelstein, D. B.: Climate impacts on
8 human settlement and agricultural activities in northern Norway revealed through
9 sediment biogeochemistry, P. Natl Acad. Sci. USA, 2012.
- 10 De Jonge, C., Hopmans, E. C., Zell, C. I., Kim, J., Schouten, S. and Sinninghe Damsté, J. S.:
11 Occurrence and abundance of 6-methyl branched glycerol dialkyl glycerol tetraethers in
12 soils: Implications for palaeoclimate reconstruction, Geochim. Cosmochim. Acta, 141,
13 97–112, 2014.
- 14 De Leeuw, J. W. and Baas, M.: Early-stage diagenesis of steroids, in: Biological markers in
15 the sedimentary record, Johns, R. B. (Ed.), Elsevier, Amsterdam, 101–123, 1986.
- 16 Diefendorf, A.F. and Freimuth, E.J.: Extracting the most from terrestrial plant-derived *n*-alkyl
17 lipids and their carbon isotopes from the sedimentary record: A review, Org. Geochem.
18 103, 1–21, 2017.
- 19 Diné, H., Schnitzer, M. and Meyhus, G. R.: Soil lipids: Origin, nature, content,
20 decomposition, and effect on soil physical properties., in: Soil Biochemistry, Bollag, J.
21 M. and Stotzky, G. (Eds.), Marcel Dekker Inc., New York, 397–429, 1990.
- 22 Dove, H., Mayes, R. and Freer, M.: Effects of species, plant part, and plant age on the *n*-
23 alkane concentrations in the cuticular wax of pasture plants., Aust. J. Agric. Res., 47,
24 1333–1347, 1996.
- 25 Dyson, W. G. and Herbin, G. A.: Variation in leaf wax alkanes in cypress trees grown in
26 Kenya, Phytochemistry, 9, 585–589, 1970.
- 27 Eckmeier, E. and Wiesenberg, G. L. B.: Short-chain *n*-alkanes (C_{16–20}) in ancient soil are
28 useful molecular markers for prehistoric biomass burning, J. Arch. Sci., 36, 1590–1596,
29 2009.
- 30 Eganhouse, R. P.: Molecular markers and environmental organic geochemistry: An overview,
31 in: ACS Symposium Series 671: Molecular Markers in Environmental Geochemistry,
32 Eganhouse, R. P. (Ed.), American Chemical Society, Washington, USA, 1–20, 1997.

- 1 Eglinton, G., Gonzalez, A. G., Hamilton, R. J. and Raphael, R. A.: Hydrocarbon constituents
2 of the wax coatings of plant leaves: A taxonomic survey, *Phytochemistry*, 1, 89–102,
3 1962.
- 4 Eglinton, G., Hamilton, R. J. and Martin-Smith, M.: The alkane constituents of some New
5 Zealand plants and their possible taxonomic implications, *Phytochemistry*, 1, 137–145,
6 1962.
- 7 Ernst, N., Peterse, F., Breitenbach, S. F. M., Syiemlieh, H. J. and Eglinton, T. I.: Biomarkers
8 record environmental changes along an altitudinal transect in the wettest place on Earth,
9 *Org. Geochem.*, 60, 93–99, 2013.
- 10 FAO: Guidelines for soil description - 4th edition, Food and Agricultural Organization of the
11 United Nations, Rome, Italy., 2006.
- 12 Feng, X. and Simpson, M. J.: The distribution and degradation of biomarkers in Alberta
13 grassland soil profiles, *Org. Geochem.*, 38, 1558–1570, 2007.
- 14 Feng, X., Xu, Y., Jaffé, R., Schlesinger, W. H. and Simpson, M. J.: Turnover rates of
15 hydrolysable aliphatic lipids in Duke Forest soils determined by compound specific ¹³C
16 isotopic analysis, *Org. Geochem.*, 41, 573–579, 2010.
- 17 Flessa, H., Amelung, W., Helfrich, M., Wiesenberg, G. L. B., Gleixner, G., Brodowski, S.,
18 Rethemeyer, J., Kramer, C. and Grootes, P. M.: Storage and stability of organic matter
19 and fossil carbon in a Luvisol and Phaeozem with continuous maize cropping: A
20 synthesis, *J. Plant Nutr. Soil Sc.*, 171, 36–51, 2008.
- 21 Frostegard, A. and Bååth, E.: The use of phospholipid fatty acid analysis to estimate bacterial
22 and fungal biomass in soil, *Biol. Fertility Soils*, 22, 59–65, 1996.
- 23 Gleixner, G.: Soil organic matter dynamics: a biological perspective derived from the use of
24 compound-specific isotopes studies, *Ecol. Res.*, 28, 683–695, 2013.
- 25 Gocke, M. I., Kessler, F., Van Mourik, J. M., Jansen, B. and Wiesenberg, G. L. B.: Paleosols
26 can promote root growth of the recent vegetation – a case study from the sandy soil-
27 sediment sequence Rakt, Netherlands, *SOIL* 2, 537–549, 2016.
- 28 Gocke, M., Kuzyakov, Y. and Wiesenberg, G. L. B.: Differentiation of plant derived organic
29 matter in soil, loess and rhizoliths based on *n*-alkane molecular proxies,
30 *Biogeochemistry*, 112, 23–40, 2013.
- 31 Gocke, M., Peth, S. and Wiesenberg, G. L. B.: Lateral and depth variation of loess organic
32 matter overprint related to rhizoliths - Revealed by lipid molecular proxies and X-ray
33 tomography, *Catena*, 112, 72–85, 2014.

- 1 González-Pérez, J.A., González-Vila, F.J., Almendros, G. and Knicker, H.: The effect of fire
2 on soil organic matter - a review, *Env. Internat.*, 30, 855–870, 2004.
- 3 Grandmougin-Ferjani, A., Dalpé, Y., Hartmann, M., Laruelle, F. and Sancholle, M.: Sterol
4 distribution in arbuscular mycorrhizal fungi, *Phytochemistry*, 50, 1027–1031, 1999.
- 5 Griepentrog, M., Bodé, S., Boeckx, P., Hagedorn, F., Heim, A. and Schmidt, M. W. I.:
6 Nitrogen deposition promotes the production of new fungal residues but retards the
7 decomposition of old residues in forest soil fractions, *Global Change Biol.*, 20, 327–340,
8 2014.
- 9 Griepentrog, M., Bodé, S., Boeckx, P. and Wiesenberg, G.L.B.: The fate of plant wax lipids
10 in a model forest ecosystem under elevated CO₂ concentration and increased nitrogen
11 deposition, *Org. Geochem.*, 98, 131–140, 2016.
- 12 Griepentrog, M., Eglinton, T. I., Hagedorn, F., Schmidt, M. W. I. and Wiesenberg, G. L. B.:
13 Interactive effects of elevated CO₂ and nitrogen deposition on fatty acid molecular and
14 isotope composition of above- and belowground tree biomass and forest soil fractions,
15 *Global Change Biol.*, 21, 473–486, 2015.
- 16 Guggenberger, G., Zech, W. and Thomas, R. J.: Lignin and carbohydrate alteration in
17 particle-size separates of an Oxisol under tropical pastures following native savanna,
18 *Soil Biol. Biochem.*, 27, 1629–1638, 1995.
- 19 Gülz, P. G. and Boor, G.: Seasonal variations in epicuticular wax ultrastructures of *Quercus*
20 *robur* leaves, *Z. Naturforsch. C*, 47, 807–814, 1992.
- 21 Gülz, P. G. and Muller, E.: Seasonal variation in the composition of epicuticular waxes of
22 *Quercus robur* leaves, *Z. Naturforsch. C*, 47, 800–806, 1992.
- 23 Gülz, P. G., Muller, E. and Prasad, R. B. N.: Developmental and seasonal variations in the
24 epicuticular waxes of *Tilia tomentosa* leaves, *Phytochemistry*, 30, 769–773, 1991.
- 25 Hamer, U., Rumpel, C. and Dignac, M. F.: Cutin and suberin biomarkers as tracers for the
26 turnover of shoot and root derived organic matter along a chronosequence of Ecuadorian
27 pasture soils, *Eur. J. Soil Sci.*, 63, 808–819, 2012.
- 28 Hamrouni, I., Ben Salah, H. and Marzouk, B.: Effects of water-deficit on lipids of safflower
29 aerial parts, *Phytochemistry*, 58, 277–280, 2001.
- 30 Hautevelle, Y., Michels, R., Malartre, F. and Trouiller, A.: Vascular plant biomarkers as
31 proxies for palaeoflora and palaeoclimatic changes at the Dogger/Malm transition of the
32 Paris Basin (France), *Org. Geochem.*, 37, 610–625, 2006.
- 33 Heim, A. and Schmidt, M. W. I.: Lignin turnover in arable soil and grassland analysed with
34 two different labelling approaches, *Eur. J. Soil Sci.*, 58, 599–608, 2007.

- 1 Hellgren, L. I. and Sandelius, A. S.: Age-dependent variation in membrane lipid synthesis in
2 leaves of garden pea (*Pisum sativum* L.), *J. Exp. Bot.*, 52, 2275–2282, 2001.
- 3 Herbin, G. A. and Robins, P. A.: Patterns of variation and development in leaf wax alkanes,
4 *Phytochemistry*, 8, 1985–1998, 1969.
- 5 Herbin, G. A. and Robins, P. A.: Studies on plant cuticular waxes—I: The chemotaxonomy
6 of alkanes and alkenes of the genus *Aloe* (Liliaceae), *Phytochemistry*, 7, 239–255, 1968.
- 7 Herbin, G. A. and Sharma, K.: Studies on plant cuticular waxes—V. The wax coatings of
8 pine needles: A taxonomic survey, *Phytochemistry*, 8, 151–160, 1969.
- 9 Hermann, N., Boom, A., Carr, A.S., Chase, B.M., West, A.G., Zabel, M. and Schefuss, E.:
10 Hydrogen isotope fractionation of leaf wax *n*-alkanes in southern African soils, *Org.*
11 *Geochem.*, 109, 1–13, 2017.
- 12 Höfle, S., Rethemeyer, J., Mueller, C. W. and John, S.: Organic matter composition and
13 stabilization in a polygonal tundra soil of the Lena Delta, *Biogeosciences*, 10, 3145–
14 3158, 2013.
- 15 Huang, X., Wang, C., Zhang, J., Wiesenberg, G. L. B., Zhang, Z. and Xie, S.: Comparison of
16 free lipid compositions between roots and leaves of plants in the Dajiuhu Peatland,
17 central China, *Geochem. J.*, 45, 365–373, 2011.
- 18 Huang, Y., Bol, R., Harkness, D. D., Ineson, P. and Eglinton, G.: Post-glacial variations in
19 distributions, ¹³C and ¹⁴C contents of aliphatic hydrocarbons and bulk organic matter in
20 three types of British acid upland soils, *Org. Geochem.*, 24, 273–287, 1996.
- 21 Huang, Y., Eglinton, G., Ineson, P., Bol, R. and Harkness, D. D.: The effects of nitrogen
22 fertilisation and elevated CO₂ on the lipid biosynthesis and carbon isotopic
23 discrimination in birch seedlings (*Betula pendula*), *Plant Soil*, 216, 35–45, 1999.
- 24 Jackson, R. B., Canadell, J., Ehleringer, J. R., Mooney, H. A., Sala, O. E. and Schulze, E. D.:
25 A global analysis of root distributions for terrestrial biomes, *Oecologia*, 108, 389–411,
26 1996.
- 27 Jackson, R. B., Mooney, H. A. and Schulze, E. D.: A global budget for fine root biomass,
28 surface area, and nutrient contents, *Proc. Natl. Acad. Sci. U. S. A.*, 94, 7362–7366, 1997.
- 29 Jaffe, R., Elisme, T. and Cabrera, A. C.: Organic geochemistry of seasonally flooded rain
30 forest soils: Molecular composition and early diagenesis of lipid components, *Org.*
31 *Geochem.*, 25, 9–17, 1996.
- 32 Jambu, P., Fustec, E. and Jacquesy, R.: Les lipides des sols: nature, origine, evolution,
33 propriétés, *Science du Sol*, 4, 229–240, 1978.

- 1 Jansen, B. and Nierop, K. G. J.: Me-ketones in high altitude Ecuadorian Andisols confirm
2 excellent conservation of plant-specific *n*-alkane patterns, *Org. Geochem.*, 40, 61–69,
3 2009.
- 4 Jansen, B., Nierop, K. G. J., Hageman, J. A., Cleef, A. and Verstraten, J. M.: The straight-
5 chain lipid biomarker composition of plant species responsible for the dominant biomass
6 production along two altitudinal transects in the Ecuadorian Andes, *Org. Geochem.*, 37,
7 1514–1536, 2006.
- 8 Jansen, B., Nierop, K. G. J., Tonneijck, F. H., Van der Wielen, F. W. M. and Verstraten, J.
9 M.: Can isoprenoids in leaves and roots of plants along altitudinal gradients in the
10 Ecuadorian Andes serve as biomarkers?, *Plant Soil*, 291, 181–198, 2007.
- 11 Jansen, B., Van Loon, E. E., Hooghiemstra, H. and Verstraten, J. M.: Improved
12 reconstruction of palaeo-environments through unravelling of preserved vegetation
13 biomarker patterns, *Palaeogeogr. Palaeoclimatol.*, 285, 119–130, 2010.
- 14 Jansen, B., de Boer, E. J., Cleef, A. M., Hooghiemstra, H., Moscol-Olivera, M., Tonneijck, F.
15 H. and Verstraten, J. M.: Reconstruction of late Holocene forest dynamics in northern
16 Ecuador from biomarkers and pollen in soil cores, *Palaeogeogr. Palaeoclimatol.*, 386, 607–619,
17 2013.
- 18 Kell, D. B.: Large-scale sequestration of atmospheric carbon via plant roots in natural and
19 agricultural ecosystems: why and how?, *Philos. T. Roy. Soc. B.*, 367, 1589–1597, 2012.
- 20 Killops, S. D. and Frewin, N. L.: Triterpenoid Diagenesis and Cuticular Preservation, *Org.*
21 *Geochem.*, 21, 1193–1209, 1994.
- 22 Kim, K. S., Park, S. H. and Jenks, M. A.: Changes in leaf cuticular waxes of sesame
23 (*Sesamum indicum* L.) plants exposed to water deficit, *J. Plant Physiol.*, 164, 1134–1143,
24 2007.
- 25 Kindler, R., Miltner, A., Thullner, M., Richnow, H. and Kaestner, M.: Fate of bacterial
26 biomass derived fatty acids in soil and their contribution to soil organic matter, *Org.*
27 *Geochem.*, 40, 29–37, 2009.
- 28 Kirkels, F. M., Jansen, B. and Kalbitz, K.: Consistency of plant-specific *n*-alkane patterns in
29 pluggen ecosystems: A review, *Holocene*, 23, 1355–1368, 2013.
- 30 Knicker, H., Hilscher, A., De la Rosa, J.M., González-Pérez, J.A. and González-Vila, F.J.:
31 Modification of biomarkers in pyrogenic organic matter during the initial phase of
32 charcoal biodegradation in soils, *Geoderma*, 197–198, 43–50, 2013.

- 1 Koch, B., Harder, J., Lara, R. J. and Kattner, G.: The effect of selective microbial degradation
2 on the composition of mangrove derived pentacyclic triterpenols in surface sediments,
3 *Org. Geochem.*, 36, 273–285, 2005.
- 4 Koch, K., Hartmann, K. D., Schreiber, L., Barthlott, W. and Neinhuis, C.: Influences of air
5 humidity during the cultivation of plants on wax chemical composition, morphology and
6 leaf surface wettability, *Environ. Exp. Bot.*, 56, 1–9, 2006.
- 7 Kögel-Knabner, I.: The macromolecular organic composition of plant and microbial residues
8 as inputs to soil organic matter, *Soil Biol. Biochem.*, 34, 139–162, 2002.
- 9 Kramer, C. and Gleixner, G.: Variable use of plant- and soil-derived carbon by
10 microorganisms in agricultural soils, *Soil Biol. Biochem.*, 38, 3267–3278, 2006.
- 11 Kreyling, J., Thiel, D., Simmnacher, K., Willner, E., Jentsch, A. and Beierkuhnlein, C.:
12 Geographic origin and past climatic experience influence the response to late spring frost
13 in four common grass species in central Europe, *Ecography*, 35, 268–275, 2012.
- 14 Krull, E., Sachse, D., Mügler, I., Thiele, A. and Gleixner, G.: Compound-specific $\delta^{13}\text{C}$ and
15 $\delta^2\text{H}$ analyses of plant and soil organic matter: A preliminary assessment of the effects of
16 vegetation change on ecosystem hydrology, *Soil Biol. Biochem.*, 38, 3211–3221, 2006.
- 17 Ladygina, N., Dedyukhina, E. G. and Vainshtein, M. B.: A review on microbial synthesis of
18 hydrocarbons, *Process Biochem.*, 41, 1001–1014, 2006.
- 19 Larkindale, J. and Huang, B. R.: Changes of lipid composition and saturation level in leaves
20 and roots for heat-stressed and heat-acclimated creeping bentgrass (*Agrostis stolonifera*),
21 *Environ. Exp. Bot.*, 51, 57–67, 2004.
- 22 Lavrieux, M., Breheret, J., Disnar, J., Jacob, J., Le Milbeau, C. and Zocatelli, R.: Preservation
23 of an ancient grassland biomarker signature in a forest soil from the French Massif
24 Central, *Org. Geochem.*, 51, 1–10, 2012.
- 25 Lavrieux, M., Jacob, J., LeMilbeau, C., Zocatelli, R., Masuda, K., Breheret, J. and Disnar, J.:
26 Occurrence of triterpenyl acetates in soil and their potential as chemotaxonomical
27 markers of Asteraceae, *Org. Geochem.*, 42, 1315–1323, 2011.
- 28 Le Milbeau, C., Lavrieux, M., Jacob, J., Breheret, J., Zocatelli, R. and Disnar, J.: Methoxy-
29 serratenes in a soil under conifers and their potential use as biomarkers of Pinaceae, *Org.*
30 *Geochem.*, 55, 45–54, 2013.
- 31 Lehmann, J. and Kleber, M.: The contentious nature of soil organic matter. *Nature*, 528, 60–
32 68, 2015.

- 1 Lei, B. K., Fan, M. S., Chen, Q., Six, J. and Zhang, F. S.: Conversion of wheat-maize to
2 vegetable cropping systems changes soil organic matter characteristics, *Science du Sol*,
3 74, 1320–1326, 2010.
- 4 Lei, G., Zhang, H., Chang, F., Pu, Y., Zhu, Y., Yang, M. and Zhang, W.: Biomarkers of
5 modern plants and soils from Xinglong Mountain in the transitional area between the
6 Tibetan and Loess Plateaus, *Quat. Internat.*, 218, 143–150, 2010.
- 7 Lejay, M., Alexis, M., Quénéa, K., Sellami, F. and Bon, F.: Organic signature of fireplaces:
8 experimental references for archaeological interpretations, *Org. Geochem.*, 99, 67–77,
9 2016.
- 10 Li, H., Madden, L. J. and Potts, B. M.: Variation in leaf waxes of the Tasmanian *Eucalyptus*
11 species 1. Subgenus *Symphyomyrtus*, *Biochem. Syst. Ecol.*, 25, 631–657, 1997.
- 12 Li, J., Huang, J., Ge, J., Huang, X. and Xie, S.: Chemotaxonomic significance of *n*-alkane
13 distributions from leaf wax in genus of *Sinojackia* species (Styracaceae), *Biochem. Syst.*
14 *Ecol.*, 49, 30–36, 2013.
- 15 Lichtfouse, E.: Heterogeneous turnover of molecular organic substances from crop soils as
16 revealed by ^{13}C labeling at natural abundance with *Zea mays*, *Naturwissenschaften*, 84,
17 23–25, 1997.
- 18 Lichtfouse, E., Berthier, G., Houot, S., Barriuso, E., Bergheud, V. and Vallaeys, T.: Stable
19 carbon isotope evidence for the microbial origin of C_{14} - C_{18} *n*-alkanoic acids in soils,
20 *Org. Geochem.*, 23, 849–852, 1995.
- 21 Lichtfouse, E., Leblond, C., Da Silva, M. and Behar, F.: Occurrence of biomarkers and
22 straight-chain biopolymers in humin: Implication for the origin of soil organic matter,
23 *Naturwissenschaften*, 85, 497–501, 1998.
- 24 Lichtfouse, E.: C-13-dating, the first method to calculate the relative age of molecular
25 substance homologues in soil, *Environ. Chem. Lett.*, 10, 97–103, 2012.
- 26 Lichtfouse, É, Bardoux, G., Mariotti, A., Balesdent, J., Ballentine, D. C. and Macko, S. A.:
27 Molecular, ^{13}C , and ^{14}C evidence for the allochthonous and ancient origin of C_{16} - C_{18} *n*-
28 alkanes in modern soils, *Geochim. Cosmochim. Acta*, 61, 1891–1898, 1997.
- 29 Lichtfouse, É, Chenu, C., Baudin, F., Leblond, C., Da Silva, M., Behar, F., Derenne, S.,
30 Largeau, C., Wehrung, P. and Albrecht, P.: A novel pathway of soil organic matter
31 formation by selective preservation of resistant straight-chain biopolymers: chemical and
32 isotope evidence, *Org. Geochem.*, 28, 411–415, 1998.
- 33 Liebig, J., Merck, E. and Mohr, F.: Das aetherische Oel der Getraide, *Annalen der Pharmacie*,
34 24, 248–251, 1837.

- 1 Liu, W., Yang, H., Ning, Y. and An, Z.: Contribution of inherent organic carbon to the bulk
2 delta C-13 signal in loess deposits from the arid western Chinese Loess Plateau, *Org.*
3 *Geochem.*, 38, 1571–1579, 2007.
- 4 Lockheart, M. J., Poole, I., Van Bergen, P. F. and Evershed, R. P.: Leaf carbon isotope
5 compositions and stomatal characters: important considerations for palaeoclimate
6 reconstructions, *Org. Geochem.*, 29, 1003–1008, 1998.
- 7 Maffei, M.: Discriminant analysis of leaf wax alkanes in the Lamiaceae and four other plant
8 families, *Biochem. Syst. Ecol.*, 22, 711–728, 1994.
- 9 Maffei, M., Badino, S. and Bossi, S.: Chemotaxonomic significance of leaf wax *n*-alkanes in
10 the Pinales (Coniferales), *J. Biol. Res.*, 1, 3–20, 2004.
- 11 Maffei, M.: Chemotaxonomic significance of leaf wax alkanes in the Gramineae, *Biochem.*
12 *Syst. Ecol.*, 24, 53–64, 1996a.
- 13 Maffei, M.: Chemotaxonomic significance of leaf wax *n*-alkanes in the Umbelliferae,
14 Cruciferae and Leguminosae (subf papilionoideae), *Biochem. Syst. Ecol.*, 24, 531–545,
15 1996b.
- 16 Maffei, M., Meregalli, M. and Scannerini, S.: Chemotaxonomic significance of surface wax
17 *n*-alkanes in the Cactaceae, *Biochem. Syst. Ecol.*, 25, 241–253, 1997.
- 18 Maffei, M., Mucciarelli, M. and Scannerini, S.: Environmental-factors affecting the lipid-
19 metabolism in *Rosmarinus officinalis* L., *Biochem. Syst. Ecol.*, 21, 765–784, 1993.
- 20 Malik, A., Blagodatskaya, E. and Gleixner, G.: Soil microbial carbon turnover decreases with
21 increasing molecular size, *Soil Biol. Biochem.*, 62, 115–118, 2013.
- 22 Mambelli, S., Bird, J. A., Gleixner, G., Dawson, T. E. and Torn, M. S.: Relative contribution
23 of foliar and fine root pine litter to the molecular composition of soil organic matter after
24 in situ degradation, *Org. Geochem.*, 42, 1099–1108, 2011.
- 25 Marseille, F., Disnar, J. R., Guillet, B. and Noack, Y.: *n*-Alkanes and free fatty acids in
26 humus and Al horizons of soils under beech, spruce and grass in the Massif-Central
27 (Mont-Lozere), France, *Eur. J. Soil Sci.*, 50, 433–441, 1999.
- 28 Martelanc, M., Vovk, I. and Simonovska, B.: Determination of three major triterpenoids in
29 epicuticular wax of cabbage (*Brassica oleracea* L.) by high-performance liquid
30 chromatography with UV and mass spectrometric detection, *J. Chrom. A*, 1164, 145–
31 152, 2007.
- 32 Matteucci, M., D'Angeli, S., Errico, S., Lamanna, R., Perrotta, G. and Altamura, M. M.: Cold
33 affects the transcription of fatty acid desaturases and oil quality in the fruit of *Olea*
34 *europaea* L. genotypes with different cold hardiness, *J. Exp. Bot.*, 62, 3403–3420, 2011.

- 1 Medina, E., Aguiar, G., Gomez, M., Aranda, J., Medina, J. D. and Winter, K.: Taxonomic
2 significance of the epicuticular wax composition in species of the genus *Clusia* from
3 Panama, *Biochem. Syst. Ecol.*, 34, 319–326, 2006.
- 4 Medina, E., Aguiar, G., Gomez, M. and Medina, J. D.: Patterns of leaf epicuticular waxes in
5 species of *Clusia*: Taxonomical implications, *Interciencia*, 29, 579–582, 2004.
- 6 Mendez-Millan, M., Dignac, M. F., Rumpel, C. and Derenne, S.: Quantitative and qualitative
7 analysis of cutin in maize and a maize-cropped soil: Comparison of CuO oxidation,
8 transmethylation and saponification methods, *Org. Geochem.*, 41, 187–191, 2010.
- 9 Mendez-Millan, M., Dignac, M. - F., Rumpel, C., Rasse, D. P., Bardoux, G. and Derenne, S.:
10 Contribution of maize root derived C to soil organic carbon throughout an agricultural
11 soil profile assessed by compound specific ^{13}C analysis, *Org. Geochem.*, 42, 1502–1511,
12 2012.
- 13 Mendez-Millan, M., Dignac, M. -, Rumpel, C., Rasse, D. P. and Derenne, S.: Molecular
14 dynamics of shoot vs. root biomarkers in an agricultural soil estimated by natural
15 abundance ^{13}C labelling, *Soil Biol. Biochem.*, 42, 169–177, 2010.
- 16 Mendez-Millan, M., Dignac, M., Rumpel, C. and Derenne, S.: Can cutin and suberin
17 biomarkers be used to trace shoot and root-derived organic matter? A molecular and
18 isotopic approach, *Biogeochemistry*, 106, 23–38, 2011.
- 19 Miltner, A., Bombach, P., Schmidt-Bruecken, B. and Kaestner, M.: SOM genesis: microbial
20 biomass as a significant source, *Biogeochemistry*, 111, 41–55, 2012.
- 21 Moingt, M., Lucotte, M. and Paquet, S.: Lignin biomarkers signatures of common plants and
22 soils of Eastern Canada, *Biogeochemistry*, 129, 133–138, 2016.
- 23 Mongrand, S., Badoc, A., Patouille, B., Lacomblez, C., Chavent, M., Cassagne, C. and
24 Bessoule, J. J.: Taxonomy of gymnospermae: multivariate analysis of leaf fatty acid
25 composition, *Phytochemistry*, 58, 101–115, 2001.
- 26 Moseley, G.: Variation in the epicuticular wax content of white and red clover leaves, *Grass*
27 *Forage Sci.*, 38, 201–204, 1983.
- 28 Moucawi, J., Fustec, E., Jambu, P., Ambles, A. and Jacquesy, R.: Biooxidation of added and
29 natural hydrocarbons in soils: effect of iron, *Soil Biol. Biochem.*, 13, 335–342, 1981.
- 30 Moucawi, J., Fustec, E., Jambu, P. and Jacquesy, R.: Decomposition of lipids in soils: Free
31 and esterified fatty acids, alcohols and ketones, *Soil Biol. Biochem.*, 13, 461–468, 1981.
- 32 Mueller-Niggemann, C. and Schwark, L.: Chemotaxonomy and diagenesis of aliphatic
33 hydrocarbons in rice plants and soils from land reclamation areas in the Zhejiang
34 Province, China, *Org. Geochem.*, 83–84, 215–226, 2015.

- 1 Naafs, D. F. W., Van Bergen, P. F., Boogert, S. J. and de Leeuw, J. W.: Solvent-extractable
2 lipids in an acid andic forest soil; variations with depth and season, *Soil Biol. Biochem.*,
3 36, 297–308, 2004.
- 4 Naafs, D. F. W., Van Bergen, P. F., de Jong, M. A., Oonincx, A. and de Leeuw, J. W.: Total
5 lipid extracts from characteristic soil horizons in a Podzol profile, *Eur. J. Soil Sci.*, 55,
6 657–669, 2004.
- 7 Ngosong, C., Gabriel, E. and Ruess, L.: Use of the signature fatty acid 16:1omega5 as a tool
8 to determine the distribution of arbuscular mycorrhizal fungi in soil, *J. Lipids*, 2012,
9 236807–236807, 2012.
- 10 Nguyen Tu, T. T., Derenne, S., Largeau, C., Bardoux, G. and Mariotti, A.: Diagenesis effects
11 on specific carbon isotope composition of plant *n*-alkanes, *Org. Geochem.*, 35, 317–329,
12 2004.
- 13 Nguyen Tu, T. T., Derenne, S., Largeau, C., Mariotti, A. and Bocherens, H.: Comparison of
14 leaf lipids from a fossil ginkgoalean plant and its extant counterpart at two degradation
15 stages: diagenetic and chemotaxonomic implications, *Rev. Palaeobot. Palynol.*, 124, 63–
16 78, 2003.
- 17 Nguyen Tu, T. T., Egasse, C., Zeller, B., Bardoux, G., Biron, P., Ponge, J., David, B. and
18 Derenne, S.: Early degradation of plant alkanes in soils: A litterbag experiment using
19 ¹³C-labelled leaves, *Soil Biol. Biochem.*, 43, 2222–2228, 2011.
- 20 Nierop, K. G. J. and Jansen, B.: Extensive transformation of organic matter and excellent
21 lipid preservation at the upper, superhumid Guandera Páramo, *Geoderma*, 151, 357–369,
22 2009.
- 23 Nierop, K. G. J., Jansen, B., Hageman, J. A. and Verstraten, J. M.: The complementarity of
24 extractable and ester-bound lipids in a soil profile under pine, *Plant Soil*, 286, 269–285,
25 2006.
- 26 Nierop, K. G. J., Van Lagen, B. and Buurman, P.: Composition of plant tissues and soil
27 organic matter in the first stages of a vegetation succession, *Geoderma*, 100, 1–24, 2001.
- 28 Nierop, K. G. J. and Verstraten, J. M.: Rapid molecular assessment of the bioturbation extent
29 in sandy soil horizons under, pine using ester-bound lipids by on-line thermally assisted
30 hydrolysis and methylation-gas chromatography/mass spectrometry, *Rapid Commun.*
31 *Mass Sp.*, 18, 1081–1088, 2004.
- 32 Nierop, K. G. J. and Verstraten, J. M.: Organic matter formation in sandy subsurface horizons
33 of Dutch coastal dunes in relation to soil acidification, *Org. Geochem.*, 34, 499–513,
34 2003.

- 1 Nødskov Giese, B.: Effects of light and temperature on the composition of epicuticular wax
2 of barley leaves, *Phytochemistry*, 14, 921–929, 1975.
- 3 Ohsaki, A., Shibata, K., Kubota, T. and Tokoroyama, T.: Phylogenetic and chemotaxonomic
4 significance of diterpenes in some *Portulaca* species (Portulacaceae), *Biochem. Syst.*
5 *Ecol.*, 27, 289–296, 1999.
- 6 Otto, A., Shunthirasingham, C. and Simpson, M. J.: A comparison of plant and microbial
7 biomarkers in grassland soils from the Prairie Ecozone of Canada, *Org. Geochem.*, 36,
8 425–448, 2005.
- 9 Otto, A. and Simpson, M. J.: Degradation and preservation of vascular plant-derived
10 biomarkers in grassland and forest soils from Western Canada, *Biogeochemistry*, 74,
11 377–409, 2005.
- 12 Otto, A. and Simpson, M. J.: Sources and composition of hydrolysable aliphatic lipids and
13 phenols in soils from western Canada, *Org. Geochem.*, 37, 385–407, 2006.
- 14 Pancost, R. D., Baas, M., Van Geel, B. and Damste, J. S. S.: Biomarkers as proxies for plant
15 inputs to peats: an example from a sub-boreal ombrotrophic bog, *Org. Geochem.*, 33,
16 675–690, 2002.
- 17 Peters, K. E., Walters, C. C. and Moldowan, J. M.: *The Biomarker Guide*, 2 ed., Cambridge
18 University Press, Cambridge, United Kingdom, 2005.
- 19 Peterse, F., Van der Meer, M. T. J., Schouten, S., Jia, G., Ossebaar, J., Blokker, J. and
20 Damste, J. S. S.: Assessment of soil *n*-alkane δD and branched tetraether membrane lipid
21 distributions as tools for paleoelevation reconstruction, *Biogeosciences*, 6, 2799–2807,
22 2009.
- 23 Peterse, F., Nicol, G. W., Schouten, S. and Damste, J. S. S.: Influence of soil pH on the
24 abundance and distribution of core and intact polar lipid-derived branched GDGTs in
25 soil, *Org. Geochem.*, 41, 1171–1175, 2010.
- 26 Peterse, F., Van der Meer, J., Schouten, S., Weijers, J. W. H., Fierer, N., Jackson, R. B., Kim,
27 J. and Damste, J. S. S.: Revised calibration of the MBT-CBT paleotemperature proxy
28 based on branched tetraether membrane lipids in surface soils, *Geochim. Cosmochim.*
29 *Acta*, 96, 215–229, 2012.
- 30 Piervittori, R., Usai, L., Alessio, F. and Maffei, M.: Surface *n*-alkane variability in *Xanthoria*
31 *parietina*, *Lichenologist*, 28, 79–87, 1996.
- 32 Prasad, R. B. N. and Gülz, P. G.: Developmental and seasonal variations in the epicuticular
33 waxes of beech leaves (*Fagus sylvatica* L.), *Z. Naturforsch. C*, 45, 805–812, 1990.

- 1 Prost, K., Birk, J.J., Lehndorff, E., Gerlach, R. and Amelung, W.: Steroid biomarkers
2 revisited - improved source identification of faecal remains in archaeological soil
3 material, PLOS ONE, 12, 1, 2017.
- 4 Quenea, K., Derenne, S., Largeau, C., Rumpel, C. and Mariotti, A.: Variation in lipid relative
5 abundance and composition among different particle size fractions of a forest soil, Org.
6 Geochem., 35, 1355–1370, 2004.
- 7 Quenea, K., Largeau, C., Derenne, S., Spaccini, R., Bardoux, G. and Mariotti, A.: Molecular
8 and isotopic study of lipids in particle size fractions of a sandy cultivated soil (Cestas
9 cultivation sequence, southwest France): Sources, degradation, and comparison with
10 Cestas forest soil, Org. Geochem., 37, 20–44, 2006.
- 11 Rao Z., Wu Y., Zhu Z, Jia G. and Henderson, A.: Is the maximum carbon number of long-
12 chain *n*-alkanes an indicator of grassland or forest? Evidence from surface soils and
13 modern plants, Chinese Sci. Bull., 56, 1714–1720, 2011.
- 14 Rao, Z., Zhu, Z., Jia, G., Henderson, A. C. G., Xue, Q. and Wang, S.: Compound specific δD
15 values of long chain *n*-alkanes derived from terrestrial higher plants are indicative of the
16 δD of meteoric waters: Evidence from surface soils in eastern China, Org. Geochem., 40,
17 922–930, 2009.
- 18 Rasse, D. P., Rumpel, C. and Dignac, M. F.: Is soil carbon mostly root carbon? Mechanisms
19 for a specific stabilisation, Plant Soil, 269, 341–356, 2005.
- 20 Rethemeyer, J., Kramer, C., Gleixner, G., Wiesenberg, G. L. B., Schwark, L., Andersen, N.,
21 Nadeau, M. J. and Grootes, P. M.: Complexity of soil organic matter: AMS ^{14}C analysis
22 of soil lipid fractions and individual compounds, Radiocarbon, 46, 465–473, 2004.
- 23 Rommerskirchen, F., Eglinton, G., Dupont, L., Guntner, U., Wenzel, C. and Rullkotter, J.: A
24 north to south transect of Holocene southeast Atlantic continental margin sediments:
25 Relationship between aerosol transport and compound-specific $\delta^{13}C$ land plant
26 biomarker and pollen records, Geochem. Geophys. Geosci., 4, 1101, 2003.
- 27 Rommerskirchen, F., Plader, A., Eglinton, G., Chikaraishi, Y. and Rullkoetter, J.:
28 Chemotaxonomic significance of distribution and stable carbon isotopic composition of
29 long-chain alkanes and alkan-1-ols in C4 grass waxes, Org. Geochem., 37, 1303–1332,
30 2006.
- 31 Rumpel, C. and Koegel-Knabner, I.: Deep soil organic matter-a key but poorly understood
32 component of terrestrial C cycle, Plant Soil, 338, 143–158, 2011.

- 1 Sachse, D., Radke, J. and Gleixner, G.: δD values of individual *n*-alkanes from terrestrial
2 plants along a climatic gradient - Implications for the sedimentary biomarker record,
3 *Org. Geochem.*, 37, 469–483, 2006.
- 4 Sachse, D., Billault, I., Bowen, G. J., Chikaraishi, Y., Dawson, T. E., Feakins, S. J., Freeman,
5 K. H., Magill, C. R., McInerney, F. A., Van der Meer, M. T. J., Polissar, P., Robins, R.
6 J., Sachs, J. P., Schmidt, H., Sessions, A. L., White, J. W. C., West, J. B. and Kahmen,
7 A.: Molecular paleohydrology: Interpreting the hydrogen isotopic composition of lipid
8 biomarkers from photosynthesizing organisms, *Annu. Rev. Earth and Pl. Sc.*, Vol 40, 40,
9 221–249, 2012.
- 10 Sanchez, F. J., Manzanares, M., de Andres, E. F., Tenorio, J. L. and Ayerbe, L.: Residual
11 transpiration rate, epicuticular wax load and leaf colour of pea plants in drought
12 conditions. Influence on harvest index and canopy temperature, *Eur. J. Agron.*, 15, 57–
13 70, 2001.
- 14 Schefuß, E., Ratmeyer, V., Stuut, J. W., Jansen, J. H. F. and Sinninghe Damsté, J. S.: Carbon
15 isotope analyses of *n*-alkanes in dust from the lower atmosphere over the central eastern
16 Atlantic, *Geochim. Cosmochim. Acta*, 67, 1757–1767, 2003.
- 17 Schmidt, M. W. I., Torn, M. S., Abiven, S., Dittmar, T., Guggenberger, G., Janssens, I. A.,
18 Kleber, M., Koegel-Knabner, I., Lehmann, J., Manning, D. A. C., Nannipieri, P., Rasse,
19 D. P., Weiner, S. and Trumbore, S. E.: Persistence of soil organic matter as an ecosystem
20 property, *Nature*, 478, 49–56, 2011.
- 21 Schouten, S., Hopmans, E.C. and Sinninghe-Damsté, J.S.: The organic geochemistry of
22 glycerol dialkyl glycerol tetraether lipids: A review, *Org. Geochem.*, 54, 19–61, 2013.
- 23 Schreiber, L., Hartmann, K., Skrabs, M. and Zeier, J.: Apoplastic barriers in roots: Chemical
24 composition of endodermal and hypodermal cell walls, *J. Exp. Bot.*, 50, 1267–1280,
25 1999.
- 26 Shelvey, J. D. and Koziol, M. J.: Seasonal and SO₂-induced changes in epicuticular wax of
27 ryegrass, *Phytochemistry*, 25, 415–420, 1986.
- 28 Shepherd, T. and Griffiths, D. W.: The effects of stress on plant cuticular waxes, *New*
29 *Phytol.*, 171, 469–499, 2006.
- 30 Shepherd, T., Robertson, G. W., Griffiths, D. W., Birch, A. N. E. and Duncan, G.: Effects of
31 environment on the composition of epicuticular wax from kale and swede,
32 *Phytochemistry*, 40, 407–417, 1995.

- 1 Simpson, M. J., Otto, A. and Feng, X.: Comparison of solid-state ¹³C nuclear magnetic
2 resonance and organic matter biomarkers for assessing soil organic matter degradation,
3 Soil Sci. Soc. Am. J., 72, 268–276, 2008.
- 4 Simpson, M. J. and Simpson, A. J.: The chemical ecology of soil organic Matter Molecular
5 Constituents, J. Chem. Ecol., 38, 768–784, 2012.
- 6 Sonibare, M. A., Jayeola, A. A. and Egunyomi, A.: Chemotaxonomic significance of leaf
7 alkanes in species of *Ficus* (Moraceae), Biochem. Syst. Ecol., 33, 79–86, 2005.
- 8 Srivastava, K., Jentsch, A., Glaser, B. and Wiesenberg, G. L. B.: Plant and soil *n*-alkane
9 composition is not affected by annual drought in temperate grassland and heathland
10 ecosystems, J. Plant Nutr. Soil Sci., doi: 10.1002/jpln.201600019.
- 11 Stevenson, F. J.: Lipids in soil, J. Am. Oil Chem., 43, 203–210, 1966.
- 12 Stevenson, F. J.: Humus Chemistry: Genesis, Composition, Reactions, John Wiley & Sons,
13 Inc., New York, 1994.
- 14 Stewart, C. E., Neff, J. C., Amatangelo, K. L. and Vitousek, P. M.: Vegetation effects on soil
15 organic matter chemistry of aggregate fractions in a Hawaiian forest, Ecosystems, 14,
16 382–397, 2011.
- 17 Thevenot, M., Dignac, M. and Rumpel, C.: Fate of lignins in soils: A review, Soil Biol.
18 Biochem., 42, 1200–1211, 2010.
- 19 Tissot, B. P. and Welte, D. H.: Petroleum Formation and Occurrence, 2 ed., Springer-Verlag,
20 Berlin, Germany, 1984.
- 21 Tonneijck, F. H., Jansen, B., Nierop, K. G. J., Verstraten, J. M., Sevink, J. and De Lange, L.:
22 Towards understanding of carbon stocks and stabilization in volcanic ash soils in natural
23 Andean ecosystems of northern Ecuador, Eur. J. Soil Sci., 61, 392–405, 2010.
- 24 Tulloch, A. P.: Composition of leaf surface waxes of *Triticum* species: Variation with age
25 and tissue, Phytochemistry, 12, 2225–2232, 1973.
- 26 Van Bon, L., Affandi, A. J., Broen, J., Christmann, R. B., Marijnissen, R. J., Stawski, L.,
27 Farina, G. A., Stifano, G., Mathes, A. L., Cossu, M., York, M., Collins, C., Wenink, M.,
28 Huijbens, R., Hesselstrand, R., Saxne, T., DiMarzio, M., Wuttge, D., Agarwal, S. K.,
29 Reveille, J. D., Assassi, S., Mayes, M., Deng, Y., Drenth, J. P. H., de Graaf, J., den
30 Heijer, M., Kallenberg, C. G. M., Bijl, M., Loof, A., Van den Berg, W. B., Joosten, L. A.
31 B., Smith, V., de Keyser, F., Scorza, R., Lunardi, C., Van Riel, P. L. C. M., Vonk, M.,
32 Van Heerde, W., Meller, S., Homey, B., Beretta, L., Roest, M., Trojanowska, M.,
33 Lafyatis, R. and Radstake, T. R. D. J.: Proteome-wide Analysis and CXCL4 as a
34 Biomarker in Systemic Sclerosis, N. Engl. J. Med., 370, 433–443, 2014.

- 1 Van Mourik, J. M. and Jansen, B.: The added value of biomarker analysis in palaeopedology;
2 reconstruction of the vegetation during stable periods in a polycyclic driftsand sequence
3 in SE-Netherlands, *Quat. Internat.*, 306, 14–23, 2013.
- 4 Vogts, A., Moossen, H., Rommerskirchen, F. and Rullkoetter, J.: Distribution patterns and
5 stable carbon isotopic composition of alkanes and alkan-1-ols from plant waxes of
6 African rain forest and savanna C3 species, *Org. Geochem.*, 40, 1037–1054, 2009.
- 7 Volkman, J. K.: Sterols and other triterpenoids: source specificity and evolution of
8 biosynthetic pathways, *Org. Geochem.*, 36, 139–159, 2005.
- 9 Von Lützow, M., Kögel-Knabner, I., Ludwig, B., Matzner, E., Flessa, H., Ekschmitt, K.,
10 Guggenberger, G., Marschner, B. and Kalbitz, K.: Stabilization mechanisms of organic
11 matter in four temperate soils: Development and application of a conceptual model, *J.*
12 *Plant Nutr. Soil Sci.*, 171, 111–124, 2008.
- 13 Wang, G., Huang, L., Xin Zhao, Niu, H. and Dai, Z.: Aliphatic and polycyclic aromatic
14 hydrocarbons of atmospheric aerosols in five locations of Nanjing urban area, China,
15 *Atmos. Res.*, 81, 54–66, 2006.
- 16 Weijers, J. W. H., Schouten, S., Spaargaren, O. C. and Damste, J. S. S.: Occurrence and
17 distribution of tetraether membrane lipids in soils: Implications for the use of the TEX86
18 proxy and the BIT index, *Org. Geochem.*, 37, 1680–1693, 2006.
- 19 Weijers, J. W. H., Bernhardt, B., Peterse, F., Werne, J. P., Dungait, J. A. J., Schouten, S. and
20 Damste, J. S. S.: Absence of seasonal patterns in MBT-CBT indices in mid-latitude soils,
21 *Geochim. Cosmochim. Acta*, 75, 3179–3190, 2011.
- 22 Wiesenberg, G. L. B., Andreeva, D. B., Chimitdorgieva, G. D., Erbajeva, M. A. and Zech,
23 W.: Reconstruction of environmental changes during the late glacial and Holocene
24 reflected in a soil-sedimentary sequence from the lower Selenga River valley, Lake
25 Baikal region, Siberia, assessed by lipid molecular proxies, *Quat. Internat.*, 365, 190–
26 202, 2015.
- 27 Wiesenberg, G. L. B., Lehndorff, E. and Schwark, L.: Thermal degradation of rye and maize
28 straw: Lipid pattern changes as a function of temperature, *Org. Geochem.*, 40, 167–174,
29 2009.
- 30 Wiesenberg, G. L. B., Schmidt, M. W. I. and Schwark, L.: Plant and soil lipid modifications
31 under elevated atmospheric CO₂ conditions: I. Lipid distribution patterns, *Org.*
32 *Geochem.*, 39, 91–102, 2008a.
- 33 Wiesenberg, G. L. B., Schneckenberger, K., Schwark, L. and Kuzyakov, Y.: Use of
34 molecular ratios to identify changes in fatty acid composition of *Miscanthus x giganteus*

1 (Greef et Deu.) plant tissue, rhizosphere and root-free soil during a laboratory
2 experiment, *Org. Geochem.*, 46, 1–11, 2012.

3 Wiesenberg, G. L. B., Schwarzbauer, J., Schmidt, M. W. I. and Schwark, L.: Plant and soil
4 lipid modification under elevated atmospheric CO₂ conditions: II. Stable carbon isotopic
5 values ($\delta^{13}\text{C}$) and turnover, *Org. Geochem.*, 39, 103–117, 2008b.

6 Wiesenberg, G. L. B., Schwarzbauer, J., Schmidt, M. W. I. and Schwark, L.: Source and
7 turnover of organic matter in agricultural soils derived from *n*-alkane/*n*-carboxylic acid
8 compositions and C-isotope signatures, *Org. Geochem.*, 35, 1371–1393, 2004.

9 Wiesenberg, G. L. B. and Gocke, M.: Reconstruction of the late Quaternary
10 paleoenvironments of the Nussloch loess paleosol sequence-Comment to the paper
11 published by Zech et al., *Quaternary Research* 78 (2012), 226–235, *Quatern. Res.*, 79,
12 304–305, 2013.

13 Wiesenberg, G. L. B., Schneckenberger, K., Kuzyakov, Y. and Schwark, L.: Plant lipid
14 composition is not affected by short-term isotopic (^{13}C) pulse-labelling experiments, *J.*
15 *Plant Nutr. Soil Sci.*, 172, 445–453, 2009.

16 Wiesenberg, G. L. B. and Schwark, L.: Carboxylic acid distribution patterns of temperate C3
17 and C4 crops, *Org. Geochem.*, 37, 1973–1982, 2006.

18 Wiesenberg, G. L. B., Schwarzbauer, J., Schmidt, M. W. I. and Schwark, L.: Source and
19 turnover of organic matter in agricultural soils derived from *n*-alkane/*n*-carboxylic acid
20 compositions and C-isotope signatures, *Org. Geochem.*, 35, 1371–1393, 2004.

21 Williams, M., Robertson, E. J., Leech, R. M. and Harwood, J. L.: Lipid metabolism in leaves
22 from young wheat (*Triticum aestivum* cv. Hereward) plants grown at two carbon dioxide
23 levels, *J. Exp. Bot.*, 49, 511–520, 1998.

24 Williams, M., Shewry, P. R. and Harwood, J. L.: The influence of the Greenhouse-Effect on
25 wheat (*Triticum aestivum* L.) Grain Lipids, *J. Exp. Bot.*, 45, 1379–1385, 1994.

26 Williams, M., Shewry, P. R., Lawlor, D. W. and Harwood, J. L.: The effects of elevated
27 temperature and atmospheric carbon dioxide concentration on the quality of grain lipids
28 in wheat (*Triticum aestivum* L) grown at 2 levels of nitrogen application, *Plant Cell*
29 *Environ.*, 18, 999–1009, 1995.

30 Wöhler F. and Liebig, J.: Zusammensetzung einiger ätherischen Oele, *Annalen der*
31 *Pharmacie*, 32, 284–287, 1839.

32 Xie, M., Wang, G., Hu, S., Han, Q., Xu, Y. and Gao, Z.: Aliphatic alkanes and polycyclic
33 aromatic hydrocarbons in atmospheric PM10 aerosols from Baoji, China: Implications
34 for coal burning, *Atmos. Res.*, 93, 840–848, 2009.

- 1 Zech, M., Buggle, B., Leiber, K., Marković, S., Glaser, B., Hambach, U., Huwe, B., Stevens,
2 T., Sümeği, P., Wiesenberg, G. and Zöller, L.: Reconstructing Quaternary vegetation
3 history in the Carpathian Basin, SE Europe, using *n*-alkane biomarkers as molecular
4 fossils - Problems and possible solutions, potential and limitations, *E&G Quat. Sci. J.*,
5 58, 148–155, 2009.
- 6 Zech, M., Leiber, K., Zech, W., Poetsch, T. and Hemp, A.: Late Quaternary soil genesis and
7 vegetation history on the northern slopes of Mt. Kilimanjaro, East Africa, *Quat. Internat.*,
8 243, 327–336, 2011.
- 9 Zech, M., Zech, R., Morras, H., Moretti, L., Glaser, B. and Zech, W.: Late Quaternary
10 environmental changes in Misiones, subtropical NE Argentina, deduced from multi-
11 proxy geochemical analyses in a palaeosol-sediment sequence, *Quat. Internat.*, 196, 121–
12 136, 2009.
- 13 Zech, R., Gao, L., Tarozo, R. and Huang, Y.: Branched glycerol dialkyl glycerol tetraethers
14 in Pleistocene loess-paleosol sequences: Three case studies, *Org. Geochem.*, 53, 38–44,
15 2012.
- 16 Zeng, F., Xiang, S., Zhang, K. and Lu, Y.: Environmental evolution recorded by lipid
17 biomarkers from the Tawan loess-paleosol sequences on the west Chinese Loess Plateau
18 during the late Pleistocene, *Environ. Earth Sci.*, 64, 1951–1963, 2011.
- 19 Zhang, Y., Togamura, Y. and Otsuki, K.: Study on the *n*-alkane patterns in some grasses and
20 factors affecting the *n*-alkane patterns, *J. Agric. Sci.*, 142, 469–475, 2004.
- 21 Zhang, Z. H., Zhao, M. X., Eglinton, G., Lu, H. Y. and Huang, C. Y.: Leaf wax lipids as
22 paleovegetational and paleoenvironmental proxies for the Chinese Loess Plateau over
23 the last 170 kyr, *Quat. Sci. Rev.*, 25, 575–594, 2006.
- 24 Zhou, J., Wang, T., Zhang, Y., Zhong, N., Medeiros, P. M. and Simoneit, B. R. T.:
25 Composition and sources of organic matter in atmospheric PM10 over a two year period
26 in Beijing, China, *Atmos. Res.*, 93, 849–861, 2009.
- 27 Zocatelli, R., Lavrieux, M., Disnar, J., Le Milbeau, C., Jacob, J. and Breheret, J. G.: Free
28 fatty acids in Lake Aydat catchment soils (French Massif Central): sources, distributions
29 and potential use as sediment biomarkers, *J. Soils Sediments*, 12, 734–748, 2012.

1 Tables

2 Table 1: Compounds frequently used as molecular proxies in soils

Compound (the ones considered in this review indicated in bold)	Most commonly used as proxy for:	Examples of recent publications ^a :	Number of articles published until 2017 (<i>publications 2007-2016</i>) ^b
Molecules of plant origin			
<i>n</i>-alkanes, <i>n</i>-alcohols (<i>n</i>-alkanol), <i>n</i>-fatty acids (<i>n</i>-alkanoic acid)	(groups of) plant species	(Zhang et al., 2006; Zeng et al., 2011; Jansen et al., 2013; Gocke et al., 2013)	<i>n</i> -alkane: 1588 (1025) <i>n</i> -alcohol: 1972 (1123); <i>n</i> -alkanol: 18 (11) <i>n</i> -fatty acids: 43 (27); <i>n</i> -alkanoic acid: 67 (41)
<i>n</i>-methyl ketones	degradation/transformation of soil organic matter	(Bai et al., 2006; Jansen and Nierop, 2009; Lei, B. et al., 2010)	<i>n</i> -methyl ketone 104 (50)
plant sterols and pentacyclic triterpenoids	(groups of) plant species	(Volkman, 2005; Jansen et al., 2007; Lavrieux et al., 2011)	plant sterol: 1682 (590) pentacyclic triterpenoid: 25 (10)
lignin monomers	coniferous species vs. broadleaf species vs. grasses and organic matter transformation	(Nierop et al., 2006; Heim and Schmidt, 2007; Thevenot et al., 2010; Simpson and Simpson,	lignin monomer: 115 (74)

		2012; Moingt et al., 2016)	
cutin and suberin monomers	root vs. aboveground biomass input	(Mendez-Millan et al., 2011; Hamer et al., 2012)	cutin monomer: 25 (17) suberin monomer: 32 (18)
Molecules of animal or bacterial origin			
Manure compounds such as coprostanol, 5β-stigmastanol, sitosterol and their epimers	Human impact, animal husbandry	(D'Anjou et al., 2012; Birk et al., 2012; Prost et al., 2017)	coprostanol: 35 (17) stigmastanol: 12 (7) sitosterol: 70 (47)
glycerol dialkyl glycerol tetraethers (GDGT)	mean ambient air temperature, paleo-elevation and soil pH	(Luo et al., 2011; Weijers et al., 2011; Peterse et al., 2012; Ernst et al., 2013; De Jonge et al., 2014)	GDGT: 148 (144)
phospholipid fatty acids (PLFA)	microbial biomass	(Kramer and Gleixner, 2006; Kindler et al., 2009; Ngosong et al., 2012; Malik et al., 2013)	Phospholipid fatty acid: 2157 (1628) PLFA: 1525 (1140)
Compound-specific stable isotope signal of one or more of the above ^c			

$\delta^{13}\text{C}$	C ₃ vs. C ₄ plants and tracing carbon transformations e.g. by free air CO ₂ enrichment (FACE)	(Feng et al., 2010; Mendez-Millan et al., 2012)	¹³ C: 13 (11)
$\delta^{15}\text{N}$	(past) land management	(Bol et al., 2005; Griepentrog et al., 2014)	¹⁵ N: 2 (2)
$\delta^2\text{H}$ (deuterium)	precipitation and paleo-elevation	(Peterse et al., 2009; Bai et al., 2011; Luo et al., 2011; Sachse et al., 2012; Hermann et al., 2017)	² H: 6 (4) deuterium: 9 (7)
$\Delta^{14}\text{C}$ (radiocarbon)	Age and contamination determination	Marschner et al., 2008; Mendez-Millan et al., 2014	¹⁴ C: 3 (1) radiocarbon: 35 (30)

1 ^aPublished from 2005 until 2017.

2 ^bAccording to ISI Web of Science, checked for ‘soil’ and ‘target compound’ in the topic of
3 articles on 27th February 2017 included in all available databases.

4 ^c‘Compound-specific’ and the respective isotope (i.e. ¹³C, ¹⁵N, ²H, and ¹⁴C respectively) were
5 used as separate keywords in addition to ‘soil’.

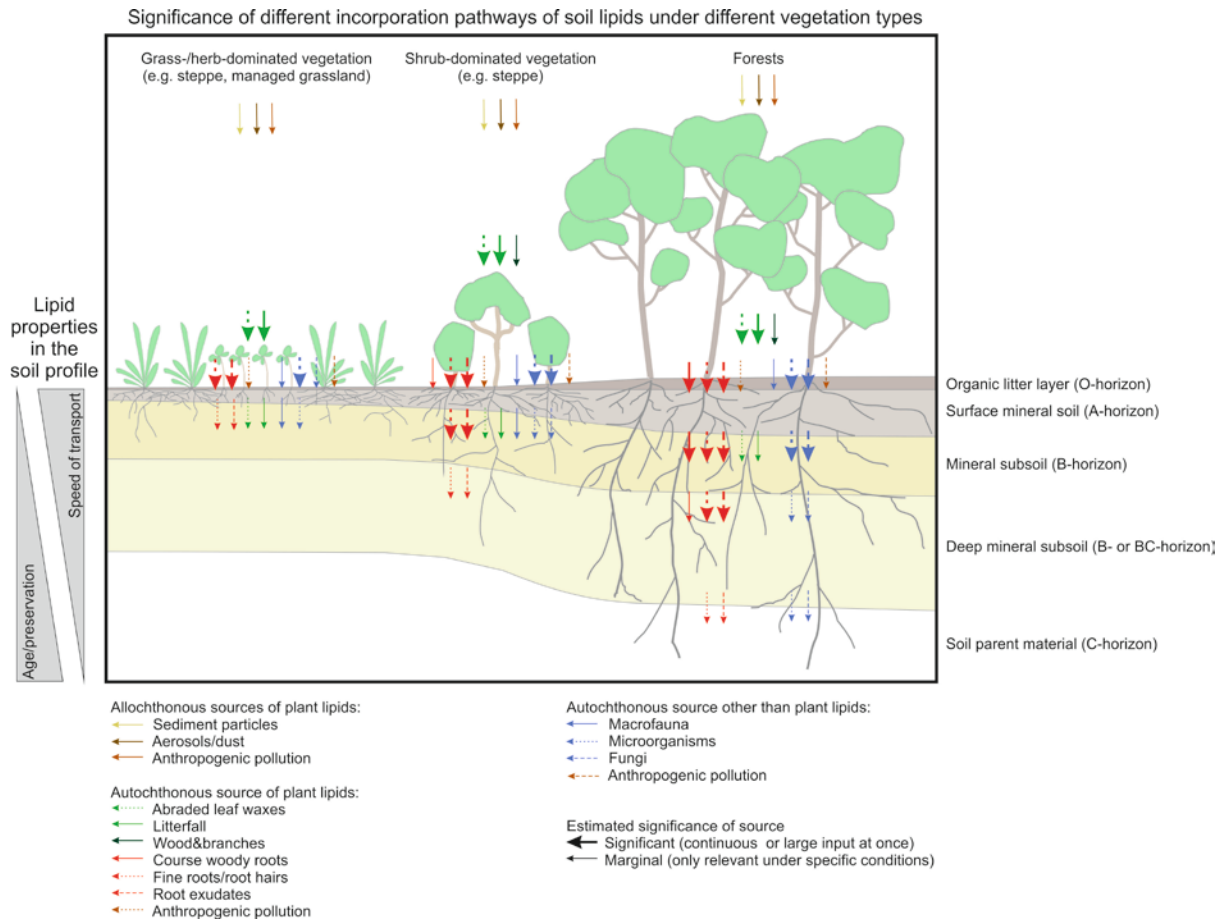
- 1 Table 2: average maximum rooting depth, biomass/depth distribution and root/shoot ratios in
 2 different biomes (Canadell et al., 1996; Jackson et al., 1996)

Biome:	Average maximum rooting depth [m]:	Average percentage of roots in the top 0.3 m:	Average root/shoot ratio:
Boreal forest	2.0±0.3	83	0.32
Cropland	2.1±0.2	70	0.10
Desert	9.5±2.4	53	4.5
Sclerophyllous shrubland and forest	5.2±0.8	67	1.2
Temperate coniferous forest	3.9±0.4	52	0.18
Temperate deciduous forest	2.9±0.2	65	0.23
Temperate grassland	2.6±0.2	83	3.7
Tropical deciduous forest	3.7±0.5	70	0.34
Tropical evergreen forest	7.3±2.8	69	0.19
Tropical grassland/savannah	15.0±5.4	57	0.70
Tundra	0.5±0.1	93	6.6

3

1 **Figures**

2 **Figure 1**



3

4

5 **Figure caption**

6 Conceptual overview of different incorporation pathways of lipids in soils originating from
7 different biological sources and anthropogenic contamination. The different sources are
8 indicated by distinct colors and lines of the arrows. The line thickness is an estimated
9 significance of individual sources, without providing quantitative measure for different
10 sources. Autochthonous sources are further distinguished by their significance in different
11 soil depths or soil horizons, respectively. Further, the transport and age/probability of
12 preservation as general properties of lipids are given at the left side of the figure.

13