



## **Soil microbial biomass and function are altered by 12 years of crop rotation**

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## 1 **Abstract**

2           Agriculture-driven declines in plant biodiversity reduce soil microbial biomass, alter  
3 microbial functions, and threaten the provisioning of soil ecosystem services. We examined  
4 whether increasing temporal plant biodiversity (by rotating crops) can partially reverse these  
5 trends and enhance microbial biomass and function. We quantified seasonal patterns in soil  
6 microbial biomass, respiration rates, extracellular enzyme activity, and catabolic potential three  
7 times over one growing season in a 12-year crop rotation study at the W.K. Kellogg Biological  
8 Station LTER. Rotation treatments varied from one to five crops in a three-year rotation cycle,  
9 but all soils were sampled under corn to isolate historical rotation effects from current crop  
10 effects. Inorganic N, the stoichiometry of microbial biomass and dissolved organic C and N  
11 varied seasonally, likely reflecting fluctuations in soil resources during the growing season.  
12 Soils from biodiverse cropping systems increased microbial biomass C by 28-112% and N by  
13 18-58% compared to monoculture corn. Rotations increased potential C mineralization by as  
14 much as 64%, and potential N mineralization by 62%, and both were related to substantially  
15 higher hydrolase and lower oxidase enzyme activities. The catabolic potential of the microbial  
16 community, assessed with community-level physiological profiling, showed that microbial  
17 communities in monoculture corn preferentially used simple substrates like carboxylic acids,  
18 relative to more diverse cropping systems. By isolating plant biodiversity from differences in  
19 fertilization and tillage, our study illustrates that crop biodiversity has overarching effects on soil  
20 microbial biomass and function that last throughout the growing season. In simplified  
21 agricultural systems, relatively small increases in plant biodiversity have a large impact on  
22 microbial community size and function.

23



24

**Keywords:** crop rotation; agriculture biodiversity; soil carbon; soil nitrogen; nitrogen mining; community-level physiological profile; mineralization; extracellular enzymes; soil microbial biomass



## 25 **Introduction**

26           Research manipulating aboveground biodiversity in grasslands has shown a strong link  
27 between plant species richness and soil functions (Tilman et al. 1997, Zak et al. 2003, Eisenhauer  
28 et al. 2010, Mueller et al. 2013). While this research has contributed to our understanding of  
29 aboveground-belowground biodiversity in natural ecosystems, it fails to capture the biodiversity  
30 dynamics in agroecosystems, where crop rotations can be used to substitute temporal for spatial  
31 biodiversity. Given that species richness at any given time in a rotated cropping system is one,  
32 the aboveground-belowground relationships dependent on diversity in agroecosystems and  
33 spatially diverse ecosystems (e.g. grasslands) may not be the same.

34           Crop rotations have been shown to have large positive effects on soil C, N, and microbial  
35 biomass (McDaniel et al., 2014a), plant pathogen suppression (Krupinsky et al. 2002), and yields  
36 (Ret al. 2008, Riedell et al. 2009). These positive effects of crop production have been  
37 colloquially referred to as the “rotation effect.” However, the mechanistic processes that link  
38 aboveground crop rotational biodiversity and belowground soil processes and contribute to the  
39 “rotation effect” remain elusive. One hypothesis explaining the benefits of crop rotations is that  
40 greater diversity of plant inputs to soil organic matter (SOM) over time enhances belowground  
41 biodiversity and soil ecosystem functioning (Hooper et al., 2000; Waldrop et al., 2006; Grandy  
42 and Robertson, 2007). Despite being low in spatial diversity, crop rotations have been shown to  
43 increase soil microbial and faunal biodiversity (Ryzkowski et al. 1998, Wu et al. 2008, Tiemann  
44 et al. 2015) and increase microbial carbon use efficiency (Kallenbach et al. 2015). One essential  
45 function of soil microbial communities is the catabolism of newly added substrates from crops.  
46 The range and efficiency of microbial catabolism has great implications for ecosystem services



47 such as sequestering C and soil fertility (Carpenter-Boggs 2000; Kallenbach et al. 2015), but also  
48 for ecosystem “dis-services” such as emission of soil-to-atmosphere greenhouse gases  
49 (McDaniel et al. 2014b). Furthermore, the partitioning of resources used in catabolism of  
50 residue and formation of SOM will likely affect long-term soil fertility (Lange et al. 2015;  
51 Kallenbach et al. 2015).

52 Soil microbial catabolic function can be measured using community-level physiological  
53 profiles (CLPP), sometimes referred to as catabolic response profiles. The basic method for  
54 measuring soil CLPP involves adding a suite of C substrates to soils and measuring the catabolic  
55 response as CO<sub>2</sub> production or O<sub>2</sub> consumption with redox indicators (i.e. Biolog; Guckert et al.  
56 1996). These C substrates are typically ecologically-relevant compounds found in soils, and are  
57 intended to represent root exudates, microbial or plant cell structures, or other more-processed  
58 soil organic molecules. By measuring the respiration response, we can establish a catabolic  
59 “fingerprint” to distinguish soil microbial communities from one another by how they catabolize  
60 these new C substrates. We can also derive catabolic evenness of that community – a measure of  
61 biodiversity. Modern agriculture’s use of monocultures could have unknown consequences for  
62 soil microbial catabolism and related processes such as SOM mineralization, but to date the  
63 effect of rotation practices and crop diversity on soil microbial functioning remains poorly  
64 understood.

65 Considering the lack of a clear understanding of how soil microbial functions are  
66 influenced by crop rotations, we sought to examine their effects on soil microbial biomass and  
67 function. We measured soil microbial catabolic potential, C and N mineralization, extracellular  
68 enzyme activities, and microbial biomass three times over one growing season in a long-term



69 crop rotation experiment at the W.K. Kellogg Biological Station (est. 2000). All soils were  
70 measured under the same crop, allowing us to separate historical rotation from current crop  
71 effects. We hypothesized that soils under more diverse crop rotations would show greater  
72 catabolic diversity and have higher measures of soil function (enzyme activities, soil microbial  
73 biomass, potentially mineralizable C and N). In addition, we hypothesized that crop rotation  
74 effects would vary seasonally, being greatest in the spring and lessen over the growing season  
75 with the emerging influence of the current crop. The rationale for this second hypothesis is that  
76 early in the season all soils are coming out of different crops from the previous year, but over the  
77 growing season under corn the soils will become more functionally similar as the immediate crop  
78 has greater influence. Alternatively, significant Rotation by Season interactions on soil  
79 microbial functioning that do not converge over the growing season point to historical effects of  
80 rotations on differences in soil microbial communities and SOM.

## 81 **Materials and Methods**

82 This study was conducted in the Cropping Biodiversity Gradient Experiment (CBGE) at  
83 the W.K. Kellogg Biological Station Long-term Ecological Research site (42° 24', 85° 24' W).  
84 The CBGE was established in 2000 and consists of crop rotations ranging from monocultures to  
85 a 5-species rotation ([http://lter.kbs.msu.edu/research/long-term-experiments/biodiversity-](http://lter.kbs.msu.edu/research/long-term-experiments/biodiversity-gradient/)  
86 [gradient/](http://lter.kbs.msu.edu/research/long-term-experiments/biodiversity-gradient/)). The crop rotations were repeated but with different rotation phases within all four  
87 blocks. For example, the corn-soy-wheat rotation is replicated three times within each block, but  
88 these replicates are planted to a different crop each year. The plot dimensions were 9.1 x 27.4 m  
89 and received the same chisel plow tillage to a depth of approximately 15 cm, and received no  
90 inputs (e.g. pesticides or fertilizers) that would have confounded the treatment effects of  
91 rotational diversity (Smith et al. 2008). Mean annual temperature and precipitation at the site are



92 9.7°C and 890 mm, respectively. The two main soil series located at the site are Kalamazoo, a  
93 fine-loamy, mixed, mesic Typic Hapludalf, and Oshtemo, a coarse-loamy, mixed, mesic Typic  
94 Hapludalf (KBS, 2012). Soil pH in the top 10 cm ranges from 4.9 to 6.1 (1:1 v of 0.01 M  
95 CaCl<sub>2</sub>).

96 Soils were collected from the following cropping systems: monoculture corn (*Zea mize*  
97 *L.*, mC), corn-soy (*Glycine max*, CS), corn-soy-wheat (*Triticum aestivum*, CSW), corn-soy-  
98 wheat with red clover cover crop (*Trifolium pretense*, CSW1), and corn-soy-wheat with red  
99 clover + rye cover crops (*Secale cereale*, CSW2). Soil sampling took place on April 27<sup>th</sup>, 2012;  
100 July 19<sup>th</sup>, 2012; and November 1<sup>st</sup>, 2012 – hereafter referred to as spring, summer, and autumn.  
101 Corn was planted in all plots on June 11<sup>th</sup> 2012. Three 5 cm soil cores (0-10 cm deep) were  
102 collected within each plot, homogenized in the field, and then put on ice and shipped to the  
103 University of New Hampshire. In the lab, field-moist soils were immediately sieved using a 2  
104 mm sieve. A sub-sample was taken from sieved soil and dried at 105 °C to determine  
105 gravimetric water content. Water-holding capacity was determined as the water content after  
106 soils were saturated and drained for 6 h.

#### 107 *Soil carbon and nitrogen parameters*

108 Five g of field-moist soil were extracted for inorganic N with 40 ml of 0.5 M K<sub>2</sub>SO<sub>4</sub>. The  
109 soil slurries were shaken for 1 h before the extracts were filtered on Whatman GF/C (5) filters  
110 and filtrate frozen and stored until analysis. Soil nitrate (NO<sub>3</sub><sup>-</sup>) and ammonium (NH<sub>4</sub><sup>+</sup>) were  
111 measured using the methods detailed in McDaniel et al. (2014c). We also used the same extracts  
112 to measure dissolved organic C and N (DOC and DON). The extracts were run on a TOC-TN  
113 analyzer (TOC-V-CPN; Shimadzu Scientific Instruments Inc., Columbia, MD, USA). Total C



114 and N were analyzed by sieving soils through 2 mm sieve, grinding and analyzing on an ECS  
115 4010 CHNSO Elemental Analyzer (Costech Analytical Technologies, Inc., Valencia, CA).

116 Potential mineralization rates of C (PMC) and net N (or PMN) estimate the quantity of  
117 potentially-mineralizable SOM at an optimal temperature and soil moisture, and reflect both the  
118 activity of the microbial community and availability of SOM (Paul et al. 1999; Robertson et al.  
119 2000). These mineralization assays provide a good indicator of the potential for a soil to provide  
120 plants with N (Stanford and Smith 1972, Robertson et al. 1999). Both PMC and PMN were  
121 measured on air-dried soils that were placed into Wheaton serum vials and brought to 50%  
122 water-holding capacity and incubated for 6 months. During this 6-month period CO<sub>2</sub> efflux was  
123 measured on a LI-820 infrared gas analyzer (LI-COR, Lincoln, NE). Efflux was measured using  
124 the change in headspace CO<sub>2</sub> concentration measured between two time points. Each soil efflux  
125 measurement began by aerating jars, capping, and injecting a time-zero sample and then a second  
126 sample between 5 hours up to 2 days later. Efflux was calculated as the difference in CO<sub>2</sub>  
127 concentration between the two time points divided by time. Measurements of PMC occurred  
128 more frequently at the beginning of the experiment (daily) and became less frequent toward the  
129 end (once every other week), for a total of 19 sampling events over 120 days. The PMN was  
130 assessed by extracting the inorganic N produced at the end of the incubation with the methods  
131 described above.

### 132 *Soil microbial parameters*

133 Soil microbial biomass C and N were determined using the modified chloroform  
134 fumigation and extraction method (Vance et al. 1987), but modified for extraction in individual  
135 test tubes (McDaniel et al. 2014c). Briefly, two sets of fresh, sieved soil (5 g) were placed in 50  
136 ml test tubes, and 1 ml of chloroform was added to one set of tubes and capped. The tubes sat



137 overnight (24 h) and were then uncapped and exposed to open air in a fume hood to allow  
138 chloroform to evaporate. Soils were then extracted in the tubes with 25 ml of 0.5 M K<sub>2</sub>SO<sub>4</sub>. The  
139 chloroform fumigated and non-fumigated extracts were run on a TOC-TN analyzer (TOC-V-  
140 CPN; Shimadzu Scientific Instruments Inc., Columbia, MD, USA). We used 0.45 (Joergensen  
141 1996) and 0.54 (Brookes et al. 1985) for the C and N extraction efficiencies.

142         Soils were analyzed for 7 extracellular enzyme activities (EEAs):  $\beta$ -1,4-glucosidase  
143 (BG),  $\beta$ -D-1,4-cellobiohydrolase (CBH),  $\beta$ -1,4-N-acetyl glucosaminidase (NAG), acid  
144 phosphatase (PHOS), Tyrosine aminopeptidase (TAP), Leucine aminopeptidase (LAP),  
145 polyphenol oxidase (PO), and peroxidase (PER). Given the large number of samples (60) and  
146 variety of measurements made at each of 3 sampling dates, soil EEAs were conducted on frozen  
147 samples within 4 weeks of sampling. Extracellular enzyme activity assays were carried out  
148 following previously published protocols (Saiya-Cork et al. 2002, German et al. 2011), but with  
149 some modifications. Briefly, 1 g of soil was homogenized with a blender in 80 ml of sodium  
150 acetate buffer at pH 5.6 (the average pH at the site). Soil slurries were pipetted into 96-well  
151 plates and then analyzed on a Synergy 2 plate reader (BioTek Instruments, Inc., Winooski, VT).  
152 For oxidoreductase enzymes, the supernatant from the slurry plates were pipetted into a clean  
153 plate to avoid interference with soil particles. Hydrolase assays were read at 360/40 and 460/40  
154 fluorescence and oxidoreductases at 450 nm absorbance. For more details on the extracellular  
155 enzyme methods see McDaniel (2014c).

156         Community-level physiological profiles (CLPP) were conducted using the MicroResp<sup>TM</sup>  
157 system (Chapman et al. 2007, Zhou et al. 2012, McDaniel et al. 2014b). The MicroResp<sup>TM</sup>  
158 system allows for high-throughput measurement of soil catabolic responses to multiple C  
159 substrates. Each soil was loaded into 96 deep-well plates using the MicroResp<sup>TM</sup> soil dispenser,



160 and then brought to 50% water-holding capacity. Thirty-one substrates were used at  
161 concentrations ranging from 7.5 to 30 mg C per g of soil H<sub>2</sub>O, as recommended by the  
162 MicroResp<sup>TM</sup> manual (Table S1). Soil and substrates were combined in analytical triplicates and  
163 a CO<sub>2</sub> detection plate (agar containing creosol red) was immediately placed onto the deep-well  
164 plate with an air tight seal provided by the MicroResp<sup>TM</sup> kit. The soil and substrates were  
165 incubated in the dark for 6 h at 25 °C. The detector plate absorbencies were read at times 0 and  
166 6 h at 540 nm on a Synergy 2 plate reader (BioTek Instruments, Inc., Winooski, VT).  
167 Absorbance data were normalized and converted to a CO<sub>2</sub> efflux rate (μg CO<sub>2</sub>-C g soil<sup>-1</sup> h<sup>-1</sup>),  
168 according to the MicroResp<sup>TM</sup> procedure (Chapman et al. 2007).

#### 169 *Data analyses*

170 Cumulative potentially mineralizable C and N were calculated in SigmaPlot v12.5 (Systat  
171 Software, Inc., San Jose, CA) using the integration macro, *area below curves*. Data not  
172 conforming to ANOVA assumptions of homogeneity of variances and normality were  
173 transformed before analyses (Zuur et al. 2010). Catabolic evenness (CE), a measure of substrate  
174 diversity, was calculated using the Simpson-Yule index,  $CE = 1/\sum p_i^2$ , where  $p_i$  is the proportion  
175 of a substrate respiration response to the total response induced from all substrates (Degens et al.  
176 2000, Magurran 2004). Metabolic quotient was calculated simply as the basal respiration over 6  
177 h (determined in the MicroResp<sup>TM</sup> method) divided by the MBC.

178 Response variables were analyzed using a 2-way analysis of variance (ANOVA), with  
179 Season and Rotation as main effects. The ANOVAs were conducted in SAS 9.3 (SAS Institute,  
180 Cary, NC) using the *proc mixed* function and post-hoc *t* tests were used to determine significant  
181 differences among means using *ls means*. Block was assigned as a random effect variable within



182 the model. Correlations between variables were made using *proc corr*, and Pearson's correlation  
183 coefficients are reported. Model effects were deemed significant if  $\alpha < 0.05$ .

184 All multi-variate data analyses were performed with R software (The R Foundation for  
185 Statistical Computing, Vienna, Austria). CLPP data were checked to ensure they conformed to  
186 principal components analysis assumptions. The *prcomp* function in the *vegan* package  
187 (Oksanen et al. 2016) was used for PCA of CLPP data. In order to correlate environmental  
188 variables with the multi-variate CLPP data we used the *envrfit* function.

## 189 **Results**

190 It was a relatively dry year at the KBS-LTER in 2012, which had an annual precipitation  
191 of 742 mm, compared to the historical mean of 870 mm (Hamilton et al. 2015). There was also  
192 an anomalous warm spell in mid- to late-March (Fig. S1). After harvest, the corn yield ( $\text{kg ha}^{-1} \pm$   
193 SE) in each treatment was as follows: mC =  $2846 \pm 152$ , CS =  $4208 \pm 575$ , CSW =  $4107 \pm 220$ ,  
194 CSW1 =  $4015 \pm 187$ , CSW2 =  $5219 \pm 1180$  (KBS-LTER 2015).

### 195 *Soil C and N biogeochemistry*

196 There were no significant Rotation or Season effects on total soil C and N, although both  
197 soil C and N tended to increase with the number of crops in rotation (Table1). Soil  $\text{NO}_3^-$ -N was  
198 the only variable that showed a significant Season X Rotation interaction ( $P < 0.001$ ). Seasonal  
199 soil  $\text{NO}_3^-$ -N concentrations were highest in summer (10.33) followed by spring (2.98), and  
200 autumn (1.28  $\text{mg kg}^{-1}$ ). Soil  $\text{NH}_4^+$ -N was generally low, but summer had more than twice the  
201 concentrations of spring and autumn. Dissolved organic C (DOC) and N (DON) was very  
202 dynamic over the year. The DOC was highest in the autumn, while DON was over six times  
203 greater in the summer than the other seasons ( $P_s < 0.001$ ). The mean DOC:DON in autumn was



204 22.5, twice that of spring and 13 times that of the summer. There were significant crop rotation  
205 effects on  $\text{NO}_3^-$ -N, DOC, and DON. During the summer the two cover crop treatments had  $\text{NO}_3^-$   
206 concentrations 67% greater than CSW and CS treatments, and 158% greater than mC. The  
207 CSW1 treatment had 112% greater DOC concentrations than mC ( $P < 0.001$ ), and two cover  
208 crop treatments had 107% greater DON than non-cover crop treatments and 211% more than the  
209 mC treatment. The potentially mineralizable pools of C and N showed significant main effects of  
210 both Season and Rotation ( $P < 0.03$ ), but no interactions. Generally, both PMC and PMN  
211 increased with increasing number of crops in rotation (Fig. 1), and PMC was highest during the  
212 autumn, while PMN was highest during the summer.

### 213 *Soil microbial dynamics*

214 The mean soil microbial biomass C (MBC) was  $332 \mu\text{g C g soil}^{-1}$  across all seasons and  
215 crop rotations, but both Season ( $P < 0.001$ ) and Rotation ( $P = 0.008$ ) had significant effects on  
216 MBC (Fig. 2). Soils collected in autumn had more than twice the MBC than those collected in  
217 spring and summer. Microbial biomass C was increased by increasing crop diversity across all  
218 seasons (Fig. 2). Increasing the number of crops in rotation increased MBC on average by 28,  
219 113, 112, and 72% in the CS, CSW, CSW1, CSW2 rotations, respectively, compared to mC  
220 (across all seasons). Microbial biomass N (MBN) also showed both Season ( $P < 0.001$ ) and  
221 Rotation ( $P = 0.005$ ) effects, but no interaction. These effects were strongest in the spring and  
222 summer (Fig. 2), but also showed an increase with increasing number of crops. Increasing the  
223 number of crops in rotation increased MBN on average by 18, 58, 54, and 50% in the CS, CSW,  
224 CSW1, and CSW2 compared to mC (across all seasons). Microbial biomass C:N showed a  
225 significant interaction ( $P = 0.013$ ), with more diverse cropping systems having greater  
226 MBC:MBN in summer and autumn, but not spring. The metabolic quotient ( $q\text{CO}_2$ ), is often used



227 as a proxy for microbial respiration efficiency (Anderson & Domsch 1989, 2010; Wardle &  
228 Ghani 1995). Season ( $P < 0.001$ ) and Rotation ( $P = 0.006$ ) both influenced  $qCO_2$ , with increased  
229 crop diversity decreasing the  $qCO_2$  by 16, 40, and 28 % in CSW, CSW1, and CSW2 compared to  
230 mC. However, the CS rotation increased  $qCO_2$  by +15 %  $qCO_2$  compared to mC (Fig. 2).

231 Soil extracellular enzymes were very dynamic over the three seasons, as evidenced by  
232 radar plots in which the area and shape for each treatment change quite drastically over the  
233 growing season (Fig. 3). A MANOVA with all eight EEAs showed significant Season ( $P <$   
234  $0.001$ ) and Rotation ( $P < 0.001$ ) main effects, but no interaction. Most individual enzymes  
235 showed only significant Rotation effects except for PO, which also showed a significant Season  
236 effect with autumn greater than the other seasons (Table 2). The soil enzyme responsible for  
237 cleaving a glucosamine from chitin (NAG) and the lignin-reducing enzyme that uses peroxide  
238 (PER) were the only enzymes that showed a significant Season X Rotation interaction ( $P <$   
239  $0.001$ ). Spring had the greatest activities of LAP, 175% greater than the average of the other  
240 seasons (Fig. 3, Table 2). In summer, we see a shift to the highest PHOS activity – 25% greater  
241 than spring and 99% greater than autumn. Season had no effect on BG or CBH but showed  
242 significant main effects of rotation, with the CSW1 treatment having an average of 42 and 50 %  
243 higher BG and CBH activity than CS and mC soils, respectively. The majority of the hydrolase  
244 enzymes were higher in the cover crop treatments compared to that of the non-cover crop  
245 treatments, especially mC (Table 2, Fig. 3). The two oxidoreductase enzymes (PO and PER)  
246 decreased with crop diversity. There were no significant main effects on the enzyme ratio used  
247 to assess C-versus-N demand (BG to NAG+LAP).

248 The community-level physiological profile (CLPP), a catabolic profile of the soil  
249 microbial communities, showed both significant Season ( $P < 0.001$ ) and Rotation ( $P = 0.003$ )



250 main effects (Figs. 4, S2; Table 3). A principal components analysis of the CLPP data showed  
251 that the summer soils corresponded with highest carboxylic acid utilization (Fig. 4), as Season  
252 was the strongest discriminating factor along principal component 1 (PC1, Table 3). However,  
253 when rotating and examining PC2 and PC3, there was a strong treatment gradient from the  
254 bottom right to upper-left quadrants of the graph (Fig. 4, right panel). The lower-diversity  
255 treatments corresponded with greater use of carboxylic acid substrates. Across seasons, summer  
256 exhibited the lowest catabolic evenness (12.9), but there was no crop rotation effect on catabolic  
257 evenness using all substrates (Table 4).

258         Due to the overwhelming influence of carboxylic acids in the PCA variation, and their  
259 possible role in abiotic reactions leading to CO<sub>2</sub> emissions (Maire et al. 2012, Pietravalle and  
260 Aspray 2013), we split the 31 substrates into two sets to analyze separately: 1) Non-carboxylic  
261 acid substrates – a total of 21 substrates, and 2) carboxylic acids by themselves – 10 substrates.  
262 Season, again, was a dominant significant effect on the MANOVAs in both groups of substrates  
263 ( $P$  values  $< 0.001$ , Fig. S3). The non-carboxylic acid CLPP showed a significant treatment effect  
264 with PC1 and PC2, and clear separation between low and high diversity cropping systems ( $P =$   
265 0.012, Fig. S3). The monoculture corn, and lower diversity treatments, associated with more  
266 complex substrates. In the carboxylic acid CLPP there was also a significant treatment effect,  
267 but with PC2 and PC3, and clear separation between low and high diversity cropping systems  
268 along PC3 ( $P = 0.035$ , Fig. S3). Interestingly, the lower diversity (especially monoculture corn),  
269 were more associated with simple carboxylic acids (Cit, Mlo, and Mli) on the positive half of  
270 PC3. When carboxylic acids were split from the substrates, crop rotation had a significant effect  
271 on catabolic evenness – decreasing the catabolic evenness both within non-carboxylic acids and  
272 carboxylic acids by as much as 4 and 13% respectively (Table 4).



273 We used the soil microbial responses of EEA and the CLPP because we assumed they  
274 would be complimentary. Indeed, this was the case. Measuring NAG enzyme and adding the  
275 Nag amine to the soils showed a somewhat tight relationship, but this relationship was not  
276 constant over the seasons. More specifically, the NAG enzyme was quite higher in the autumn  
277 compared to summer and spring, and showed a steeper linear relationship with the CO<sub>2</sub> response  
278 after adding the Nag amine to soils (Fig. S4). Additionally, when the CLPP substrates were  
279 grouped by guild and correlated with EEA there were strong relationships (Fig. S5). For  
280 example, total amino acid catabolic response positively correlated well with LAP+TAP enzymes  
281 ( $r^2 = 0.35$ ,  $P < 0.001$ ) meaning that high activity of these enzymes in soils corresponded with  
282 high relative use of these substrates when added to soils, compared to other substrates added to  
283 the soil. This suggests that the LAP and TAP enzymes strongly reflect demand for N-bearing  
284 amino acids in soils. However, the catabolic response of the ‘Complex’ guild was negatively  
285 correlated with PO ( $r^2 = 0.29$ ,  $P < 0.001$ ).

#### 286 *Relationships between soil biogeochemical factors, microbial functioning and yield*

287 Over the three seasons many soil biogeochemical factors correlated with microbial  
288 catabolic potential, both with individual C substrate guilds and catabolic evenness (Table 5).  
289 Abiotic factors such as pH and sand content correlated with the specific use of particular  
290 substrates. Soil pH positively correlated with N-containing and complex substrates, but strongly  
291 negative with carboxylic acids. Sand content negatively correlated with amino acids and  
292 carbohydrates, but positively with carboxylic acids. The microbial response to amino acids and  
293 amines correlated best with NO<sub>3</sub><sup>-</sup>-N and many of the specific enzyme activities, showing  
294 negative relationships which indicated a linkage between demand for N and usage of N-bearing  
295 substrates (i.e. when supply is high, demand and usage of N substrates is low). Soil NO<sub>3</sub><sup>-</sup>-N was



296 also significantly negatively correlated with catabolic evenness. We used the soil microbial  
297 responses of EEA and the CLPP because we assumed they would be complimentary. For  
298 example, adding N-acetyl glucosamine in the CLPP should be related to  $\beta$ -1,4-N-acetyl  
299 glucosaminidase (NAG) enzyme activity. Indeed, this was the case. Measuring NAG enzyme  
300 and adding the Nag amine to the soils showed a somewhat tight relationship (Fig. S4).  
301 Additionally, when the CLPP substrates were grouped by guild and correlated with EEA there  
302 were strong relationships (Fig. S5). For example, total amino acid catabolic response positively  
303 correlated well with LAP+TAP enzymes ( $r^2 = 0.35$ ,  $P < 0.001$ ) meaning that high activity of  
304 these enzymes in soils corresponded with high relative use of these substrates when added to  
305 soils, compared to other substrates added to the soil. This suggests that the LAP and TAP  
306 enzymes strongly reflect demand for N-bearing amino acids in soils. However, the catabolic  
307 response of the ‘Complex’ guild was negatively correlated with PO ( $r^2 = 0.29$ ,  $P < 0.001$ ). Soil  
308 PMN was better correlated with crop yields ( $r^2 = 0.61$ ) than  $\text{NO}_3^-$  in early spring (Fig. S6).

### 309 Discussion

310 Increasing biodiversity in this long-term crop rotation experiment has altered the soil  
311 microbial dynamics across an entire growing season. This is despite the fact that the soils in our  
312 study were all under the same crop (corn) for the season, indicating that observed differences  
313 among soils reflect long-term rotation effects rather than the current crop. Microbial biomass C,  
314 N, potential mineralization, and catabolic potential were all altered by crop rotations, although  
315 the rotation effect for some of these indicators of microbial functioning also depends upon the  
316 season. Soil microbial biomass and activity are now widely recognized as pillars of soil health  
317 (Doran and Zeiss 2000). Our results clearly indicate that practices like diversifying  
318 agroecosystems (through crop rotations) enhances this aspect of soil health, and this is also likely



319 linked to changes in SOM dynamics (Tiemann et al. 2015) as well as the observed differences in  
320 yield among crop rotations (Fig. S6).

### 321 *Seasonal dynamics and N limitation*

322           Season strongly influenced the measured pools of labile C and N (Table 1), as well as the  
323 microbial biomass size and functioning within this agroecosystem (Figs. 1-4). We showed that  
324 the greatest microbial biomass and activity occurred in autumn, but that potentially mineralizable  
325 N peaked in summer. In perennial and annual cropping systems in Iowa, potentially  
326 mineralizable N declined from spring to late summer (Hargreaves and Hofmockel 2013); in  
327 addition, extracellular enzyme activities peaked in July but there was little effect of the cropping  
328 system. Season was shown to affect microbial biomass and potentially mineralizable C and N  
329 pools in a wheat-sorghum-soybean rotation in south-central Texas (Franzluebbers et al. 1994,  
330 1995, Franzluebbers 2002), but timing for peak values differed depending on the study and  
331 cropping systems, likely reflecting different climates and soil types. The frequently observed  
332 late-summer spike in microbial biomass and activity may be related to higher temperatures  
333 during this time period; however, even within agroecosystems, the timing for maximal microbial  
334 biomass varies substantially, although few microbial maxima are reported in winter (Wardle,  
335 1992). Our findings highlight the dynamic nature of soil microbial biomass and activity,  
336 especially with regards to the supply and demand of N (e.g. microbial C:N, substrate utilization,  
337 and extracellular enzyme activities), which is likely a limiting nutrient in these agroecosystems  
338 that are receiving no exogenous N inputs.

339           The summer warrants discussion because the sample was collected after a prolonged  
340 period of hot and dry days, but right after a large rainfall event. This rainfall event ( $> 18 \text{ mm d}^{-1}$ ,



341 Fig. S1) increased the volumetric water content in the 0-10 cm of a nearby soil by over 54%  
342 from the lowest value of the year (0.1, data shared from Hamilton et al. 2015), and we know  
343 from previous research that drying-wetting cycles are important soil biogeochemical drivers  
344 (Borken and Matzner, 2009) and can alter microbial structure and functioning (Fierer et al. 2003,  
345 Schimel et al. 2007, Tiemann and Billings 2011, McDaniel et al. 2014b). Indeed, the summer  
346 showed several signs of the soil microbial community being impacted by a rapid dry-wet event:  
347 lower overall microbial biomass C, extremely high  $\text{NO}_3^-$ -N concentrations (Table 2), high  
348 potential N mineralization (Fig. 1), high extracellular enzyme activities per unit of microbial  
349 biomass (Fig. S2, presumably a result of lysed intracellular enzymes, Burns et al. 2013), and the  
350 particularly strong response of the summer soils to carboxylic acids (a highly-labile class of  
351 compounds used by fast-growing, opportunistic microbes, that would be found after a  
352 disturbance such as a dry-wet event, Figs. 4 and S3). Dry-wet cycles may drive microbial C and  
353 N to be reallocated to stress-response compounds instead of growth or reproduction, making C  
354 and N more vulnerable to loss from soils (Schimel et al. 2007). We captured one of these dry-  
355 wet events during one of the driest summers in the Kellogg Biological Station LTER's history  
356 and we show high soil inorganic N concentrations and altered microbial dynamics relative to the  
357 other dates. Climate change may increase the frequency and magnitude of these rapid dry-wet  
358 cycles (Groffman et al. 2001, McDaniel et al. 2014d), and thus may have long-term impacts on  
359 soil microbial functioning and biogeochemistry.

360 In the autumn we found several lines of evidence that indicate soil microbes are N, rather  
361 than C, limited. These lines of evidence include: lowest soil inorganic N concentrations, low  
362 potentially mineralizable N, high microbial biomass C:N and DOC:DON ratios, and high TAP  
363 and NAG enzymes relative to other enzymes (although interestingly not LAP), and finally strong



364 respiration response to the addition of amines and amino acids (Fig. 4). The unusually high  
365 microbial biomass C:N in autumn was very surprising (mean of 18.2 versus 7.5 and 9.9 in spring  
366 and summer), but microbial biomass C:N has been known to reach as high as 30 in laboratory  
367 conditions (Schimel et al. 1989). Additionally, the few days before and after the collection of the  
368 autumn sample were unusually cold (Fig. S1), and cold temperatures and freezing can cause  
369 accumulation of carbohydrates in fungi (Tibbits et al. 2002), which could widen microbial C:N  
370 ratio. Regardless of environmental conditions widening microbial biomass C:N, it is likely that  
371 N limitation is a major factor considering this experiment receives no fertilizer N for 12 years.

#### 372 *Crop biodiversity and soil microbial functioning*

373 Both soil microbial biomass and functioning were strongly affected by increased crop  
374 diversity through rotation. This rotation effect on soil microbial biomass and functioning were  
375 largely independent of the season, as indicated by the limited number of observed Season X Crop  
376 Rotation interactions. The exception to this was microbial biomass C:N and two extracellular  
377 enzyme activities (NAG and PER), which are likely indicative of the enhanced ability of soil  
378 microbes under diverse rotations to process, provision, and retain N. These findings are  
379 consistent with plant biodiversity studies which find increased aboveground diversity enhances  
380 soil microbial biomass and functioning in natural (Stephan et al. 2000, Zak et al. 2003, Lange  
381 2015) and agricultural ecosystems (Lupwayi et al. 1998; Xuan et al. 2012; Tiemann et al. 2015).

382 In our study, the strong response of soil microbial functioning to crop diversity is  
383 particularly interesting given that all treatments were under the same crop (corn) over the entire  
384 growing season. Some studies have shown that the current plant species identity often trumps  
385 biodiversity legacy in controlling belowground microbial structure and functioning (Stephan et



386 al. 2000, Wardle et al. 2003, Bartelt-Ryser et al. 2005). Conversely, several studies have pointed  
387 to weak or no influence of current plant species on soil microbial structure and functioning  
388 (Costa et al. 2006, Kielak et al. 2008). The question of whether plant species identity versus  
389 spatial and temporal diversity has a stronger control on soil biota remains a critical question in  
390 terrestrial ecology. Here we show strong evidence for a biodiversity “carry-over” effect (Bartelt-  
391 Ryser et al. 2005), where the legacy of crop rotation affects soil microbial biomass and  
392 functioning of the current year, even though the soils are all planted under the same crop species.

393 We hypothesized that increasing biodiversity in agroecosystems through rotation would  
394 result in soil microbial communities that are more diverse, and thus would more evenly use  
395 added C substrates (i.e. increase catabolic evenness, or decrease the variation in use among  
396 substrates). This hypothesis stems from the “plant diversity hypothesis”, which posits that soil  
397 community and functional biodiversity is linked to plant biodiversity mostly through the  
398 diversity of plant inputs to SOM (Lodge, 1997; Hooper et al., 2000; Waldrop et al., 2006;  
399 Korboulewsky et al., 2016). However, in our study, we found no evidence that crop rotational  
400 diversity increased overall soil catabolic evenness. There is some evidence that crop rotations  
401 can alter soil bacterial catabolic diversity, or the ability to use different C substrates (Lupwayi et  
402 al. 1998, Larkin 2003, Govaerts et al. 2007), however all of these studies used Biolog which has  
403 several limitations (Preston-Mafham et al. 2002). The MicroResp™ system’s main benefit is  
404 that it adds C substrates directly to the soil instead of transferring an inoculum from a soil slurry.  
405 The discrepancy between our study and these studies other studies may be due to methodological  
406 differences between Biolog and MicroResp™. Our lack of evidence for an aboveground-  
407 belowground link to catabolic potential aligns with findings from other studies that have found



408 functional diversity measures of soil microbes are not related to plant diversity (Bartelt-Ryser et  
409 al. 2005, Jiang et al. 2012), nor plant species in general (McIntosh et al. 2013).

410           However, it is important to note that in our study when a subset of the C substrates were  
411 analyzed (all non-carboxylic), we found that increased crop diversity actually decreased  
412 catabolic evenness. This is unusual considering soils from this same study, but collected a year  
413 prior, showed increases of soil biodiversity (Shannon-Weiner index or  $H'$ ) with increased crop  
414 diversity when measuring phospholipid fatty acids (Tiemann et al. 2015); and diversity has been  
415 found to be strongly, positively related to evenness in plants and animals (Stirling & Wilsey  
416 2001). Our finding of no change in (or lower) catabolic evenness with increasing crop  
417 biodiversity is also contradictory to the findings of Degens et al. (2000), whom showed that  
418 management practices that decrease soil C are associated with low catabolic evenness since we  
419 show a general trend of increasing soil C.

420           Perhaps the incongruity between the positive effect of crop rotation on  $H'$ , but slightly  
421 negative effect on catabolic evenness, could lie in the difference between who is targeted by the  
422 catabolic response profile. Most bacteria are thought to be generalists with regard to the use of C  
423 substrates (Mou et al. 2008; Goldfarb et al. 2011), whereas fungi may show more specialization  
424 (Hanson et al. 2008; Treseder et al. 2015). Not to mention the catabolic response method used  
425 here is probably more favorable to detecting response of bacteria rather than fungi, because fungi  
426 are more sensitive to the disturbance of preparing the soils (Frey et al., 1999) and bacteria have  
427 faster growth and reproduction (i.e. respiration was measured for only 6 h). Thus a high  
428 catabolic evenness may not be a good indicator of soil biodiversity in soils with high relative  
429 bacterial biomass and activity, which is typical of many agroecosystem soils (Strickland &  
430 Rousk 2010). Instead, a low catabolic evenness may actually be better suited to detect C-use



431 specialization in bacterial-dominated microbial communities. In support of this idea, we have  
432 evidence from these same soil samples that crop diversity significantly decreased  $H'$  for bacterial  
433 16S rRNA by as much as 5 % compared to monoculture corn (Peralta et al. *in preparation*).  
434 These inconsistencies, especially between methods of measuring soil microbial diversity, are  
435 highlighted in a recent quantitative review (Venter et al. 2016) but overall crop rotations tend to  
436 increase soil biodiversity by 3 % and richness by 15 %. Regardless of belowground diversity  
437 trends, crop rotations did create functionally distinct microbial communities in our study (Fig. 4).  
438 We still do not have a complete understanding of how crop rotations alter soil microbial  
439 diversity, nor (arguably more importantly) how these changes in belowground diversity might  
440 provide beneficial soil ecosystem services like increasing soil C or mineralizing more N to  
441 increase crop yields.

442         One trend that emerges across the suite of 31 C substrates is that crop rotations altered the  
443 preference for substrates of differing quality (i.e. complex versus simple C substrates). The soils  
444 from monoculture corn corresponded to greater use of simple C substrates (especially carboxylic  
445 acids), and showed less response to the suite of N-containing and complex substrates. This  
446 finding corroborates a previous study we conducted using whole-plant residues, in which we  
447 showed diverse crop rotations resulted in greater decomposition of low quality crop residues  
448 (e.g. corn and wheat, McDaniel et al. 2014c). Further, when looking only within the relatively  
449 labile carboxylic acid substrates, microbial communities in the less diverse crop rotations (mC,  
450 and CS to a lesser extent) responded to more labile, low-molecular weight carboxylic acids (e.g.  
451 citric, malonic, and malic acid), while soil microbes from more biodiverse crop rotations  
452 responded more to complex, higher-molecular weight carboxylic acids (e.g. caffeic, tartaric, and  
453 vanillic acids - Fig. S4). The strong effects of crop biodiversity on both overall usage of and the



454 types of catabolized carboxylic acids is not surprising due to the small, yet dynamic pool of these  
455 compounds in soil (Strobel 2001). Since soil microbial function (as measured by CLPPs) is an  
456 aggregate measure of both the community composition and available resources, it is impossible  
457 to tease out which (or both) have changed due to increased crop biodiversity. However, our  
458 overall findings indicate that increased aboveground biodiversity through crop rotations and  
459 cover crops appears to facilitate soil microbial communities' use of complex C substrates relative  
460 to simple ones.

#### 461 *Conclusions*

462 As the growing population is increasingly reliant on soils for food, fiber, and fuel we will  
463 either need to consume less, put more land into production, or better use the land we already  
464 have in production. Putting more land in production will likely result in declines in local and  
465 global biodiversity. Thus, it is critical to incorporate biodiversity through any means possible  
466 into the existing managed ecosystems – even including biodiversity through time as with crop  
467 rotations. Here we show that both microbial biomass and function are strongly influenced by  
468 cropping diversity. In fact, the influence of crop rotations on soil microbes and functioning lasts  
469 over an entire growing season and even when all soils are under the same crop. Crop rotations  
470 clearly enhance soil microbial biomass and activity, which are now considered a pillar of soil  
471 health. Furthermore, this rotation effect on soils also appears to facilitate microbes in supplying  
472 more N to crops (Fig. S6). Overall, our study highlights the importance of incorporating  
473 biodiversity into agroecosystems by including more crops in rotation, especially cover crops, to  
474 enhance beneficial soil processes controlled by soil microbes.

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Table 1. Soil carbon (C) and nitrogen (N) pools by season and crop rotation

Season	Crop Rotation	Total Organic C g kg <sup>-1</sup>	Total N	NO <sub>3</sub> <sup>-</sup> -N mg kg <sup>-1</sup>	NH <sub>4</sub> <sup>+</sup> -N	DOC	DON	C:N	DOC:DON
Spring									
	mC	8.1 (0.8)	0.8 (0.1)ab	2.66 (0.79)	0.06 (0.01)B	14 (4)bb	5 (1)bb	9.8 (0.3)	2.8 (0.2)B
	CS	7.8 (1.2)	0.8 (0.1)ab	2.97 (1.13)	0.06 (0.01)B	11 (1)abB	5 (1)bb	10.3 (0.4)	2.1 (0.2)B
	CSW	7.0 (0.6)	0.7 (0.1)b	2.67 (0.39)	0.10 (0.02)B	21 (8)abB	6 (1)abB	10.4 (0.4)	4.2 (1.9)B
	CSW1	8.7 (0.4)	0.9 (0.1)a	3.10 (0.66)	0.10 (0.02)B	44 (18)ab	8 (1)ab	9.6 (0.2)	5.4 (2.6)B
	CSW2	8.2 (1.4)	0.8 (0.1)ab	3.49 (0.62)	0.12 (0.03)B	26 (7)abB	8 (2)ab	10.2 (0.2)	3.3 (0.4)B
Summer									
	mC	7.9 (0.8)	0.8 (0.1)ab	5.58 (0.67)c	0.08 (0.02)A	35 (4)bb	18 (1)ba	10.2 (0.4)	2.0 (0.1)C
	CS	7.6 (0.9)	0.8 (0.1)ab	9.47 (1.96)b	0.08 (0.01)A	32 (4)abB	33 (7)ba	9.8 (0.1)	1.0 (0.1)C
	CSW	7.6 (0.7)	0.8 (0.0)b	7.76 (0.75)b	0.08 (0.01)A	43 (7)abB	28 (4)abA	9.7 (0.3)	1.6 (0.3)C
	CSW1	8.1 (0.8)	0.9 (0.1)a	16.68 (0.87)a	0.37 (0.22)A	88 (32)ab	76 (8)ba	9.0 (0.2)	1.2 (0.4)C
	CSW2	8.7 (1.1)	0.9 (0.1)ab	12.14 (4.03)ab	0.34 (0.12)A	54 (7)abB	68 (13)ba	9.5 (0.1)	0.8 (0.1)C
Autumn									
	mC	8.1 (0.6)	0.7 (0.1)ab	1.31 (0.15)	0.07 (0.02)B	58 (21)ba	5 (1)bb	11.4 (0.3)	14.3 (7.3)A
	CS	7.7 (1.1)	0.7 (0.1)ab	1.44 (0.28)	0.06 (0.01)B	46 (15)abA	5 (1)bb	10.9 (1.0)	9.6 (3.2)A
	CSW	7.4 (0.8)	0.7 (0.1)b	1.28 (0.30)	0.08 (0.02)B	117(77)abA	6 (2)abB	10.6 (0.6)	15.6 (5.2)A
	CSW1	9.6 (0.6)	0.9 (0.0)a	1.41 (0.06)	0.05 (0.01)B	102 (27)ba	7 (1)ab	10.6 (0.5)	17.1 (7.2)A
	CSW2	8.9 (0.9)	0.9 (0.1)ab	0.96 (0.15)	0.05 (0.01)B	190 (42)abA	6 (1)ab	10.4 (0.4)	30.4 (4.0)A
ANOVA Factor		<i>P</i> values							
Season		0.756	0.769	< <b>0.001</b>	<b>0.004</b>	< <b>0.001</b>	< <b>0.001</b>	0.213	< <b>0.001</b>
Rotation		0.298	<b>0.040</b>	< <b>0.001</b>	0.084	<b>0.038</b>	< <b>0.001</b>	0.223	0.947
Season*Rotation		0.994	0.928	< <b>0.001</b>	0.071	0.965	0.221	0.746	0.192

Note: Significant comparisons are shown among Rotations (lowercase) and Season (capital) with letters.



Table 2. Soil extracellular enzyme activities (EEA) expressed as nano-moles of product per hour per gram of dry soil.

Season	Rotation	Soil Extracellular Enzyme Activity (nmol hr <sup>-1</sup> g <sup>-1</sup> soil)							
		BGase	CBHase	LAPase	NAGase	PHOSase	TAPase	PPOase	PERase
Spring									
	mC	94 (8)b	27 (2)b	24 (4)bA	27 (2)ab	133 (19)bC	10 (1)abA	140 (47)B	614 (12)a
	CS	107 (18)b	28 (5)b	28 (4)abA	20 (2)b	129 (20)bC	11 (0)abA	100 (30)B	634 (53)a
	CSW	118 (12)ab	31 (4)ab	26 (8)abA	33 (2)ab	152 (7)abC	12 (2)bA	92 (27)B	602 (59)ab
	CSW1	148 (5)a	50 (5)a	43 (5)abA	47 (3)a	188 (17)aC	16 (1)aA	87 (13)B	516 (24)b
	CSW2	153(13)ab	56 (12)ab	33 (5)aA	48 (5)a	208 (8)aC	16 (1)aA	137 (61)B	562 (24)b
Summer									
	mC	100 (5)b	37 (3)b	7 (2)bB	43 (4)	270 (42)bA	9 (2)abB	174 (67)B	676 (88)a
	CS	111 (17)b	43 (10)b	14 (3)abB	44 (7)	291 (25)bA	9 (1)abB	140 (50)B	580 (124)b
	CSW	102 (7)ab	47 (12)ab	14 (2)abB	47 (3)	280 (13)abA	7 (2)bB	96 (29)B	578 (68)b
	CSW1	146 (12)a	61 (10)a	20 (3)abB	69 (10)	370 (45)aA	14 (1)aB	236 (91)B	317 (144)bc
	CSW2	132 (17)ab	62 (14)ab	13 (4)aB	59 (9)	400 (56)aA	12 (1)aB	126 (73)B	392 (97)c
Autum									
	mC	111 (9)b	44 (6)b	5 (3)bB	67 (13)	238 (57)bB	14 (3)abA	330 (77)A	543 (113)a
	CS	110 (17)b	42 (8)b	8 (1)abB	55 (7)	209 (36)bB	11 (2)abA	234 (64)A	461 (103)bc
	CSW	115 (19)ab	49 (15)ab	9 (2)abB	54 (9)	245 (34)abB	14 (2)bA	176 (18)A	517 (150)b
	CSW1	138 (10)a	59 (6)a	8 (1)abB	63 (13)	277 (42)abB	18 (2)aA	300 (30)A	396 (76)c
	CSW2	117 (15)ab	46 (8)ab	17 (3)aB	63 (2)	308 (24)abB	18 (2)aA	202 (51)A	336 (49)c
ANOVA Factor		<i>P</i> values							
	Season	0.775	0.063	<0.0001	<0.0001	<0.0001	0.003	<0.0001	<0.0001
	Rotation	0.017	0.006	0.007	<0.0001	0.0003	0.002	0.224	<0.0001
	Season*Rotation	0.852	0.839	0.314	<0.0001	0.967	0.647	0.837	<0.0001

Note: Significant comparisons are shown among Rotations (lowercase) and Season (capital) with letters.



Table 3. Analysis of variance of results from the principal components analysis of community-level physiological profile (Fig. 4).

ANOVA <sup>§</sup> Parameter	PC1		PC2		PC3		PC4		PC5		MANOVA (Total)	
Proportion of variance	38.7		17.7		14.5		9		3.8		83.7	
ANOVA Factor	F	<i>P</i> value	F	<i>P</i> value	F	<i>P</i> value	F	<i>P</i> value	F	<i>P</i> value	F	<i>P</i> value
Season	<b>64.02</b>	<b>&lt; 0.001</b>	<b>22.57</b>	<b>&lt; 0.001</b>	<b>5.4</b>	<b>0.008</b>	0.68	0.510	<b>10.33</b>	<b>&lt; 0.001</b>	<b>33.28</b>	<b>&lt; 0.001</b>
Crop Rotation	0.69	0.605	<b>3.03</b>	<b>0.028</b>	<b>12.82</b>	<b>&lt; 0.001</b>	0.36	0.834	1.81	0.146	<b>2.19</b>	<b>0.003</b>
Season*Rotation	0.16	0.995	1.22	0.311	0.55	0.81	0.88	0.544	0.27	0.973	0.65	0.949
Significant comparisons <sup>‡</sup>	1=3≠2		1=2≠3, CS ≠ CSW1		1=2≠3, mC=CS≠CSW= CSW2				1≠2=3,			

§ Degrees of freedom: Season = 2, Crop Rotation = 4, Season\*Rotation = 8.

‡ Significant comparison abbreviations: 1 = spring, 2 = summer, 3 = autumn



Table 4. Catabolic evenness by season and crop rotation (showing full suite of C substrates, without carboxylic acids, and carboxylic acids only).

Season	Rotation	Catabolic Evenness		
		Full	No Carboxylic Acids	Carboxylic Acids Only
<b>Spring</b>				
	mC	24.37 (0.79)A	20.20 (0.05)A	7.60 (0.23)A
	CS	23.79 (0.91)A	19.80 (0.15)A	7.21 (0.13)ab
	CSW	22.98 (0.63)A	19.65 (0.15)B	6.56 (0.35)B
	CSW1	24.28 (0.44)A	18.95 (0.19)ab	6.91 (0.12)ab
	CSW2	24.52 (0.72)A	19.75 (0.24)B	6.90 (0.31)B
<b>Summer</b>				
	mC	14.99 (1.61)B	18.95 (0.59)A	4.91 (0.54)A
	CS	12.86 (1.77)B	20.20 (0.18)A	4.32 (0.38)ab
	CSW	12.10 (1.02)B	19.82 (0.54)B	3.93 (0.20)B
	CSW1	13.83 (1.65)B	18.59 (0.83)ab	4.34 (0.50)ab
	CSW2	12.78 (0.92)B	19.24 (0.51)B	3.75 (0.11)B
<b>Autumn</b>				
	mC	25.81 (0.79)A	19.62 (0.16)A	8.47 (0.24)A
	CS	25.82 (0.55)A	19.11 (0.22)A	8.41 (0.22)ab
	CSW	25.71 (0.74)A	18.98 (0.28)B	8.12 (0.61)B
	CSW1	27.41 (0.63)A	18.63 (0.12)ab	8.90 (0.24)ab
	CSW2	26.08 (0.67)A	18.17 (0.28)B	8.11 (0.08)B
<b>ANOVA Factor</b>				
	Season	< <b>0.001</b>	<b>0.002</b>	< <b>0.001</b>
	Treatment	0.357	<b>0.035</b>	<b>0.028</b>
	Season*Treatment	0.928	0.058	0.807

Note: Significant comparisons are shown among Rotations (lowercase) and Season (capital) with letters.



Table 5. Pearson correlation coefficients between soil properties and community-level physiological profile (CLPP) parameters.

Soil Variable	Substrate Guilds					Catabolic Evenness		
	Amino acids	Amine	Carboxylic Acids	Carbohydrates	Complex	Full	No Carboxylic Acids	Only Carboxylic Acids
Water content	ns	ns	ns	ns	ns	0.40	ns	<b>0.52</b>
pH	0.27	<b>0.43</b>	<b>-0.41</b>	ns	<b>0.53</b>	<b>0.68</b>	ns	<b>0.74</b>
Sand	<b>-0.36</b>	ns	0.28	-0.27	ns	ns	ns	ns
Silt	0.30	ns	ns	ns	ns	ns	ns	ns
Clay	ns	ns	ns	ns	ns	ns	-0.33	ns
Total C	ns	ns	ns	ns	ns	ns	<b>-0.40</b>	ns
Total N	ns	ns	ns	ns	ns	ns	<b>-0.40</b>	ns
C-to-N ratio	ns	0.27	ns	ns	0.30	<b>0.45</b>	ns	<b>0.53</b>
NH <sub>4</sub> <sup>+</sup>	ns	-0.31	0.33	ns	<b>-0.37</b>	<b>-0.40</b>	ns	<b>-0.38</b>
NO <sub>3</sub> <sup>-</sup>	<b>-0.58</b>	<b>-0.55</b>	<b>0.66</b>	-0.30	<b>-0.72</b>	<b>-0.74</b>	ns	<b>-0.70</b>
PMC	ns	0.29	ns	ns	ns	ns	<b>-0.63</b>	ns
PMN	ns	-0.27	0.32	ns	<b>-0.55</b>	<b>-0.49</b>	ns	<b>-0.52</b>
MBC	0.31	<b>0.49</b>	<b>-0.37</b>	ns	ns	<b>0.41</b>	<b>-0.38</b>	<b>0.47</b>
MBN	<b>0.36</b>	0.34	<b>-0.37</b>	<b>0.42</b>	ns	<b>0.36</b>	ns	0.31
MBC:MBN	ns	<b>0.40</b>	ns	ns	ns	0.31	-0.34	<b>0.40</b>
BGase	ns	<b>-0.43</b>	<b>0.30</b>	ns	ns	-0.29	0.32	-0.28
CBHase	-0.32	<b>-0.47</b>	<b>0.39</b>	-0.27	ns	-0.33	ns	-0.28
LAPase	ns	-0.29	ns	ns	ns	ns	<b>0.49</b>	ns
TAPase	ns	<b>-0.37</b>	ns	ns	ns	ns	ns	0.37
NAGase	<b>-0.35</b>	<b>-0.56</b>	<b>0.47</b>	<b>-0.39</b>	-0.29	<b>-0.46</b>	<b>0.29</b>	<b>-0.41</b>
PHOSase	<b>-0.45</b>	<b>-0.66</b>	<b>0.56</b>	<b>-0.46</b>	-0.34	<b>-0.63</b>	<b>0.34</b>	<b>-0.60</b>
PPOase	-0.38	-0.33	0.37	-.31	ns	ns	ns	ns
PERase	-0.40	-0.54	0.42	-0.37	ns	-0.30	0.43	ns

Note: Only significant correlations are shown ( $P$  values < 0.05), ns = non-significant



## FIGURES

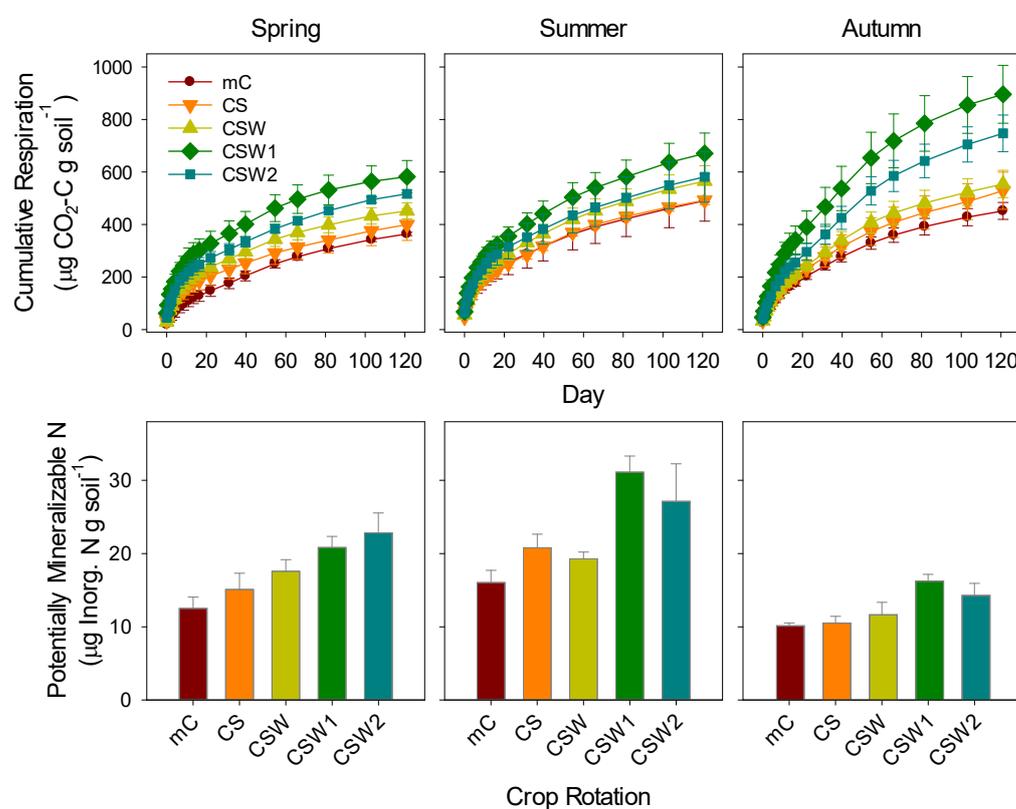


Figure 1. Potentially mineralizable carbon (top row) and potentially mineralizable nitrogen (bottom row). Crop rotation abbreviations are: monoculture corn (mC), corn-soy (CS), corn-soy-wheat (CSW), corn-soy-wheat with red clover cover crop (CSW1), and corn-soy-wheat with red clover + rye cover crops (CSW2). Means are shown and error bars are standard errors (n = 4).

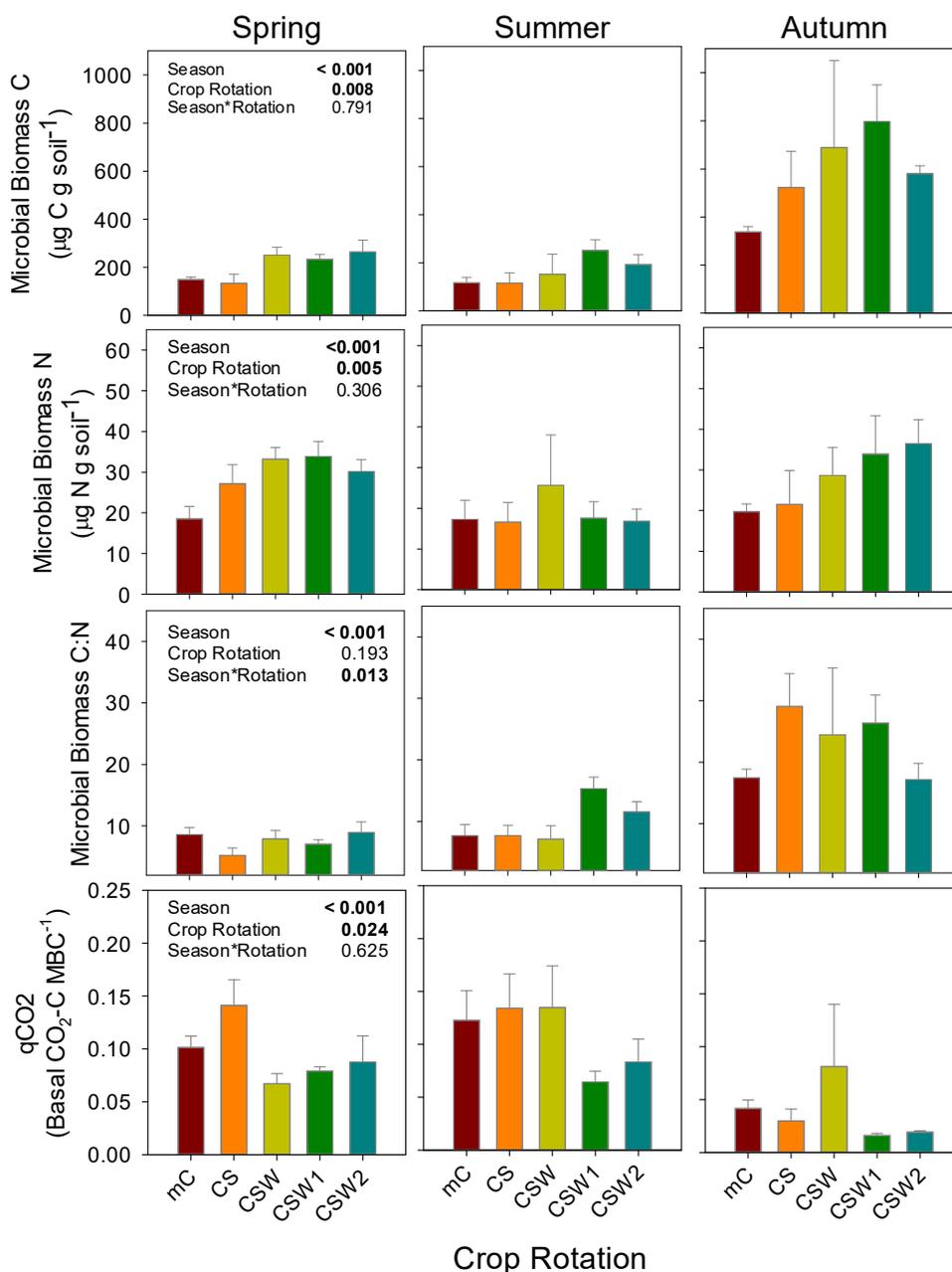


Figure 2. Soil microbial biomass parameters by season and crop rotation. See Fig.1 for crop rotation abbreviations. Means are shown and error bars are standard errors (n = 4).

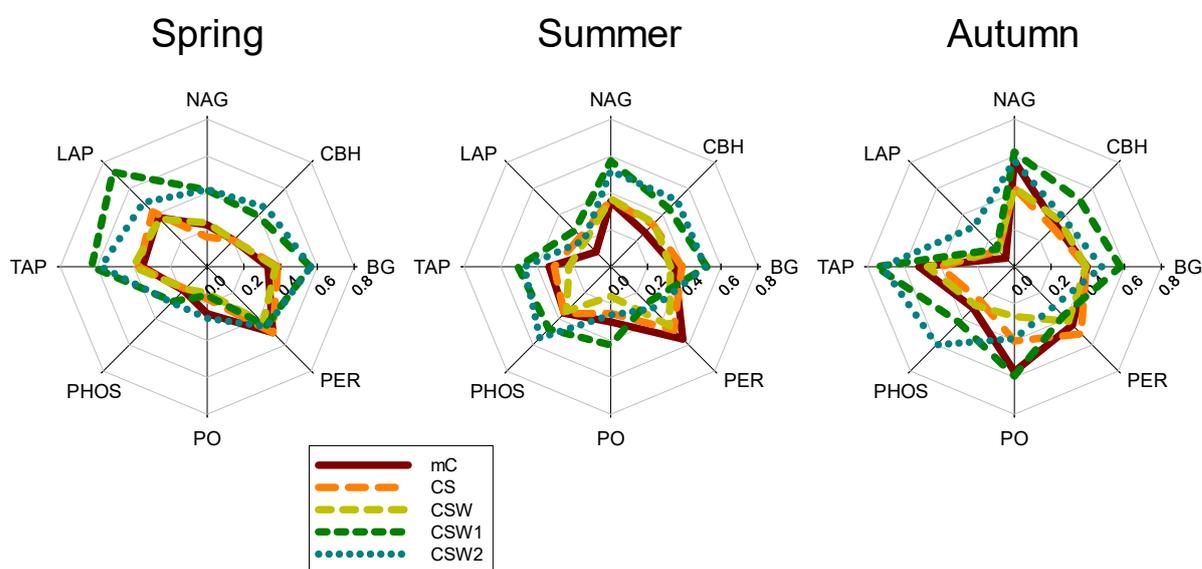


Figure 3. Extracellular enzyme activities (EEA) normalized for the maximum value during each season. EEA abbreviations are:  $\beta$ -1,4,-glucosidase (BG),  $\beta$ -D-1,4-cellobiohydrolase (CBH),  $\beta$ -1,4,-N-acetyl glucosaminidase (NAG), acid phosphatase (PHOS), Tyrosine aminopeptidase (TAP), Leucine aminopeptidase (LAP), phenol oxidase (PO), and peroxidase (PER). See Fig.1 for crop rotation abbreviations.

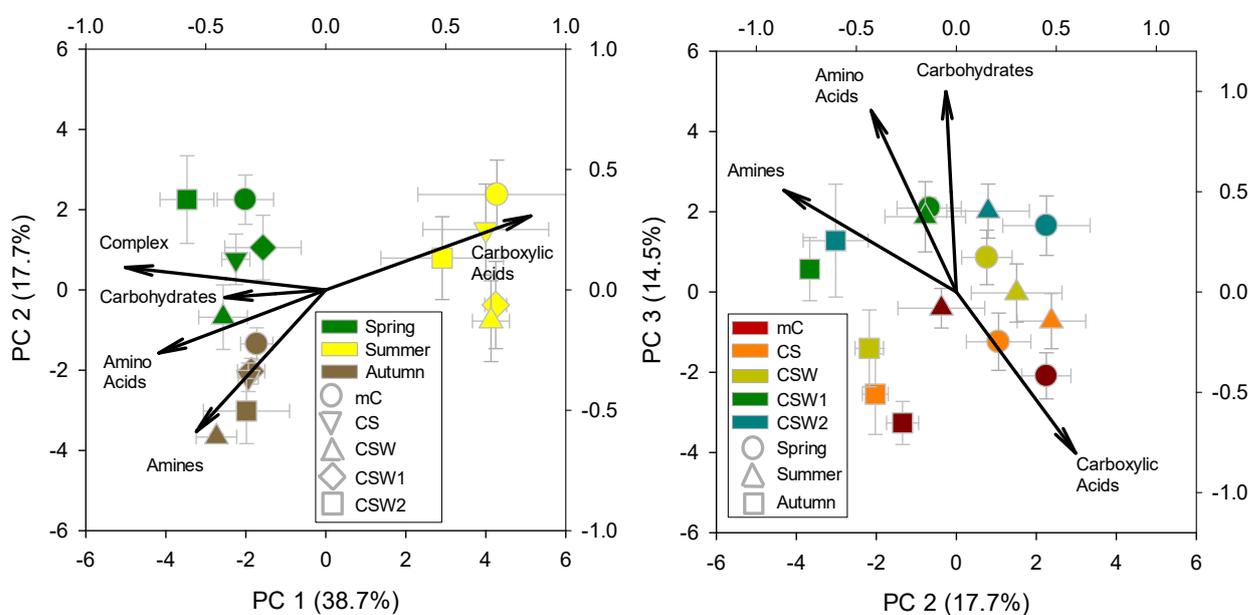


Figure 4. Principal components analysis (PCA) on all 31 substrates. *Left Panel:* Principal components 1 and 2, where Season is dominant discriminating factor ( $P < 0.001$ ) and *Right Panel:* Principal components 2 and 3 where Rotation is highlighted as a dominant discriminating factor. See also Table 5 for PCA and ANOVA results. Means are shown and error bars are standard errors ( $n = 4$ ). See Fig.1 for crop rotation abbreviations.