

1 On the rebound: soil organic carbon stocks can bounce back to near
2 forest levels when agroforests replace agriculture in southern India

3

4 **H. C. Hombegowda^{1, 2*}, Oliver van Straaten^{3*}, Michael Köhler¹ and Dirk Hölscher¹**

5 [*] {The authors contributed equally to this work}

6 [1] {Tropical Silviculture and Forest Ecology, Georg-August-Universität Göttingen, Germany}

7 [2] {Indian Council of Agricultural Research, New Delhi, India}

8 [3] {Soil Science of Tropical and Subtropical Ecosystems, Georg-August-Universität
9 Göttingen, Germany}

10 Corresponding author: H. C. Hombegowda (hombegowdaars@gmail.com)

Abstract

Tropical agroforestry has an enormous potential to sequester carbon while simultaneously producing agricultural yields and tree products. The amount of soil organic carbon (SOC) sequestered is influenced by the type of the agroforestry system established, the soil and climatic conditions and management. In this regional scale study, we utilized a chronosequence approach to investigate how SOC stocks changed when the original forests are converted to agriculture, and then subsequently to four different agroforestry systems (AFSs): homegarden, coffee, coconut and mango. In total we established 224 plots in 56 plot clusters across four climate zones in southern India. Each plot cluster consisted of four plots: a natural forest reference plot, an agriculture reference and two of the same AFS types of two ages (30-60 years and >60 years). The conversion of forest to agriculture resulted in a large loss the original SOC stock (50-61%) in the top meter of soil depending on the climate zone. The establishment of homegarden and coffee AFSs on agriculture land caused SOC stocks to rebound to near forest levels, while in mango and coconut AFSs the SOC stock increased only slightly above the agriculture stock. The most important variable regulating SOC stocks and its changes was tree basal area, possibly indicative of organic matter inputs. Furthermore, climatic variables such as temperature and precipitation, and soil variables such as clay fraction and soil pH were likewise all important regulators of SOC and SOC stock changes. Lastly, we found a strong correlation between tree species diversity in homegarden and coffee AFSs and SOC stocks, highlighting possibilities to increase carbon stocks by proper tree species assemblies.

Key words: land-use change, carbon sequestration, homegarden, coffee, coconut, mango

1 **1 Introduction**

2 Land-use changes in the tropics are responsible for approximately 10% of the human
3 induced greenhouse gas emissions and are expected to remain the second largest source of
4 carbon (C) emissions in the near future (Achard et al., 2014). Considering that tropical forest
5 soils store a similar amount of organic carbon (692 Gt in the top 3-m; Jobbágy and Jackson,
6 2000) as the atmosphere (589 Gt C; Ciais, et al., 2013), and that tropical climates foster rapid
7 organic matter decomposition, land-use changes can result in strong carbon fluxes into or
8 out of the soil. The conversion of tropical forests to agriculture causes both a release of
9 stored soil organic carbon, often in the form of carbon dioxide, but it also results in a decline
10 in soil productivity. To reduce carbon emissions from agriculture while simultaneously
11 maintaining agricultural productivity it is necessary to identify and implement simple and
12 cost effective measures to store and capture carbon. In this context agroforestry practices,
13 which integrate trees into agricultural systems, offer a unique opportunity to sequester
14 atmospheric carbon while also growing food, diversifying incomes (e.g. from sale of wood,
15 fruit and staple foods), and simultaneously providing numerous environmental benefits.
16 These include mitigating soil erosion (Montagnini and Nair, 2004), improving soil structure
17 (Lal, 2007), pumping up nutrients from the subsoil (Das and Chaturvedi, 2008) and
18 sequestering atmospheric carbon (Lal, 2007; Nair et al., 2009). Agroforestry systems (AFSs)
19 have higher SOC sequestration rates than conventional agricultural systems (Nair et al.,
20 2009) as the trees they have comparatively higher litter inputs and are capable of inserting
21 carbon deep in the soil with their root systems (Montagnini and Nair, 2004).

22 Furthermore, tree species diversity in AFSs can have a large impact on organic matter
23 turnover as diverse species mixtures can add different qualities of organic matter which
24 correspondingly influence soil microbial communities and decomposition processes (Six et
25 al., 2002; Acker et al., 2002). Niche differentiation and resource partitioning may lead to a
26 better use of space and nutrient uptake and thus increase ecosystem carbon inputs (Thakur.,
27 et al., 2015) . Although it is recognized that AFSs have many benefits, their C sequestration
28 potential, especially belowground, remains largely unexplored (Montagnini and Nair, 2004).
29 Concentration and SOC turnover rates in AFSs vary significantly with biophysical site
30 properties such as climate (Liu et al., 2011), vegetation, land-use types (Cadotte, 2013; Saha
31 et al., 2010), soil type and texture (Six et al., 2002; Chaplot et al., 2010), land management
32 (Hevia et al., 2003) and their interactions (Powers and Schlesinger, 2002). It is generally

1 recognized that temperature and precipitation are the most important variables regulating
2 SOC (Chaplot et al., 2010; Liu et al., 2011). Since both affect the type of vegetation cover, the
3 quantity of biomass production and the rate of SOM turnover (Hevia et al., 2003).

4 It is estimated that there are approximately 25.3 million hectares (Dhyani et al., 2013) of
5 AFSs established across India, whereby the type of AFS established depends on the
6 biophysical site conditions and the socioeconomic status of the owners. Despite this,
7 regional-scale studies evaluating the impacts the establishment these land-use types have
8 on SOC stocks remain relatively scarce (Albrecht and Kandji, 2003; Mutuo et al., 2005, Saha
9 et al., 2010). Here in this study, we quantified SOC changes associated with the conversion of
10 forest to agriculture and subsequently from agriculture to four different AFS types
11 (homegarden, coffee, coconut and mango). Our plots were established across southern India
12 in a broad range of biophysical conditions, ranging from semi-arid to humid climates and in
13 soils with low and high activity clays. The objectives of this study were as follows:

- 14 (i) to quantify SOC stocks and changes to SOC stocks along a forest – agriculture –
15 AFS trajectory, and
- 16 (ii) to determine the biophysical drivers regulating SOC stocks and its changes.

17 **2 Materials and methods**

18 **2.1 Study area**

19 The study was conducted across three states in Southern India (Kerala, Karnataka and
20 Andhra Pradesh; Fig. 1), in an area that has a long-standing history of diverse agroforestry-
21 based land-use practices. The region has a tropical monsoon climate, with a rainy season
22 from May to October and a distinct dry season from November to April. The mean annual
23 precipitation (MAP) ranges from 627 to 3422 mm and mean annual temperature (MAT)
24 ranges between 21.9°C in the highlands to 27.2°C at lower elevations (Hijmans et al., 2005).
25 The soils were classified as Luvisols, Acrisols or Nitisols.

26 **2.2 Sampling design and site selection**

27 In this study we investigated how SOC stocks changed when forests are converted to
28 agriculture and subsequently to agroforestry systems in four different climatic zones along a
29 precipitation gradient (humid, moist sub-humid, dry sub-humid and semi-arid; based on

classification by ICAR (1984)). Each AFS type was sampled in two of four climatic zones based on its relative importance in terms of area, production and income in the region (Table 1). Using a chronosequence approach we established 56 plot clusters (Fig. 1). A plot cluster consisted of four land-use types: a natural forest reference and an agriculture reference and two AFS plots at different ages, medium (30-60 years) and old (>60 years). Seven plot clusters were established for each AFS type in two climatic zones. In total 224 plots were set up (4 AFSs x 2 climate zones x 7 clusters x 4 land-use types).

The plot clusters were always centered around the natural forest reference plot and located within a maximum distance of 1-km of each other. To reduce edge effects, the forest plots were selected at least 40 m from the forest edge. Plot pairs (either forest with agriculture or agriculture with AFS) were carefully selected to ensure soil and climate conditions were similar: 1. All plots were located on similar landscape positions on flat to gently sloping terrain (average: 3%; maximum: 7%). 2. Using a feel test we compared the subsoil texture of all possible candidate sites and chose only those with comparable soil textures. An *a posteriori* texture analysis revealed that there were small differences in surface clay percent at 10-30 cm. For the forest to agriculture comparison, the clay content difference was $-5.5 \pm 0.5\%$ ($P > 0.05$), and for the agriculture to AFS the difference was $2.3 \pm 0.4\%$ ($P > 0.05$). 3. Selected sites were all well drained, had deeply weathered soils (no stones) and had not been limed.

2.3 Land-use systems and forest reference

Each AFS type consisted of a unique combination of trees and crops (Table 1) and originated from former agricultural land with the exception of 19 coffee plots and one homegarden plot which replaced forests directly. Pictures of the investigated land-use types in the different climate zones are found in Fig. A1 in Appendix A.

The homegarden AFS has a multilayered canopy, consisting of a diverse tree-admixture of different ages and sizes. The multipurpose trees grown here are found in association with shrubs and herbaceous species (Kumar et al., 1994). The majority of households in the humid and moist sub-humid climate zones manage homegarden AFS to (partially) satisfy their fruits, spice and vegetable needs (such as cassava, banana and ginger). For management, farmers loosen the soil once a year during vegetable cultivation (Appendix A, Figure A1).

Coffee AFSs are mainly grown in the humid and moist sub-humid region of the Western Ghats. Coffee is grown in the understory of both native shade tree species and planted trees and is often inter-cropped with spices such as pepper, cardamom, cinnamon, clove and nutmeg (Appendix A, Figure A1). Litterfall from shade trees and pruning products mean that this system receives substantial organic matter inputs. The soils in coffee AFS are typically hand tilled once every two to three years.

Coconut is primarily grown by smallholder farmers in southern India. The four southern states Karnataka, Andhra Pradesh, Tamil Nadu and Kerala together produce 92% of India's total coconut production. In the humid region of Kerala state, coconut is extensively grown with vegetables, whereas in the dry sub-humid region it is grown together with food grains (such as maize, turmeric and finger millet) (Appendix A, Figure A1). These crops are cultivated at different stages of coconut plantation development depending on the amount of incoming light. Since most of the coconut plant parts (leaves, stems) are useful to the farmers for various household uses, little organic matter is left onsite. In dry sub-humid zone, farmers plough the land two to three times a year for agriculture crop production.

Mango is an important commercial fruit crop in India which ranks first among world's mango producing countries, accounting for about 40% of the world production (Sekhar et al., 2013). In Karnataka and Andhra Pradesh states, mango is predominantly grown in drier climate (dry sub-humid and semi-arid) and is predominately cultivated by smallholder farmers. Many farmers adopt wide row spacing and grow agriculture crops between the rows (Appendix A, Figure A1). Since it normally takes 12-15 years to establish a closed canopy, farmers utilize this duration to cultivate agriculture crops, such as finger millet and maize. In later stages, those crops are no longer profitable and farmers switch to grow fodder and short rotation crops for the household consumption.

Agriculture under humid and moist sub-humid climate zones is only practiced by smallholder farmers. Here, staple foods like tubers and vegetables for subsistence use are grown, with only organic matter inputs as nutrient supplement sources (Appendix A, Figure A1). In the dry sub-humid and semi-arid climate zones commercial agriculture crops are grown with a high input of both fertilizers and pesticides. All agricultural fields are typically ploughed two to four times a year, often with a small tractor. During agriculture establishment the forests

1 were cleared using hand tools and most above-ground biomass was removed for domestic
2 use. Sites were not burnt.

3 The forest plots of our study were most often community managed and considered “sacred
4 groves”. Due to religious reasons people do not remove any wood from there. The
5 remaining forests which we sampled were government owned; these too were relatively
6 undisturbed. All forests were located in or within the vicinity of the village and were highly
7 protected. Evergreen forests were found in humid climate, moist deciduous forest in moist
8 sub-humid climate zone, dry deciduous forest in dry sub-humid and scrub forest in semi-arid
9 climate (Appendix A, Figure A1).

10 **2.4 Sampling and lab analyses**

11 In each 20 x 20m plot we took soil samples using a soil auger from 12 fixed locations around
12 the plot at predefined soil depths (0-10, 10-30, 30-60 and 60-100 cm). The samples for each
13 respective depth were pooled and thoroughly mixed. A soil pit (1 x 1 x 1 m) was dug in the
14 center of the plot for soil bulk density determination by embedding a cylindrical core (165
15 cm³, diameter of 5.3 cm, height of 7.5 cm) at 5, 20, 45 and 80 cm depths and replicated
16 twice per depth.

17 Total soil carbon and nitrogen concentrations were analyzed using a CN Elemental Analyser
18 (Vario EL III, Elementar, Hanau, Germany). As the entire region is underlain by high grade
19 metamorphic rocks and granites of the Indian Shield no carbonates were expected in these
20 soils and no attempts were made to remove them. The soil carbon stock (Mg C ha⁻¹) was
21 calculated by multiplying the carbon concentration (g kg⁻¹ of soil) with bulk densities of the
22 respective depth interval (kg m⁻³) and the layer thickness (m) and up-scaled to one hectare.
23 Total soil carbon stocks for the top meter of soil were calculated as the sum of all depth
24 intervals. To ensure comparability of plot pairs and to avoid overestimation of SOC stock
25 changes we used the bulk density data of the respective forest plots to calculate the soil
26 carbon stock of the agroforestry and agriculture plots within each the cluster (Veldkamp et
27 al., 1994). Additionally, we determined the pH of air-dried soil in a 1:2.5 soil-to-water
28 solution for all sampling depths, and soil texture for two depths (0-10, 10-30 cm) using the
29 pipette sedimentation method.

At each plot, we measured tree basal area for all trees with a diameter at breast height greater than 10 cm. These tree species were identified to the species level. Furthermore, we recorded information on slope, elevation and geographical coordinates of each plot. Through informal interviews with the land-owners we got information on current and past land-uses and their management practices. Meteorological data such as mean annual temperature and mean annual precipitation for the selected plots was retrieved from the WorldClim database (Hijmans et al., 2005).

2.5 Statistical analysis

To verify that plots were satisfactorily selected in the field and that the soils of the plot pairs were inherently similar we did an *a posteriori* comparison of soil clay percentages in the subsoil (10-30 cm) of the plot pairs using a paired t-test analysis. To estimate the size of SOC stock changes following land-use change (either from forest to agriculture or from agriculture to AFS), we calculated the difference in SOC stocks between plot pairs. The percent-difference in SOC stocks was then expressed as the relative change to the respective reference SOC stock (forest reference for agriculture; agriculture reference for AFSs) ($\text{Relative change} = (\text{SOC}_{\text{converted}} - \text{SOC}_{\text{reference}}) / \text{SOC}_{\text{reference}} \times 100$). The influence of climate and site variables on SOC stocks and relative SOC stock changes was evaluated by linear and non-linear regression analyses across AFSs for single variables and with stepwise linear multivariate analyses for each system.

The residuals of all models were checked for normality with QQ-plots; models were considered significant at the $P \geq 0.05$ level. All statistical analyses were done using the software package R - version 2.15.0 (R Development Core Team, 2014).

3 Results

3.1 Land-use change impacts on SOC stocks

Among the land-use types investigated, SOC concentrations (data not shown) and SOC stocks were highest in natural forest, lowest in agriculture and intermediate for the different AFSs (Fig. 2). Deforestation for agriculture resulted in a strong decrease in SOC stocks across all climate zones, creating a loss of 50% to 61% of the original SOC stock. The resulting SOC

stocks of AFSs however were dependent on AFS type. While SOC stocks in homegarden and coffee AFSs rebounded to near forest levels, the SOC stocks in both coconut and mango AFSs increased only marginally compared to the agricultural reference (Fig. 2 and 3).

Of the AFSs studied, SOC stocks in the top meter of soil were highest in coffee ($156 \pm 10 \text{ Mg C ha}^{-1}$) followed by homegarden ($151 \pm 5 \text{ Mg C ha}^{-1}$), coconut ($98 \pm 7 \text{ Mg C ha}^{-1}$) and lowest in mango AFSs ($76 \pm 3 \text{ Mg C ha}^{-1}$). AFS establishment on agricultural land caused SOC stocks to increase significantly in all AFSs, with SOC gains ranging from ~45% in coconut in dry sub-humid zones to ~103% in homegarden in humid zones (Fig. 3). Furthermore, significant SOC stock gains were measured at all soil depths (with the exception of coffee at 60-100 cm). Homegardens exhibited the highest overall SOC stock gains, with relative changes within each depth layer being relatively uniform throughout the soil profile. This was followed by coffee, which in contrast, had highest SOC stock gains at the soil surface and decreased with depth. Lastly, SOC stock gains in mango and coconut AFSs were comparatively low, but constant throughout the soil profile (Fig. 3). When expressed in terms of overall changes throughout the whole soil profile, most SOC gains were concentrated at the soil surface (see grey bar in Fig. 3). Nevertheless, when considering the whole soil profile (which has a much bigger volume), there are large SOC gains below 30 cm. For homegardens, 58% of the gains occurred below 30 cm, for coffee 26%, for coconut 59%, and for mango 50%.

The climate zone had a marginal influence on overall SOC stocks of a given AFS where only coconut and mango AFSs showed a tendency towards higher SOC stocks in the respective wetter climate zone (Fig. 2). Also the SOC stocks of the AFSs did not differ among the two age categories sampled (Fig. 2).

3.2 Predicting SOC stocks and relative changes in SOC stocks

In undisturbed forest ecosystems 82% of the variance in SOC stocks could be significantly explained by five variables (MAP, MAT, basal area, clay fraction and soil pH) using stepwise multivariate regression analyses (Table 2). While tree basal area and clay fraction exhibited positive linear correlations with SOC stock, both soil pH and MAP exhibited parabolic and inverse parabolic relationships respectively (Fig. 4a-d). For pH, SOC stocks were lowest at a near-neutral conditions, while SOC stocks peaked between 2000 and 3000 mm yr⁻¹ MAP.

Basal area was the single best predictor of forest SOC stocks. This is evident in both, the scatterplot graphs in Fig. 4b and from its dominant position in the multivariate regressions for all land uses investigated (Table 2), except in agriculture which had no trees. Likewise, soil clay fraction was also an important predictor of SOC and was present in all AFSs regression equations except homegardens. Although MAT was an influential predictor of forest SOC stocks in the stepwise regression (found in 3 of the 6 prediction models; Table 2), yet taken as single predictor its influence on SOC stocks remains insignificant. SOC stocks in both homegarden and coffee AFSs are further positively correlated to the Shannon Wiener species diversity index of trees (Fig. 5). Coconut and mango AFSs however were not included as they are monocultures in terms of tree or palm species admixture.

The SOC stock losses attributed to the conversion of forests to agriculture could be predicted by two variables: clay fraction and MAT (Table 2). Thereafter, when agriculture plots were converted to AFSs, SOC stock changes could be predicted by MAP, MAT, basal area, clay fraction and soil pH in varying importance for the individual AFSs and for all AFSs combined (Table 2). While MAP, basal area, and clay fraction all exhibited positive linear correlations with SOC stock change, soil pH exhibited a parabolic relationship with SOC change. For the latter, SOC stock losses were highest in acidic soils, lowest near neutral pH, and again higher in slightly alkaline soils (Fig. 4e-h).

4 Discussion

4.1 SOC stocks in natural forests

The SOC stock is a balance of incoming carbon from organic matter and carbon losses either through decomposition processes or dissolved organic carbon (DOC) leaching (Davidson and Janssens, 2006; Raich et al., 2006). Both the carbon inputs and outputs however are strongly affected by ecosystem productivity, vegetation type, climate, clay mineralogy, soil pH, nutrient availability, soil aggregates and texture (Lal, 2004; Six et al., 2002; Chaplot et al., 2010; Don et al., 2011).

Although we did not measure organic matter inputs directly, we found a very strong correlation between SOC stocks and plot basal area, which could be indicative of organic matter litter inputs (Fig. 4b; Lebrete et al., 2001). Once organic matter enters the soil, the soil's physical characteristics and biochemical environment plays an important role in how

SOC is either stabilized, mineralized or leached (Raich and Schlesinger, 1992; Six et al., 2002; Zinn et al., 2007a). The clay mineralogy and soil clay fractions play a critical role in how much SOC can be stored by the soil, its residence times and its susceptibility to land-use change (Zinn et al., 2007b). Clays can stabilize organic matter particles through clay-humus complexation, or physically protect organic matter molecules from further mineralization by trapping them in clay macroaggregates (Sollins et al., 1996). The strong positive correlation between SOC stocks and clay fraction we measured reflect the importance of these processes on SOC storage (Table 2, Fig. 4c). Next, the parabolic relationship we observed between soil pH and forest SOC stocks (Fig. 4d) reflect how the soil biochemical environment is critically important for soil microbial communities that decompose organic matter (Motavalli et al., 1995). The higher SOC stocks found in the acidic and basic soils in this study indicates that the unfavorable biochemical environment retards microbial communities that decompose the organic matter. At a near-neutral soil pH, conditions were ideal for microbial communities and accordingly decomposition rates were high, resulting in little SOC accumulation.

As expected, regions of higher rainfall stored more SOC than drier zones (Fig. 2). This is likely because the natural forests in the humid zone have higher net primary production given the favorable year-round water availability compared to forests in the drier regions. Unexpectedly however, at very high levels of precipitation ($>3000 \text{ mm y}^{-1}$) SOC stocks declined again (Fig. 4a). We suspect that this decrease at high precipitation is related to the corresponding lower ecosystem biomass (using basal area as a proxy; Fig. A2, Appendix A). Both the decrease in SOC stocks and basal area at high precipitation could be explained by the torrential monsoon rains. Although the overall rainfall amount may be higher, its intensity and distribution over time causes much of it to runoff, which is then not available to plants when they need it. Meanwhile, it has been reported that the leachate from these soils may also contain high DOC concentrations (Lal, 2003).

Although MAT is a significant predictor of forest SOC stocks, its influence can only be evaluated in interaction with other variables (Fig. A3a, Appendix A). Although one would expect a strong correlation between MAT and MAP, it was not present here. While MAT is mainly driven by altitude, MAP depends on more complex weather phenomena that differ among climate zones. This might also explain why there is no clear relationship between MAT and SOC when other variables are not included in the analysis.

4.2 Agriculture establishment causes up to 61% SOC losses

In comparison to a large pool of studies on this land-use conversion conducted in the tropics, the SOC losses detected here for the different climate zones were at the higher end (50-61%). Results from two meta-analyses report SOC stock decreases on average between 18% (Powers et al., 2011) and 25% (Don et al., 2011) for this land-use conversion in the tropics. Powers et al. (2011) however report that large SOC losses are possible (to a maximum of 76%), but the losses depend on the clay mineralogy and precipitation regimes. Nevertheless, in our study we measured significant decreases in SOC stocks irrespective of soil type and precipitation regimes. We primarily attribute these large SOC losses to the frequent tilling and low organic matter inputs. Tillage exposes SOC to microbial activity through the destruction of aggregates which as a result makes SOC complexes vulnerable to decomposition (Six, et al., 2002; Mangalassery et al., 2013). Furthermore, fine grained particles (such as clay) and associated organic matter can also be lost through soil erosion, runoff and leaching (Gonzalez and Laird, 2003) due to a lack of soil protection measures especially following plowing at the onset of the monsoon (Dourte et al., 2012). Associated with the erosional clay losses is a corresponding reduction of the soil's carbon storage potential. However, since all our plots were established on flat to gentle slopes, and because we were only on-site for one day we did not measure soil erosion.

As previously reported by van Straaten et al. (2015), we also found that SOC stock losses were proportional to the initial forest SOC stock, whereby the higher the SOC stock was initially, the larger the corresponding SOC stock loss when converted to agriculture (Fig. A4, Appendix A).

4.3 SOC stocks rebound when agroforestry systems are established

Increases in SOC stocks resulting from agroforestry establishment are ultimately attributed to higher organic matter inputs from above- and belowground sources (leaves, wood, roots, fungi, animals, etc.; Montagnini and Nair, 2004) and a reduction of SOC losses from decomposition and leaching. Similar to the carbon stocks in forests, MAP, basal area, clay fraction and pH also control carbon stock changes when agricultural land is converted to agroforestry (Table 2, Fig. 4 e-h). Furthermore, the accumulation of SOC depends highly on the quality of incoming litter (Lemma et al., 2006) and is reflected in the significantly higher soil C:N ratios in homegarden, coconut and mango AFSs (Fig. A5, Appendix A). In comparison

1 to litter from agricultural crops, which generally have low C:N ratios and decompose rapidly,
2 organic matter inputs from trees are generally of poorer quality (higher C:N ratios) because
3 of the higher lignin and polyphenolic contents, which in turn results in slower decomposition
4 rates and more SOC accumulation (Davidson and Janssens, 2006).

5 The large SOC stocks increases in the subsoil (Fig. 3) indicate that belowground carbon
6 inputs from roots and/or leaching of organic acids and soluble humus fractions to deeper
7 layers are important processes for SOC accumulation. Considering the trees of the four AFSs
8 have deeper rooting profiles than agricultural crops and that often more than half of the
9 carbon assimilated by trees is transported belowground for root production (Montagnini and
10 Nair, 2004; Poeplau and Don, 2013), it is no surprise that SOC stocks also increased
11 substantially at depth. Furthermore, tillage activities will have mixed soils in the top 30-cm
12 and therein homogenized soil carbon concentrations to a certain extent (Yang and Kay,
13 2001). However, the size of the SOC stock change hinges on the type of AFS established and
14 its management practices. While all AFSs gained carbon compared to agriculture, the
15 amounts gained varied strongly between the different types (Fig. 2). While coconut and
16 mango SOC stocks increased just marginally above the agriculture reference, homegarden
17 and coffee SOC stocks rebounded to forest levels. Clearly, the carbon cycling dynamics of
18 both homegarden and coffee AFSs resemble that of natural forests since both AFSs support
19 many different tree species of different ages, have varied stand structures and have high
20 basal areas.

21 However, not only carbon input, but also losses are a function of AFS type, especially in
22 terms of the plantation management schemes. In coconut and mango AFSs (especially in the
23 dry regions), the low SOC stock increases are linked to the removal of crop residues
24 (including leaves) from the site which are used as fodder or fuel. In coconut, even the tree
25 leaves are removed and utilized.

26 The effect of climate zone on SOC stock changes can only be quantified within each AFS
27 type. While the climate zone did not affect SOC stock changes in either coffee or
28 homegarden AFSs, only a slight difference were found for coconut and mango (Fig. 2). This is
29 however more likely attributed to the implementation of different management practices
30 and cannot be disentangled from climate itself. For instance, in humid climates coconut
31 farmers utilize a “planted fallow system” which includes a fallow period in the cropping cycle

1 where organic matter is reintegrated into the soil. Such systems have been shown to
2 improve soil fertility and maintain SOC stocks (Salako et al., 1999). In contrast, in the dry sub-
3 humid climate farmers use a continuous cropping system which inevitably results in lower
4 organic matter inputs and therein lower SOC stocks. The age of each respective AFS also had
5 no significant effect on SOC stocks (Fig. 2), indicating that soil carbon had already reached a
6 new equilibrium within the first 30-60 years. This is consistent with literature which report
7 that a new SOC equilibrium can be attained in 20 to 40 years following land-use conversion
8 (Detwiler, 1986; de Blécourt et al., 2013; Chiti et al., 2014).

9 Lastly, tree species diversity (Shannon Wiener index) correlated strongly positively with SOC
10 stocks in homegarden and coffee AFSs (Fig. 5), highlighting the role that species and
11 resource complementarity have in maximizing biomass production (Cadotte, 2013).
12 Interestingly, both homegarden and coffee AFSs investigated in this study had similar tree
13 diversities as natural forests. Since plant diversity was also positively correlated with basal
14 area (data not shown) it is possible that the favorable climatic conditions where these AFSs
15 exist can allow both high species diversity and high ecosystem productivity. Literature has
16 shown that plant diversity is integrally linked to ecosystem productivity (Cadotte, 2013) and
17 ecosystem resource utilization (Tilman et al., 2012), which both affect SOC storage potential
18 (Thakur et al., 2015). In contrast, monocultures have been shown to have lower organic
19 inputs than species diverse systems (Cardinale et al., 2007)

20 **5 Conclusions**

21 Agroforestry systems provide a unique opportunity to produce food and tree products, while
22 also improving livelihoods, protecting and improving soils and, as discussed here, to
23 sequester carbon. Nevertheless, not all AFSs provide the same benefits. Soils in homegarden
24 and coffee plantations for instance can sequester much more carbon than coconut or mango
25 AFSs. Additionally, the soil carbon sequestration potential of AFSs can be maximized by
26 cultivating a broad range of different tree species, minimizing tillage activities and leaving
27 crop residue on site.

29 **Acknowledgements**

30 We kindly acknowledge financial support from the International Fellowship Program of the
31 Indian Council of Agricultural Research (ICAR), Government of India. We also thank Prof. Dr.

1 N. B. Prakash at the University of Agricultural Sciences in Bangalore, for his support and
2 guidance and our field assistants Jagadish, M. R. and Sathya.

4 **References**

5 Achard, F., Beuchle, R., Mayaux, P., Stibig, H.-J., Bodart, C., Brink, A., Carboni, S., Desclée, B.,
6 Donnay, F., Eva, H. D., Lupi, A., Raši, R., Seliger, R. and Simonetti, D.: Determination of
7 tropical deforestation rates and related carbon losses from 1990 to 2010, *Glob. Change Biol.*,
8 20(8), 2540–2554, doi:10.1111/gcb.12605, 2014.

9 Acker, S. A., Halpern, C. B., Harmon, M. E. and Dyrness, C. T.: Trends in bole biomass
10 accumulation, net primary production and tree mortality in *Pseudotsuga menziesii* forests of
11 contrasting age, *Tree Physiol.*, 22(2-3), 213–217, doi: 10.1093/treephys/22.2-3.213, 2002.

12 Albrecht, A. and Kandji, S. T.: Carbon sequestration in tropical agroforestry systems, *Agric.*
13 *Ecosyst. Environ.*, 99(1-3), 15–27, doi:10.1016/S0167-8809(03)00138-5, 2003.

14 Cadotte, M. W.: Experimental evidence that evolutionarily diverse assemblages result in
15 higher productivity, *Proc. Natl. Acad. Sci.*, 110(22), 8996–9000,
16 doi:10.1073/pnas.1301685110, 2013.

17 Cardinale, B. J., Wright, J. P., Cadotte, M. W., Carroll, I. T., Hector, A., Srivastava, D. S.,
18 Loreau, M. and Weis, J. J.: Impacts of plant diversity on biomass production increase through
19 time because of species complementarity, *Proc. Natl. Acad. Sci.*, 104(46), 18123–18128,
20 2007.

21 Chaplot, V., Bouahom, B. and Valentin, C.: Soil organic carbon stocks in Laos: spatial
22 variations and controlling factors, *Glob. Change Biol.*, 16(4), 1380–1393, doi:10.1111/j.1365-
23 2486.2009.02013.x, 2010.

24 Chiti, T., Grieco, E., Perugini, L., Rey, A. and Valentini, R.: Effect of the replacement of
25 tropical forests with tree plantations on soil organic carbon levels in the Jomoro district,
26 Ghana, *Plant Soil*, 375(1-2), 47–59, doi:10.1007/s11104-013-1928-1, 2014.

27 Ciais, P., Sabine, C., Bala, G., Bopp, L., Brovkin, V., Canadell, J., Chhabra, A., DeFries, R.,
28 Galloway, J., Heimann, M., Jones, C., Quéré, C. L., Myneni, R. B., Piao, S. and Thornton, P.:
29 Carbon and Other Biogeochemical Cycles. In: *Climate Change 2013: The Physical Science*

1 Basis. Contribution of Working Group I to the Fifth Assessment Report of the
2 Intergovernmental Panel on Climate Change [Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor,
3 S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley (eds.)]. Cambridge
4 University Press, Cambridge, United Kingdom and New York, NY, USA, pp. 465–570, 2013.
5 doi:10.1017/CBO9781107415324.015, 2013.

6 Das, D. K. and Chaturvedi, O. P.: Root biomass and distribution of five agroforestry tree
7 species, *Agrofor. Syst.*, 74(3), 223–230, doi:10.1007/s10457-008-9159-9, 2008.

8 Davidson, E. A. and Janssens, I. A.: Temperature sensitivity of soil carbon decomposition and
9 feedbacks to climate change, *Nature*, 440(7081), 165–173, doi:10.1038/nature04514, 2006.

10 De Blécourt, M., Brumme, R., Xu, J., Corre, M. D. and Veldkamp, E.: Soil Carbon Stocks
11 Decrease following Conversion of Secondary Forests to Rubber (*Hevea brasiliensis*)
12 Plantations, *PLoS ONE*, 8(7), e69357, doi:10.1371/journal.pone.0069357, 2013.

13 Detwiler, R. P.: Land-use change and the global carbon cycle: the role of tropical soils,
14 *Biogeochemistry*, 2(1), 67–93, doi:10.1007/BF02186966, 1986.

15 Dhyani, S. K., Handa, A. K. and Uma,: Area under agroforestry in India: An assessment for
16 present status and future perspective, *Indian J. of Agroforestry*, 15(1), 1-11, 2013.

17 Don, A., Schumacher, J. and Freibauer, A.: Impact of tropical land-use change on soil organic
18 carbon stocks - a meta-analysis: Soil organic carbon and land-use change, *Glob. Change Biol.*,
19 17(4), 1658–1670, doi:10.1111/j.1365-2486.2010.02336.x, 2011.

20 Dourte, D., Shukla, S., Singh, P. and Haman, D.: Rainfall intensity-duration-frequency
21 relationships for Andhra Pradesh, India: changing rainfall patterns and implications for runoff
22 and groundwater recharge, *Journal of Hydrologic Engineering*, 18(3), 324-330,
23 doi/abs/10.1061/(ASCE)HE.1943-5584.0000625, 2012.

24 Gonzalez, J. M. and Laird, D. A.: Carbon Sequestration in Clay Mineral Fractions from C-
25 Labeled Plant Residues, *Soil Science Society of America Journal*, 67(6), 1715-1720,
26 doi:10.2136/sssaj2003.1715, 2003.

27 Hevia, G. G., Buschiazzi, D. E., Hepper, E. N., Urioste, A. M. and Antón, E. L.: Organic matter
28 in size fractions of soils of the semiarid Argentina. Effects of climate, soil texture and
29 management, *Geoderma*, 116(3-4), 265–277, doi:10.1016/S0016-7061(03)00104-6, 2003.

1 Hijmans, R. J., Cameron, S. E., Parra, J.L., Jones, P.G. and Jarvis, A.: Very high resolution
2 interpolated climate surfaces for global land areas, *International Journal of Climatology*, 25,
3 1965-1978, 2005.

4 ICAR.: Handbook of agriculture fourth edition, Indian Council of Agricultural Research
5 publishing, New Delhi, ISBN No.: 81-7164-050-8, 1984.

6 Jobbágy, E. G. and Jackson, R. B.: The vertical distribution of soil organic carbon and its
7 relation to climate and vegetation. *Ecological applications*, 10(2), 423-436,
8 doi.org/10.1890/1051-0761, 2000.

9 Kumar, B. M., George, S. J. and Chinnamani, S.: Diversity, structure and standing stock of
10 wood in the homegardens of Kerala in peninsular India, *Agrofor. Syst.*, 25(3), 243–262, 1994.

11 Lal, R.: Soil erosion and the global carbon budget, *Environ. Int.*, 29(4), 437–450,
12 doi:10.1016/S0160-4120(02)00192-7, 2003.

13 Lal, R.: Soil carbon sequestration to mitigate climate change, *Geoderma*, 123(1-2), 1–22,
14 doi:10.1016/j.geoderma.2004.01.032, 2004.

15 Lal, R.: Soil carbon sequestration in natural and managed tropical forest ecosystems, in:
16 *Environmental services of agroforestry systems*, Montanini, F. (Eds.), International Book
17 Distributing Co., Lucknow, India, 1-30, 2007.

18 Lebrete, M., Nys, C. and Forgeard, F.: Litter production in an Atlantic beech (*Fagus sylvatica* L.)
19 time sequence. *Annals of forest science*, 58(7), 755-768, 2001.

20 Lemma, B., Kleja, D. B., Nilsson, I. and Olsson, M.: Soil carbon sequestration under different
21 exotic tree species in the southwestern highlands of Ethiopia, *Geoderma*, 136(3-4), 886–898,
22 doi:10.1016/j.geoderma.2006.06.008, 2006.

23 Liu, Z., Shao, M. and Wang, Y.: Effect of environmental factors on regional soil organic
24 carbon stocks across the Loess Plateau region, China, *Agric. Ecosyst. Environ.*, 142(3-4), 184–
25 194, doi:10.1016/j.agee.2011.05.002, 2011.

26 Mangalassery, S., Sjögersten, S., Sparkes, D. L., Sturrock, C. J. and Mooney, S. J.: The effect of
27 soil aggregate size on pore structure and its consequence on emission of greenhouse gases,
28 *Soil Tillage Res.*, 132, 39–46, doi:10.1016/j.still.2013.05.003, 2013.

1 Montagnini, F. and Nair, P. K. R.: Carbon sequestration: an underexploited environmental
2 benefit of agroforestry systems, *Agrofor. Syst.*, 61(1-3), 281–295, 2004.

3 Motavalli, P. P., Palm, C. A., Parton, W. J., Elliott, E. T. and Frey, S. D.: Soil pH and organic C
4 dynamics in tropical forest soils: evidence from laboratory and simulation studies, *Soil*
5 *Biology and Biochemistry*, 27(12), 1589-1599, 1995.

6 Mutuo, P. K., Cadisch, G., Albrecht, A., Palm, C. A. and Verchot, L.: Potential of agroforestry
7 for carbon sequestration and mitigation of greenhouse gas emissions from soils in the
8 tropics, *Nutr. Cycl. Agroecosystems*, 71(1), 43–54, doi:10.1007/s10705-004-5285-6, 2005.

9 Nair, P. K. R., Kumar, B. M. and Nair, V. D.: Agroforestry as a strategy for carbon
10 sequestration, *J. Plant Nutr. Soil Sci.*, 172(1), 10–23, doi:10.1002/jpln.200800030, 2009.

11 Poeplau, C. and Don, A.: Sensitivity of soil organic carbon stocks and fractions to different
12 land-use changes across Europe, *Geoderma*, 192, 189-201, 2013.

13 Powers, J. S. and Schlesinger, W. H.: Relationships among soil carbon distributions and
14 biophysical factors at nested spatial scales in rain forests of northeastern Costa Rica,
15 *Geoderma*, 109(3), 165–190, doi:10.1016/S0016-7061(01)00119-7, 2002.

16 Powers, J. S., Corre, M. D., Twine, T. E. and Veldkamp, E.: Geographic bias of field
17 observations of soil carbon stocks with tropical land-use changes precludes spatial
18 extrapolation, *Proc. Natl. Acad. Sci.*, 108(15), 6318–6322, doi:10.1073/pnas.1016774108,
19 2011.

20 Raich, J. W. and Schlesinger, W. H.: The global carbon dioxide flux in soil respiration and its
21 relationship to vegetation and climate, *Tellus*, 44, 81-99, 1992.

22 Raich, J. W., Russell, A. E., Kitayama, K., Parton, W. J. and Vitousek, P. M.: Temperature
23 influences carbon accumulation in moist tropical forests, *Ecology*, 87(1), 76–87, 2006.

24 R Development Core Team: A language and environment for statistical computing, R
25 Foundation for statistical computing, Vienna, Austria, URL <http://www.R-project.org/>, 2014.

26 Saha, S. K., Ramachandran Nair, P. K., Nair, V. D. and Mohan Kumar, B.: Carbon storage in
27 relation to soil size-fractions under tropical tree-based land-use systems, *Plant Soil*, 328(1-2),
28 433–446, doi:10.1007/s11104-009-0123-x, 2010.

1 Salako, F. K., Babalola, O., Hauser, S. and Kang, B. T.: Soil macroaggregate stability under
2 different fallow management systems and cropping intensities in southwestern Nigeria,
3 *Geoderma*, 91(1), 103–123, doi:10.1371/journal.pone.0121368, 1999.

4 Sekhar, C., Selvarajan, M., Pounraj, A., Prahadeeswaran, M. and Vidhyavathi, A.: Production
5 and export of Mango in India - A paradigm to the developing nations, *Am. Int. J. Res.*
6 *Humanit. Arts Soc. Sci.*, 4(1), 78–87, 2013.

7 Six, J., Conant, R. T., Paul, E. A. and Paustian, K.: Stabilization mechanisms of soil organic
8 matter: implications for C-saturation of soils, *Plant and soil*, 241(2), 155-176,
9 10.1023/A:1016125726789, 2002.

10 Six, J., Bossuyt, H., Degryze, S. and Denef, K.: A history of research on the link between
11 (micro) aggregates, soil biota, and soil organic matter dynamics, *Soil and Tillage Research*,
12 79(1), 7-31, 2004.

13 Sollins, P., Homann, P. and Caldwell, B. A.: Stabilization and destabilization of soil organic
14 matter: mechanisms and controls, *Geoderma*, 74, 65– 105, doi:10.1016/S0016-
15 7061(96)00036-5, 1996.

16 Thakur, M. P., Milcu, A., Manning, P., Niklaus, P. A., Roscher, C., Power, S., Reich, P. B.,
17 Scheu, S., Tilman, D., Ai, F, Guo, H., Ji, R., Pierce, S., Ramirez, N. G., Richter, A. N., Steinauer,
18 K., Strecker, T., Vogel, A. and Eisenhauer, N.: Plant diversity drives soil microbial biomass
19 carbon in grasslands irrespective of global environmental change factors, *Global change*
20 *biology*, doi:10.1111/gcb.13011, 2015.

21 Tilman, D., Reich, P. B. and Isbell, F.: Biodiversity impacts ecosystem productivity as much as
22 resources, disturbance, or herbivory, *Proc. Natl. Acad. Sci.*, 109(26), 10394–10397,
23 doi:10.1073/pnas.1208240109, 2012.

24 van Straaten, O., Corre, M. D., Wolf, K., Tchienkoua, M., Cuellar, E., Matthews, R. and
25 Veldkamp, E.: Conversion of lowland tropical forests to tree cash-crop plantations loses up
26 to half of stored soil organic carbon, *Proc. Natl. Acad. Sci.*, 112(32), 9956-9960 , 2015.

27 Veldkamp, E.: Organic carbon turnover in three tropical soils under pasture after
28 deforestation, *Soil Sci. Soc. Am. J.*, 58(1), 175–180, 1994.

- 1 Yang, X. M. and Kay, B. D.: Impacts of tillage practices on total, loose-and occluded-
2 particulate, and humified organic carbon fractions in soils within a field in southern Ontario,
3 Can. J. Soil Sci., 81(2), 149–156, doi:10.4141/S00-015, 2001.
- 4 Zinn, Y. L., Lal, R., Bigham, J. M. and Resck, D. V. S.: Edaphic controls on soil organic carbon
5 retention in the Brazilian cerrado: soil structure, Soil Sci. Soc. Am. J., 71, 1215-1224, 2007a.
- 6 Zinn, Y. L., Lal, R., Bigham, J. M. and Resck, D. V. S.: Edaphic controls on soil organic carbon
7 retention in the Brazilian cerrado: texture and mineralogy, Soil Sci. Soc. Am. J., 71, 1204–
8 1214, 2007b.
- 9

1 **Tables**

2

3 Table 1. Environmental variables and structural characteristics of agroforestry systems in the study area (mean \pm SE).

4

AFS	Climatic zone *	Tree / palm species	Soil classification [†]	Elevation (m asl)	MAP [‡] (mm)	MAT ^c (°C)	Clay content [¶] (%)	Tree density [¶] (tree/ha)
Homegarden	Humid	<i>Artocarpus spp.</i> , <i>Mangifera indica</i> ,	Nitisols (n= 4) Acrisols (n=3)	90 \pm 17	3422 \pm 277	27.2 \pm 0.15	20.8 \pm 1.4	362 \pm 21
	Moist sub-humid	<i>Myristica spp.</i> and mixed species	Nitisols (n=6) Luvisols (n=1)	940 \pm 17	1744 \pm 52	22.1 \pm 0.15	22.6 \pm 1.2	328 \pm 17
Coffee	Humid	<i>Artocarpus spp.</i>	Nitisols (n=4) Acrisols (n=2), Luvisols (n=1)	956 \pm 42	2718 \pm 90	21.9 \pm 0.15	23.1 \pm 1.2	291 \pm 14
	Moist sub-humid	<i>Terminalia spp.</i> and mixed forest species	Nitisols (n=6) Luvisols (n=1)	981 \pm 36	1666 \pm 61	22.0 \pm 0.19	22.1 \pm 1.0	285 \pm 14
Coconut	Humid	<i>Cocos nucifera</i>	Nitisols (n=4) Acrisols (n=3)	90 \pm 17	3422 \pm 277	27.2 \pm 0.15	21.4 \pm 1.3	166 \pm 6
	Dry sub-humid		Luvisols (n=4) Nitisols (n=3)	860 \pm 20	688 \pm 49	23.5 \pm 0.15	19.4 \pm 1.4	177 \pm 6
Mango	Dry sub-humid	<i>Mangifera indica</i>	Luvisols (n=6) Nitisols (n=1)	875 \pm 29	703 \pm 29	23.3 \pm 0.19	17.9 \pm 1.3	104 \pm 6
	Semi-arid		Luvisols (n=5) Nitisols (n=2)	844 \pm 34	627 \pm 28	24.6 \pm 0.26	16.9 \pm 1.3	116 \pm 6

5 * ICAR, 1984

6 [†] FAO world reference base soil classification derived from Harmonized World Soil Database (FAO, 2009)7 [‡] Derived from WorldClim dataset (Hijmans et. al., 2005) (n=7)8 [¶] Mean \pm standard error; data derived from respective agroforestry plots (n=14)

9

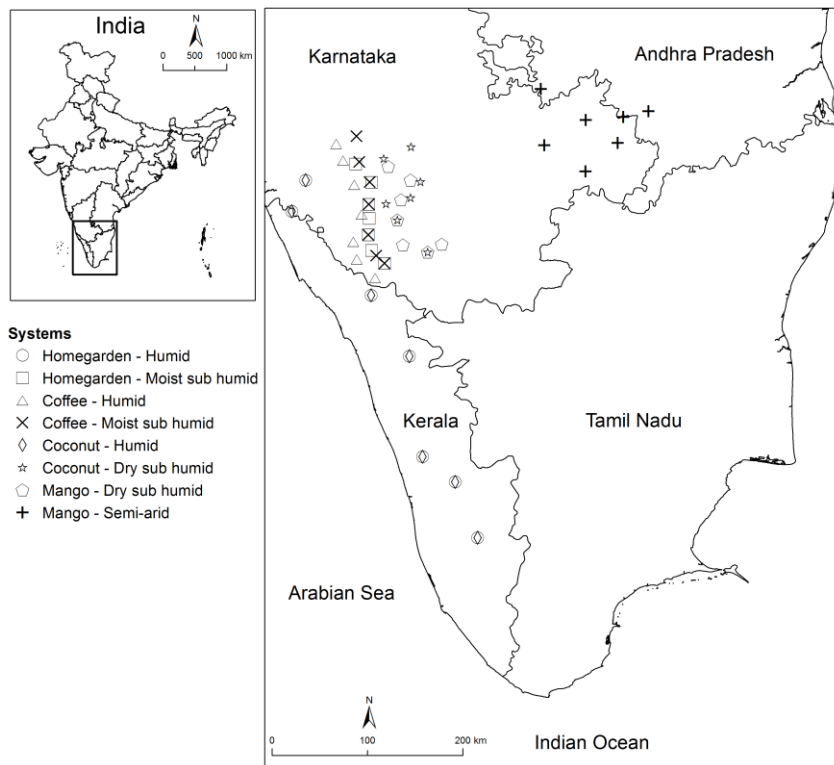
1 Table 2. Multivariate regression models predicting SOC stocks in different land-use systems and the relative changes in SOC stock (%) from either forest or
2 agriculture references in the 100 cm soil profile. MAP is mean annual precipitation (mm), MAT is mean annual temperature (°C), Clay fraction (%) and basal
3 area (m² ha⁻¹).

Land-use	Statistical model	R ²	n
<i>SOC stock (Mg C ha⁻¹)</i>			
Forest	SOC stock = 71.7 + 4.9(Basal area) – 4.3(MAT) + 1.6(Clay) + 9.0(pH) + 0.006(MAP)	0.82**	56
Homegarden	SOC stock = 95.1 + 3.5(Basal area)	0.29**	28
Coffee	SOC stock = 680.5 + 4.5(Basal area) + 4.2(Clay) – 34.7(MAT) + 0.03(MAP)	0.64**	28
Coconut	SOC stock = –186.7 + 11.5(MAT) – 1.1(pH)	0.40**	28
Mango	SOC stock = 42.3 + 4.3(Basal area)	0.34**	28
Agriculture	SOC stock = –25.1 + 4.4(Clay) + 0.006(MAP)	0.47**	56
<i>Relative change in SOC stock (%) from forest to agriculture</i>			
Agriculture	ΔSOC = -144.8 + 2.1(Clay) + 2.1(MAT)	0.34**	56
<i>Relative change in SOC stock (%) from agriculture to AFS</i>			
All AFSs	ΔSOC stock = 228.8 + 2.7(Basal area) -8.9(MAT) + 0.01(MAP)	0.15**	92
Homegarden	ΔSOC stock = 228.3 + 7.1(Clay) – 4.8(Basal area) – 37.1(pH)	0.29**	27
Coffee	ΔSOC stock = -8.7 - 7.1(Basal Area) + 7.7(Clay)	0.44	9
Coconut	ΔSOC stock = -223.5 + 6.3(Basal area) + 0.002(MAP) + 24.7(pH)	0.24†	28
Mango	ΔSOC stock = -66.7 + 13.9(Basal area) - 0.2(MAP) + 30.6(pH) - 3.0(Clay)	0.58**	28

4 † marginally significant at p ≤ 0.1, * significant at p ≤ 0.05 and ** highly significant at p ≤ 0.01

1 Figures caption

2



3

4 Figure 1. Geographical location of the plot clusters in southern India. Each point represents a
5 cluster of four plots: natural forest, agriculture and the AFSs at two ages (medium and old).

6

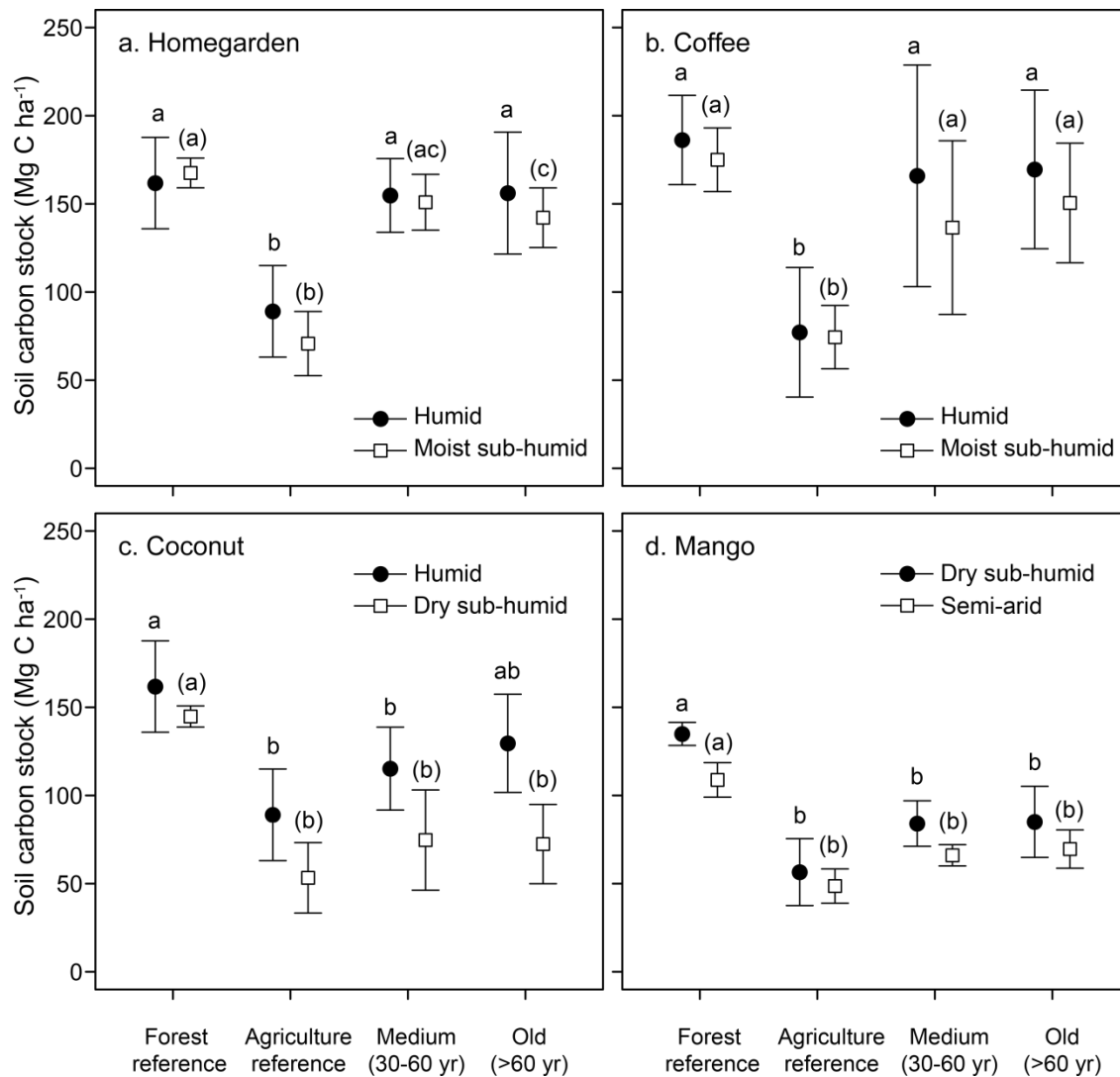


Figure 2. Absolute SOC stocks (0-100 cm) in the four different land-use systems. Each graph shows the respective forest and agriculture references and the two AFSs ages categories (medium and old) in the two different climatic zones. Point represents the mean of seven plots, while error bars indicate the 95% confidence intervals based on the Student's T distribution. Within each graph the letters indicate significant differences (P<0.05; one-way ANOVA) between land-use systems for each climate zone.

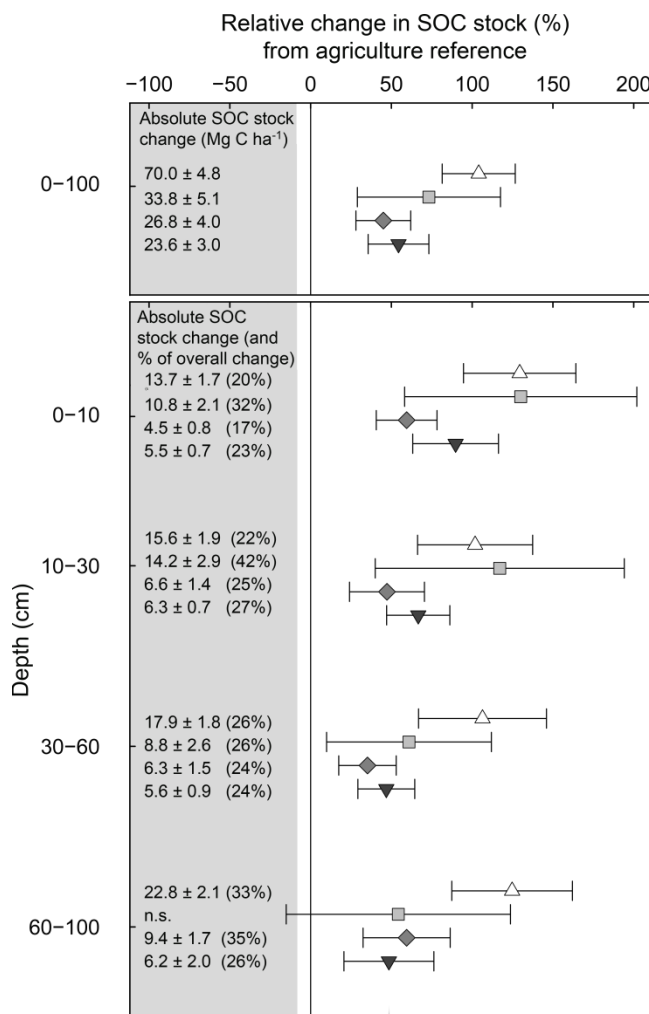


Figure 3. Relative change in SOC stock in the 1-m soil profile from agriculture to agroforestry systems (homegarden (Δ ; $n=27$), coffee (\square ; $n=9$), coconut (\blacklozenge ; $n=28$) and mango (\blacktriangledown ; $n=28$). Error bars indicate the 95% confidence intervals based on the Student's T distribution. The numbers in the grey shaded area show the absolute changes in SOC stocks (Mg C ha^{-1}) (n.s. = not significant). The numbers in brackets indicate the corresponding carbon change in percent of the overall change in the soil profile.

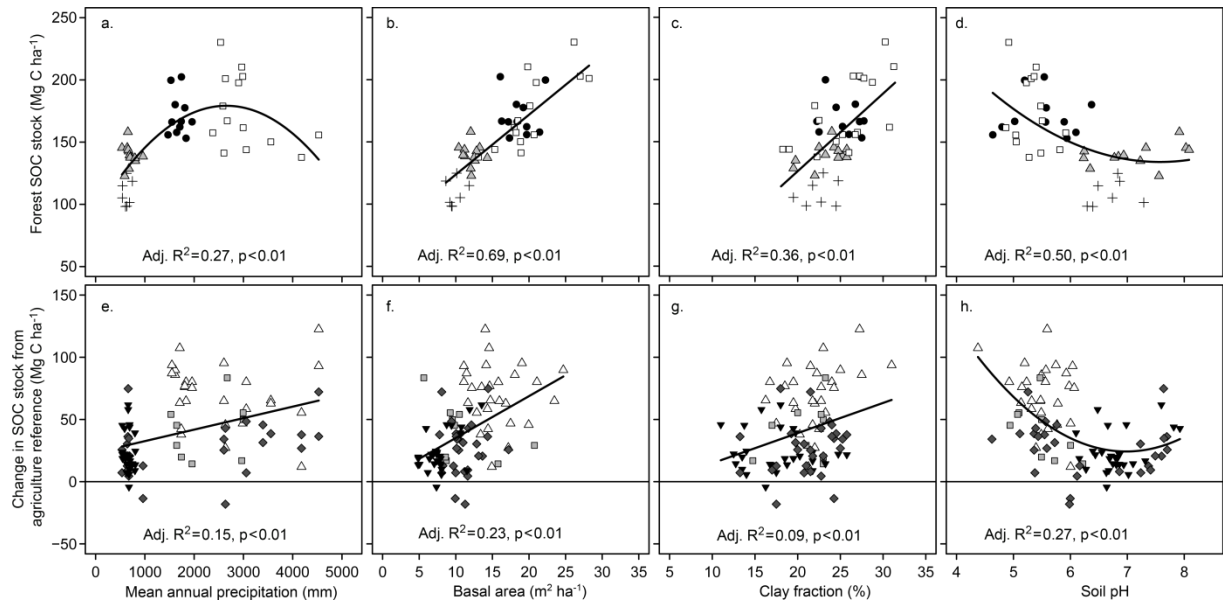
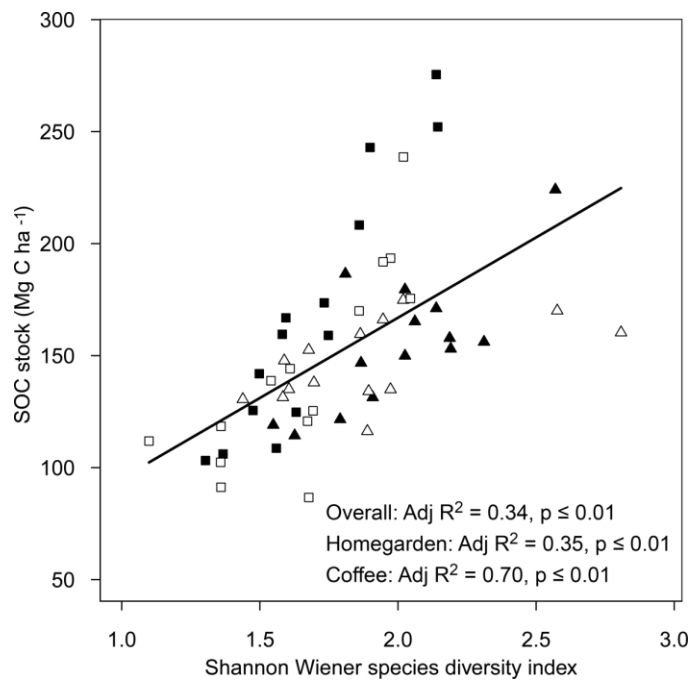


Figure 4. Scatterplots (a-d) showing the relationship forest SOC stock (top 1m) exhibits with MAP (mm), basal area (m² ha⁻¹), soil texture (clay percent), and soil pH across humid (□), moist-sub-humid (●), dry sub-humid (△) and semi-arid (+) climate zones. Scatterplots (e-h) show the relationship SOC stock changes (from agriculture reference) have with the same soil and biophysical variables in homegarden (△), coffee (□), coconut (◆) and mango (▼) AFSs.



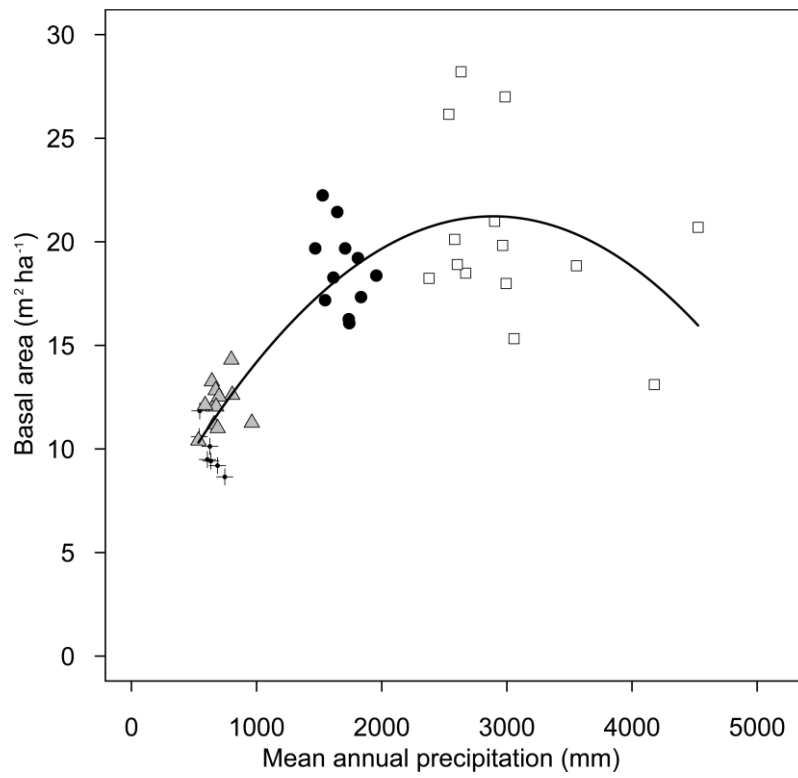
1
2 Figure 5. Scatterplot showing the relationship SOC stock of homegarden and coffee AFSs
3 exhibits with plot Shannon Wiener species diversity index: homegarden in the humid zone
4 (▲) and moist sub-humid zone (△) and coffee in the humid zone (■) and moist sub-humid
5 zone (□).

1 Appendix A



2

3 Figure A1. Photos of the six land use types investigated in two climatic zones.



1

2 Figure A2. Scatterplot showing the relationship between basal area ($\text{m}^2 \text{ha}^{-1}$) and MAP (mm)
 3 across humid (□), moist-sub-humid (●), dry sub-humid (△) and semi-arid (+) climate zones.

4

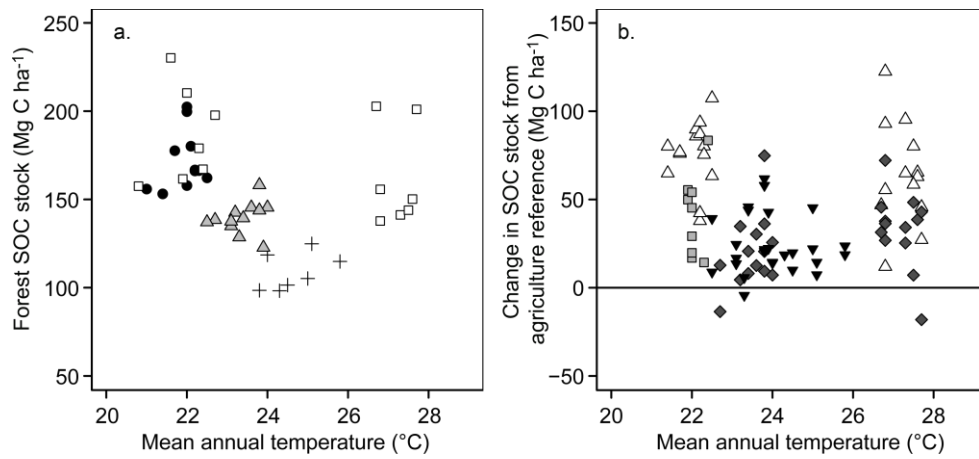
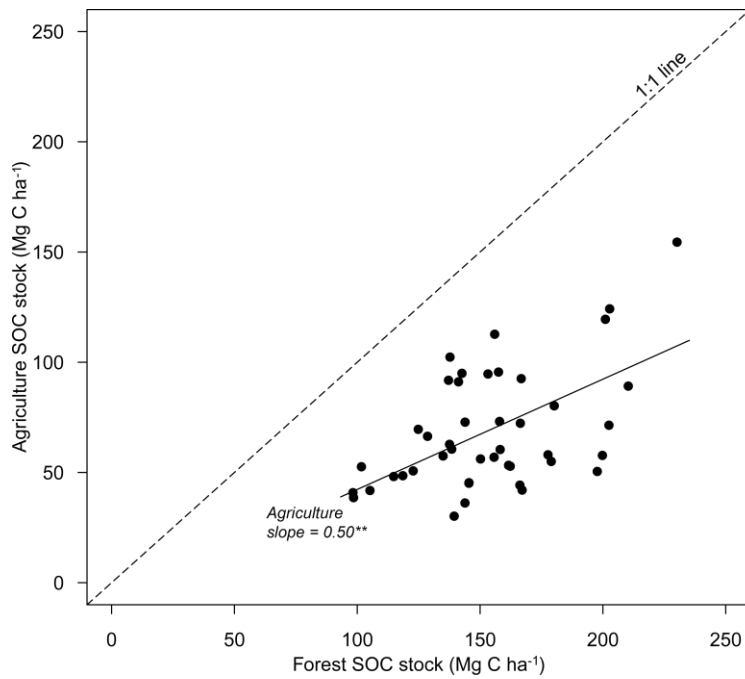


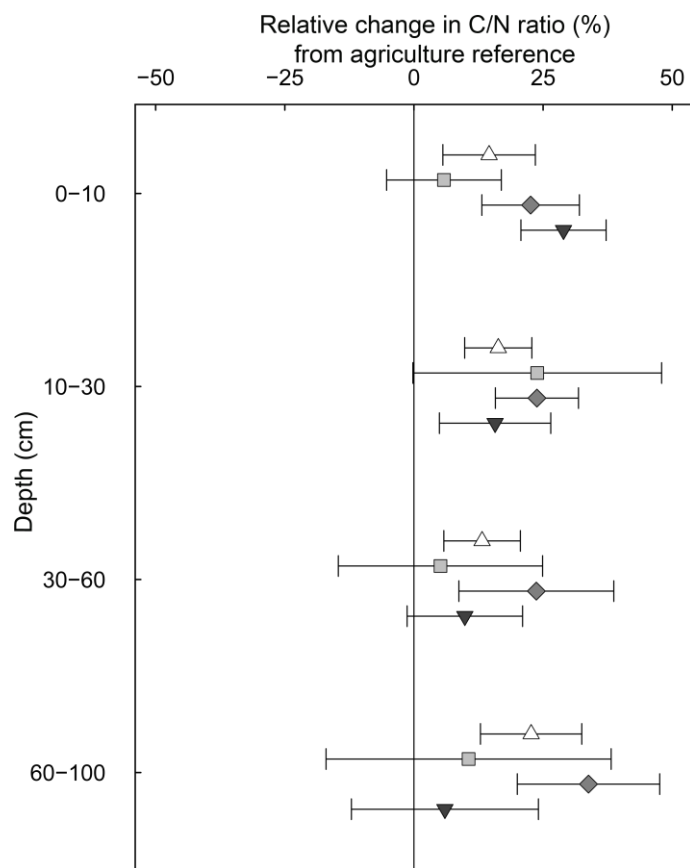
Figure A3. Scatterplots showing the relationship (or lack thereof) between MAT (°C) and (a) forest SOC stocks across humid (□), moist-sub-humid (●), dry sub-humid (△) and semi-arid (+) climate zones and (b) changes in SOC stocks in in homegardens (△), coffee (□), coconut (◆) and mango (▼) AFS.



1

2 Figure A4. a) Agriculture SOC stocks (●) in comparison to forest reference SOC stock in the
 3 100 cm soil profile. The slope of the linear regression (m) that differed significantly from one
 4 highlight an uneven response of carbon loss or gain as affected by the initial SOC stocks. The
 5 stars, ** indicates that the linear regression slope is significantly different from one (P
 6 <0.01).

7



1
2 Figure A5. Relative change in C/N ratio in the 1-m soil profile from agriculture to agroforestry
3 systems (homegarden (Δ; n=27), coffee (□; n=9) coconut (◆; n=28) and mango (▼; n=28).
4 Error bars indicate the 95% confidence intervals based on the Student's T distribution.