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# Soil surface roughness: comparing old and new measuring methods and application in a soil erosion model

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# Abstract

Quantification of soil roughness, i.e. the irregularities of the soil surface due to soil texture, aggregates, rock fragments and land management, is important as it affects surface storage, infiltration, overland flow and ultimately sediment detachment and erosion. Roughness has been measured in the field using both contact methods, such as roller chain and pinboard, and sensor methods, such as stereophotogrammetry and terrestrial laser scanning (TLS). A novel depth sensing technique, originating in the gaming industry, has recently become available for earth sciences; the Xtion Pro method. Roughness data obtained using various methods are assumed to be similar; this assumption is tested in this study by comparing five different methods to measure roughness in the field on 1 m<sup>2</sup> agricultural plots with different management (ploughing, harrowing, forest and direct seeding on stubble) in southern Norway. Subsequently, the values were used as input for the LISEM soil erosion model to test their effect on the simulated hydrograph on catchment scale. Results show that statistically significant differences between the methods were obtained only for the fields with direct drilling on stubble; for the other land management types the methods were in agreement. The spatial resolution of the contact methods was much lower than for the sensor methods (10 000 versus at least 57 000 points per m<sup>2</sup> respectively). In terms of costs and ease of handling in the field, the Xtion Pro method is promising. Results from the LISEM model indicate that especially the roller chain underestimated the RR values and the model thereby calculated less surface runoff than measured. In conclusion: the choice of measurement method for roughness data matters and depends on the required accuracy, resolution, mobility in the field and available budget. It is recommended to use only one method within one study.

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## 1 Introduction

Soil surface roughness is defined as the irregularities of the soil surface, caused by factors such as soil texture, aggregate size, rock fragments, vegetation cover and land management. Soil roughness affects surface storage, infiltration, overland flow and ultimately sediment detachment and erosion (Amoah et al., 2013). As soil roughness is usually too small to be captured by conventional topographic mapping or DEMs, it is also referred to as soil micro-topography and can be subdivided into four different categories (Takken, 2000): (i) micro-relief variations or grain roughness, mainly determined by the soil type, (ii) random roughness, related to soil aggregates, (iii) oriented roughness, describing the systematic variations in topography due to e.g. tillage, and (iv) higher order roughness representing elevation variations in the field, such as embankments or field borders. In this study, we focus on the second type, random roughness.

Since the late 1950s (Zobeck and Onstad, 1987), soil surface roughness has been quantified using various indices, mostly based on statistical indicators of surface elevation point measurements (Doren and Linden, 1986; Lehrsch et al., 1988). One of the most commonly used soil roughness indices is the Random Roughness defined by Allmaras et al. (1966) as the natural logarithm of the standard deviation of multiple height measurements after removing the effects of slope and oriented roughness and the 10 % upper and lower extreme values. Cremers et al. (1996) showed that the standard deviation of height measurements after removing slope effects could be used equally well. This latter definition of random roughness (RR) is used in this study.

The RR index was found to vary significantly with changes in land use, tillage orientation and over time (Cremers et al., 1996). The index is used in several equations to estimate surface depression storage, i.e. the fraction of surface covered with water and amount of excess rainfall needed before runoff starts, for instance in physically-based hydrological or soil erosion models. In such models, roughness is either included explicitly by performing hydraulic calculation at high spatio-temporal resolutions (Esteves

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et al., 2000; Fiedler and Ramirez, 2000), but more commonly by using a roughness parameter as proxy value in the runoff process in the simulation. Arguably, one of the reasons that erosion modelling is still not satisfactory (e.g. Takken et al., 1999; de Vente et al., 2013) is the lack of input data that capture the heterogeneity of the area of interest (Kværnø, 2011; Sheikh et al., 2010). Therefore, one way to improve erosion model performance is to improve the accuracy and precision of the model input data such as random roughness, using new sensor measurement methods and to test how sensitive the model is to variations in input data.

Soil roughness has been measured in the field using various methods (García Moreno et al., 2010), some of which capture a profile or transect of elevation data, while others cover an area, typically extending a few square meters. These methods differ in obtainable accuracy and resolution, but also in practical aspects, such as costs and workability in the field. Jester and Klik (2005) compared measured roughness data from pin-profilers and a roller chain as contact methods and from laser scanner and stereophotogrammetry as sensor methods. They did not compare between the contact and non-contact methods and also not consider a specific purpose for which the resulting roughness data would be used.

The current study extends the study by Jester and Klik (2005), by directly comparing roughness values measured using five different methods on different agricultural surfaces. The measurement methods evaluated were the contact methods pinboard and roller chain and the sensor methods TLS, stereophotogrammetry and the new Xtion Pro method (ASUS, 2013). This latter depth sensing technology originates in the gaming industry and has recently been used in earth science studies (Mankoff and Russo, 2012). For example, Hammerle et al. (2014) used it to model a karst cave, Tortini et al. (2014) used it to assess volcano deformation and Brouwer (2013) used it to assess tree metrics. The advantage over state-of-the-art terrestrial laser scanners is the increased mobility, especially relevant for erosion monitoring at the landscape scale, and the relatively low costs (James and Quinton, 2014; Mankoff and Russo, 2012). Next to direct comparison of the measured roughness values from the five methods, they

were also compared in terms of their effect on the output of the distributed, process based soil erosion model LISEM (Limburg Soil Erosion Model, De Roo et al., 1996). The objectives of the study were to obtain a quantitative comparison of the different measurement methods and to quantify the propagation of the differences in an erosion model.

## 2 Methodology

### 2.1 Study area

This study was conducted in the Skuterud catchment (4.5 km<sup>2</sup>), located east of Ås, approximately 30 km south of Oslo, Norway (Fig. 1). The landscape is undulating with elevations ranging from 92 to 150 m a.s.l. Slopes are mainly between 2–10 %, with the steepest slopes in the eastern and western part of the catchment. Annual average temperature is 5.3 °C and annual average precipitation is 785 mm. Sixty percent of the catchment area is cultivated land, while about thirty percent is forest, mainly coniferous or mixed coniferous-deciduous, 7 % is urban settlement and roads and 2 % is covered by forested peatland located in the southernmost depression in the catchment. Geologically, the area mainly consists of marine deposits. Coarse marine shore deposits dominate on the fringe of the cultivated areas and in the forest. The dominant soil types in the central part are marine silt loam and silty clay loam soils (Albeluvisols and Stagnosols). In the areas with shore deposits sandy and loamy sand soils dominate (Arenosols, Umbrisols, Podzols, Cambisols and Gleysols). Loamy and sandy loam soils are found in the transition zones between marine and shore deposits. Monitoring of water and sediment discharge has been carried out since 1993 at the catchment outlet (northernmost point in the catchment). In 2008, continuous monitoring started at the outlet of the Gryteland sub-catchment (Fig. 1). The sub-catchment covers approximately 0.27 km<sup>2</sup> and is dominated by relatively more sandy soils and forest in comparison with the main catchment (Kværnø and Stolte, 2012). Arable land in the catchment

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is mostly used for the production of cereals, i.e. winter and spring wheat, barley and rye. If not left under stubble over winter, soils are tined, cultivated, mouldboard ploughed and/or harrowed in autumn. Some of the surface roughness measurements were conducted in another catchment near Leirsund, Lillestrøm, 40 km north of Skuterud. This catchment was characterized by relatively more flat slopes and a higher clay content than in Skuterud.

## 2.2 Data collection

Soil roughness data was collected on four typical land units in the catchment in the period between March and June 2013. Two tilled surfaces: harrowed and ploughed, and two untilled surfaces: direct drilling on stubble and forest were analyzed. In addition, three texture classes were defined initially to discriminate between land units, but preliminary data analysis showed no significant differences between the harrowed surfaces on two different texture classes. Therefore, no further differentiation in land units was made based on texture. On each land unit at least three random locations were selected. As the field work was conducted in late spring, when seedbed preparation had almost been completed throughout the catchment, there were hardly any mouldboard ploughed fields left, so part of the measurements on ploughed surface had to be done in a similar catchment north of Skuterud. On each location a plot of 1 m × 1 m was marked on which soil roughness measurements were taken (Fig. 2).

Soil roughness was measured using five different methods (Table 1; Fig. 3): roller chain, pinboard, stereophotos, TLS and the Xtion Pro. Four of these methods have been used before (Aguilar et al., 2009; Allmaras et al., 1966; Barneveld et al., 2013; Haubrock et al., 2009; Jester and Klik, 2005; Saleh, 1993), while the Xtion Pro is a novel technique (Mankoff and Russo, 2012). Two methods were contact methods (roller chain and pinboard) and three were non-contact, sensor methods (stereophotos, TLS and the Xtion Pro). For the sensor methods, each technique was applied on an area of 1 m<sup>2</sup>. For the pinboard method, two transects were recorded along the direction of cultivation (e.g. ploughing) on a 1 m<sup>2</sup> area, while for the roller chain, 4 replicates along the direction



of cultivation on the 1 m<sup>2</sup> area were recorded. Recordings were taken only along the tillage direction, to avoid having to correct for the systematic roughness caused by the tillage. When roughness was isotropic (forest and direct drilling on stubble), one or two replicates were taken in both directions on the plot.

The roller chain (Fig. 3a) used in this study had joints of 4 mm and links of 6 mm and a total length of 1 m. The chain was carefully laid out on the surface and the length the chain covered was measured with a ruler. The ratio between the distance-over-surface (1 m) over the Euclidean distance (measured by ruler) of the chain was used to calculate the Cr index, as a measure for the roughness. The Cr index is calculated using Eq. (1) (Saleh, 1993):

$$Cr = \left(1 - \frac{L2}{L1}\right) \times 100, \quad (1)$$

where  $L1$  is the distance-over-surface [m] (here 1) and  $L2$  is the Euclidian distance [m]. The pinboard (Fig. 3b) used in this study was a wooden frame of 1 m width with 50 metal pins, each 3 mm in diameter, with 18.5 mm spacing between pins. While the pinboard was placed on the surface and a white screen was placed behind, a digital camera recorded an image from which the height of each individual pin later was estimated using the software program ImageJ (Rasband, 2013). The standard deviation of all recorded pin heights is a measure for the roughness.

Terrestrial laser scanners (TLS) have been used increasingly during the last 20 years in soil surface survey studies to obtain 3-D data of micro-topography (e.g. Aguilar et al., 2009; Barneveld et al., 2013; Haubrock et al., 2009; Takken, 2000). TLS height data was acquired using a Leica ScanStation2 (Fig. 3c). The scanner sends out pulsed laser beams which by a time-of-flight principle estimates distances (Leica-Geosystems AG, 2007). At the sensor head the pulsed beam is 4 mm in diameter, which increases to 6 mm at 50 m distance (Barneveld et al., 2013). The signals were processed using the licensed Cyclone Software (Leica-Geosystems AG, 2014). Since a very limited area (1 m<sup>2</sup>) had to be covered a single scan was enough to obtain a high accurate

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Stereophoto images were constructed using a digital camera, Olympus PEN E-P3 with a M. Zuiko Digital ED 12 mm f2.0 lens (Fig. 3d) and the Photomodeler software (Eos systems Inc., 2011). The program links several images using the black-and-white markers (Fig. 2). In each plot several images with slightly different view angles were linked to establish the height differences. The method only works properly with angles up to 45 degrees.

In 2010 a new depth-sensing technology was developed under the name Kinect. It was originally intended for stationary indoor use in combination with a gaming console. The technology proved to be quite accurate in depth estimates and became of interest for the geophysical research (Kramer et al., 2012; Mankoff and Russo, 2012). Subsequently, similar sensors started to be developed parallel to the development of the Kinect. In this study the Xtion Pro was used (Fig. 3e). The Xtion Pro has an emitter that sends out an infrared pattern, which is reflected and recorded by the sensor on the device. This is used in a triangulation process to estimate distances (Khoshelham and Elberink, 2012). The Xtion Pro was fixed to a wooden handle to be able to take recordings close enough to the soil surface. The sensor was connected to a laptop in order to communicate with the Skanect software (Tisserand and Burrus, 2013). The Xtion Pro recording had to be done without direct sunlight, as the natural light interferes with the infrared signal.

For all sensor methods, roughness was calculated as the standard deviation of the point heights in each point cloud.

An H-flume was installed to measure overland flow. Water depth (logged at 10 min intervals) in the flume was measured using an ultrasonic sensor, and a conversion from depth to discharge was performed using a height-to-discharge relationship. Volume proportional samples from surface and drain discharge were taken during discharge episodes (Stolte and Kværnø, 2013).



## 2.3 Data processing

For each data sampling location, one roughness estimate per method was calculated. For the roller chain this was an average of four measurements and for the pinboard an average of two. As mentioned in the introduction we used the random roughness (RR) index for comparison, which is calculated based on height measurements. Thus, to be able to compare the roller chain with the other four methods a conversion of the Cr index to the RR index was necessary. From the pinboard it is possible to estimate both the RR-index and Cr index by summing up the Euclidean distance between consecutive pins. By plotting the two different indices from the same measurements in a scatter plot a regression between the two indices could be established. This was used to convert the Cr index to RR values (Jester and Klik, 2005). For the correlation please refer to Appendix A.

For each of the sensors a different software program was employed to transform the measurements to data files, however all data output was in xyz-coordinates (i.e. a point cloud) in an ASCII-file format. To process data in an efficient way, a routine was developed in R (version 3.0.1, using R studio version 0.97.511). This routine placed the coordinate system of each point cloud in the center of the cloud, converted units to meters, cut out an area of  $0.8 \times 0.8$  m to avoid side disturbance of the plot that would distort the roughness estimates, removed any slope effect by fitting a plane through the point cloud and extracted height data by calculating the distance from each point to the plane. Removing outliers defined as heights larger than three times the standard deviation of the measurements. In contrast to other studies (Jester and Klik, 2005; Taconet and Ciarletti, 2007; Vaudour et al., 2014) the RR was calculated directly from the point cloud data points instead of by defining a grid. This was done to obtain the highest possible resolution and accuracy. If a grid was defined, the sub-grid microtopography would have been averaged out because of the large minimum cell size required in order to full point coverage of the plot.

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## 2.4 Comparison of measurement techniques

RR results were compared between the different measuring techniques according to four aspects: accuracy, precision, data resolution and the price of the different devices. Accuracy, defined as how well a method reproduces the soil surface, was assessed using the TLS data as reference, assuming that TLS data is the most accurate measurement device in outdoor environments (Aguilar et al., 2009; James and Quinton, 2014). Precision, defined as a measure for the data reproducibility of each method was assessed using the standard deviation of multiple measurements on the same plot. Further, an ANOVA test was used to assess how well the methods agreed about an RR value for a land unit. Spatial resolution was defined as number of points per square meter or per meter length (for the 2-D methods). Finally, the price of acquiring or using the different methods was compared.

## 2.5 Random roughness in the LISEM erosion model

The Limburg Soil Erosion Model (LISEM) is an event-based spatially distributed model, originally developed to simulate the effect of local measures on soil loss in the province of Limburg in the Netherlands. The model incorporates, where possible, physical principles rather than empirical relations (Jetten and De Roo, 2001). LISEM simulates the surface runoff and sediment transport during and after a single rain event, using a standard surface water balance. It incorporates interception, infiltration and storage in micro-depressions before the transport capacity of the runoff is calculated. The transport capacity of the runoff together with splash erosion is used to estimate sediment detachment. The theoretical framework of the model has been described in detail in e.g. De Roo et al. (1996a) and Jetten (2002). Sensitivity analysis of the input parameters showed that the model is most sensitive to the soil saturated hydraulic conductivity, followed by RR, slope gradient and Manning's  $n$  (De Roo et al., 1996b). In the model, RR is used to calculate surface storage in micro-depressions. In this study, the erosion model was run using the different average RR values resulting from each measurement

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method and the simulated hydrographs at the sub-catchment outlet were compared. By assessing how well the hydrographs of the rain event fitted the hydrograph obtained with the RR measured with TLS, the accuracy of each method relative to the TLS was evaluated.

### 2.5.1 Surface storage in LISEM

Surface ponding of precipitation cannot directly be modeled due to lack of high resolution data of the micro-depressions. In LISEM, ponding on the surface is simulated using the concept of Maximum Depressional Storage (MDS, cm). MDS is defined as the threshold value for a given area above which the surface micro-depressions will overflow. When this value is reached in any cell, each additional raindrop will directly result in overland runoff out of the cell. MDS is related to RR (Kamphorst et al., 2000) as follows (Eq. 2):

$$\text{MDS} = 0.234\text{RR} + 0.01\text{RR}^2 + 0.012\text{RR} \times S, \quad (2)$$

where RR is random roughness [cm] and  $S$  is terrain slope [%]. As MDS is an expression for the maximum storage in any given cell it is recognized that some runoff will occur before this threshold. This is modelled using the concept of Start Depressional Storage (SDS, cm). When the SDS threshold is reached some runoff occurs from a given cell. SDS is arbitrarily defined to be reached when 10 % of a cell is ponded. The fraction of ponded surface ( $f_{\text{ps}}$ ) is estimated using Eq. (3):

$$f_{\text{ps}} = 1 - \exp\left(-\left(1.406 \times \text{RR}^{-0.942}\right) \times h\right), \quad (3)$$

where  $h$  [mm] is the water depth at the surface. Thus, when  $f_{\text{ps}}$  equals 0.1 in a given cell runoff is assumed to start. Reworking the equation with  $f_{\text{ps}} = 0.1$  gives the threshold water height (SDS) by which runoff will start given the RR of that cell:

$$\text{SDS} = \frac{\ln(0.9)}{-(1.406 \times \text{RR}^{-0.942})}. \quad (4)$$

As long as the actual water height  $h$  is smaller than SDS no runoff occurs. When the water height is larger than SDS but smaller than MDS, runoff is calculated using Eq. (5):

$$(h - \text{SDS}) \times \left( 1 - \exp\left(-h \times \frac{(h - \text{SDS})}{(\text{MDS} - \text{SDS})}\right) \right), \text{ (for } \text{SDS} < h < \text{MDS}) \quad (5)$$

5 and when  $h$  is larger than MDS runoff is calculated using Eq. (6):

$$h - \text{MDS}, \text{ (for } h > \text{MDS}). \quad (6)$$

Runoff is a function of RR and slope, after interception and infiltration have been subtracted (Jetten, 2002). The fraction of the water height which runs off will increase when RR decreases, i.e. when the soil surface becomes less rough. The smoother the surface the more water runs off. This is logical following physical principles, but the concept of SDS and MDS is a necessary approximation based on empirical observations.

## 2.5.2 Simulated rain event and model settings

A rain event from 4 September 2009 was used for the simulations, from which surface discharge was recorded with the flume at the outlet of Gryteland sub-catchment. This rain event lasted approximately 12 h, with a peak rain intensity of  $18 \text{ mm h}^{-1}$  and a total precipitation of 18.6 mm. The LISEM model has been calibrated for the Gryteland sub-catchment using a rain event from August 2010 (Kværnø and Stolte, 2012). We used the same model settings for our simulations, including a spatial resolution of 10 m and model time step of 30 s. The saturated hydraulic conductivity was re-calibrated, and was multiplied by a factor of 3 for the clay soils in the Gryteland catchment (versus a factor 4.5 for the 2010 event) and Manning's  $n$  value for arable land was set to 0.45 to match the measured hydrograph of 4 September 2009. The effect of the RR input values on the simulated outlet hydrograph was subsequently tested using the average values for each of the five measurement methods as input.



## 3 Results and discussion

### 3.1 Comparison of methods

Figure 4 shows the roughness values obtained with four different methods (the Xtion Pro, stereophotos, pinboard and roller chain) on the different land units (harrowed, ploughed, forest and direct drilling on stubble). The methods seem to be in better agreement on the harrowed and ploughed surfaces, concerning both mean and spread of data, whereas the two plots with vegetation show more varying RR values. The results of the ANOVA test indicated that there were significant differences between the RR values on plots with stubble. For each of the other surfaces the methods were in agreement according to the ANOVA test. The precision of the methods varied between the methods, but it cannot be concluded that the precision of one method was higher than another as the time constraint of the field work only allowed for few replicates and a varying number of replicates for each land unit. The limited spread of the roller chain data on forest and stubble indicates that the surface was very homogenous or that the roller chain was not able to capture differences present on these surfaces.

The vegetation residuals were not filtered out prior to the RR calculation even though RR is meant to express the soil surface roughness and not the general roughness. This might partly explain the disagreement between measurement on the forest and stubble plots, where vegetation residuals can distort the measurement. This would mainly be an issue for the non-contact methods and the chain, whereas the pinboard has the advantage that the pins can fit in between straws and still reproduce the height difference of the soil surface. For the stereophoto method an explanation for the relatively low values and spread of RR data is that the plant residuals have a too fine structure to be reconstructed resulting in an interpolated surface between some points, filtering out details. This explanation could apply to the Xtion Pro as well, but the RR value measured with the Xtion Pro on forest and stubble is significantly higher than that of the stereophoto method, indicating that the Xtion Pro does not have the same issues as the stereophoto method. For the chain, the explanation for the relatively low values and

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spread of RR data can be that the size of the joints is too large to fit in between smaller soil aggregates, resulting in a smoother surface as reproduced by the roller chain than in reality. This means that the resolution of the chain method is restricting its applicability to surfaces that have relatively large soil aggregates, as also concluded by Jester and Klik (2005). Another restriction for the chain method is that it is not measuring height differences but surface length differences (Skidmore, 1994).

Figure 5 shows the RR values for one representative plot on a harrowed field. Since roughness for each point cloud is calculated as the standard deviation of the full point cloud, only one number can be presented for the sensor methods. The numeric difference between the TLS, taken here as the most accurate method, and the other methods is smallest for the pinboard (+3 %) and Xtion Pro (−7 %), intermediate for the stereophoto (−16 %) and largest for the chain method (+35 %). Although it is not possible to base the argument on statistics, the results seems to indicate that the Xtion Pro is in better agreement with the TLS than the stereophoto and therefore more accurate. Likewise the pinboard is in better agreement with the TLS than the chain. To further assess these first results in detail and to be able to base them on statistics, more replicates are needed.

Each of the RR values acquired by non-contact methods is the standard deviation of a minimum of 57 000 height point measurements on 1 m<sup>2</sup> and thus it can be argued that it is a representative number for this land unit. While most sensors are reported with accuracy estimates, the actual accuracy of the field data are affected by environmental factors, which ideally should be known before using the method (Barneveld et al., 2013). For future studies it will be useful to make thorough accuracy tests under field conditions, for instance with objects of known dimensions placed in the field. This will especially be interesting for the Xtion Pro in order to fully assess the potential for this sensor in earth science studies.

### 3.2 Data resolution

For the roller chain the size of the joints represents the degree of detail the chain is able to measure, however the resolution is not easy to evaluate. For the pinboard the resolution was 100 points per meter or 10 000 points per m<sup>2</sup>. In contrast, each of the sensor methods had at least 57 000 points per m<sup>2</sup>. While the TLS had a resolution of 2 million points per m<sup>2</sup>, the Xtion Pro and stereophotos had fluctuating resolutions based on the physical properties of the soil surface and the environmental conditions. For example, on a ploughed plot the shading caused by steep furrows, caused gaps in the stereophoto data. The very irregular surfaces, such as the ploughed surfaces, also caused trouble for the Xtion Pro in some cases, since the software program is made for smooth surfaces and thus it seemed that the software program smoothed out deeper cavities in the surface. From the software documentation (Tisserand and Burrus, 2013) it was not possible to obtain further information regarding this.

It should be kept in mind that the various sensors use different techniques for estimating distance and that different algorithms are applied in the data processing. Although the software programs sometimes report accuracy estimates, it was unknown which approaches and algorithms are used to process the measurements of both the TLS (Leica-Geosystems AG, 2014) and the Xtion Pro (Tisserand and Burrus, 2013). Although these routines are improving the quality of the recorded data, it is not guaranteed that the data from different devices are treated consistently, which could be problematic when the data is compared. The quality of data produced by sensors will improve when the algorithms have been better documented and is an issue that requires further attention.

### 3.3 Price of devices

Table 2 provides an overview of the approximate expenses related to the various methods. The contact methods are significantly cheaper than the sensor methods. Of the sensor methods, the Xtion Pro is by far the cheapest. It should however be kept in

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mind that a digital camera can be used for many other purposes and as such, the comparison is only indicative. In addition, the Xtion Pro needs a computer with a high-resolution graphics card to be able to analyze the data. The price for the computer is not included in the comparison. Likewise, it is possible that the TLS can be rented rather than purchased and in that case becomes cheaper than indicated here.

### 3.4 Roughness values as input to an erosion model

The hydrograph output from the erosion model, based on model runs with RR input values from the different methods, is shown in Fig. 6. The RR values obtained with the Xtion Pro and stereophotos were lower than the TLS (Fig. 5) and produced hydrographs with a slightly higher peak discharge (Xtion Pro:  $43 \text{ L s}^{-1}$  and Stereophoto:  $47 \text{ L s}^{-1}$  vs. TLS:  $40 \text{ L s}^{-1}$ ). Similarly, the pinboard and the chain method produced higher roughness values which resulted in lower peak discharges, most remarkably for the chain method, which simulated the peak discharge to be only  $15 \text{ L s}^{-1}$ . This was less than half the amount of the measured discharge and the discharge simulated with the calibrated model using the TLS RR data.

These results clearly show the importance of the RR input values on the simulated hydrograph and the effect that different measurement techniques have. To ensure consistency in a field campaign, it will be crucial to use the same device. The model sensitivity to RR is determining how accurate the data need to be – this means that a standardized method and number of replicates should be stated in the model user manual and different measurement methods within the same study (area) should be avoided.

It should be remembered that this result is based on a small homogenous catchment. The effect in larger areas with more heterogonous RR values should be assessed as well.

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4 Conclusions

This study is, to our knowledge, the first that directly compared five different techniques to measure random roughness; one of which is the new method Xtion Pro which originates in the gaming industry and is here applied to measure roughness in an earth sciences context. It can be concluded that for measuring soil roughness the choice of methods matters, as proven by the statistical tests and the differences in model output in this study. Which method is the best choice depends on required accuracy and resolution, mobility in the field and available budget. In general the resolution of the contact methods (roller chain and pinboard) seems to restrict their applicability, not to mention how time-consuming and impractical they were in the field. In contrast, the Xtion Pro proved to be a useful, cheap and accurate sensor, a promising alternative to TLS and stereophotos, although more research is needed regarding accuracy and software.

The TLS and stereophoto technique each have their advantages. The digital camera gives the user most mobility but requires expensive expert software to process. The TLS ensures highest accuracy and resolution in outdoor environments but is not practical to use in extensive field measurement campaigns because of all the required equipment.

The choice of roughness measurement device is important when the data is used as input for an erosion model, as shown for the LISEM model in this study. With respect to the inconsistencies that can be expected between roughness measurement devices, it would be useful to attach information regarding the measuring device and number of necessary replicates to the model guidelines.

The sensitivity of LISEM to RR is determined by equations based on empirical relations found during field trials. Thus, the method and device should match the ones used in this empirical study (Kamphorst et al., 2000), to be in line with the values that were used to establish the equation, rather than getting the highest possible accuracy or resolution. Interestingly, in Kamphorst et al. (2000) roughness was determined with a pinboard and a laser scanner with all roughness values averaged together.

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With the sensor methods, a representative 3-D image of the microtopography could be created which potentially incorporates more information regarding surface roughness than is included in the RR index. Thus, while the high resolution sensors get cheaper and more mobile it is worth considering if overland flow should be estimated by other indices or methods than the RR index and the maximum storage capacity concept.

## Appendix A

Figure A1 shows a scatter plot of the Cr index calculated based on measurements from the roller chain against the Cr index calculated based on the Euclidean distances between consecutive pins from the pinboard. A linear regression can be established with a satisfying  $R^2$  value.

Figure A2 shows the non-linear regression fitted to the scatter plot of the Cr index from the chain against the RR measured with the pinboard. Three outliers were excluded from the regression to create a better fit.

*Acknowledgements.* The authors gratefully acknowledge the contribution of Christophe Moni (development of point cloud processing routine in R) at Bioforsk. The research was financed by the SIS-Catchy project of Bioforsk.

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**Table 1.** Land unit, number of plots and point height measurements per method.

Land unit	Measurement device	Number of readings
Harrowed	Roller chain	22 transects
Harrowed	Pinboard	10 transects × 50 points = 500 points
Harrowed	Stereophotos	7 plots, average 38 197 points
Harrowed	Xtion Pro	6 plots, average 76 099 points
Harrowed	Laser scanner	1 plot, 2 134 146 points
Ploughed	Roller chain	30 transects
Ploughed	Pinboard	12 transects × 50 points = 600 points
Ploughed	Stereophotos	5 plots, average 36 740 points
Ploughed	Xtion Pro	3 plots, average 80 027 points
Direct seeding on stubble	Roller chain	18 transects
Direct seeding on stubble	Pinboard	6 transects × 50 points = 300 points
Direct seeding on stubble	Stereophotos	3 plots, average 89 678 points
Direct seeding on stubble	Xtion Pro	3 plots, average 67 651 points
Forest	Roller chain	18 transects
Forest	Pinboard	6 transects × 50 points = 300 points
Forest	Stereophotos	3 plots, average 87 541 points
Forest	Xtion Pro	3 plots, average 58 794 points
Forest	Laser scanner	1 plots, 2 193 589 points

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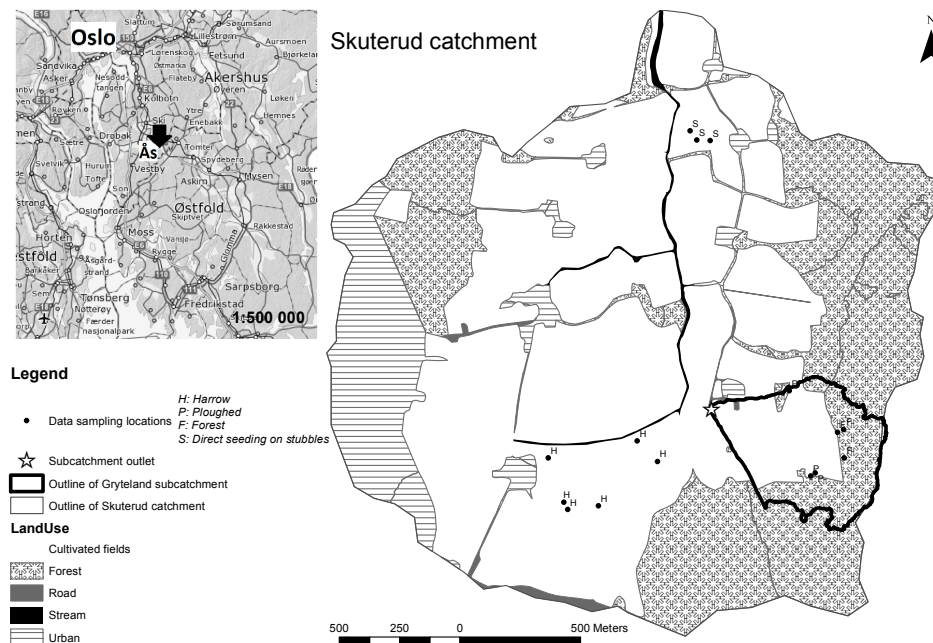
**Table 2.** Approximate prices related to roughness measurement devices. All prices are in euro. Note that the prices might be subject to change between countries around the world, temporary offers and prices fluctuations after this study was done.

Technique	Price of device	Price of used software	Total
Roller chain	74	–	74
Pinboard (home made)	30	freeware (ImageJ)	30
Stereophotogrammetry	1000	2145	3145
Leica Terrestrial Laser scanner	60 000	(price includes software)	60 000
Xtion Pro	100	100 (Skanect)	200



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**Figure 1.** Outline and land use of Skuterud catchment. The sub-catchment Gryteland is outlined in black. Data sampling locations are indicated. The inserted map in upper left corner indicates the location of the catchment. Land use map obtained from Bioforsk. Norway map obtained from: <http://www.skogoglandskap.no/kart/kilden>.

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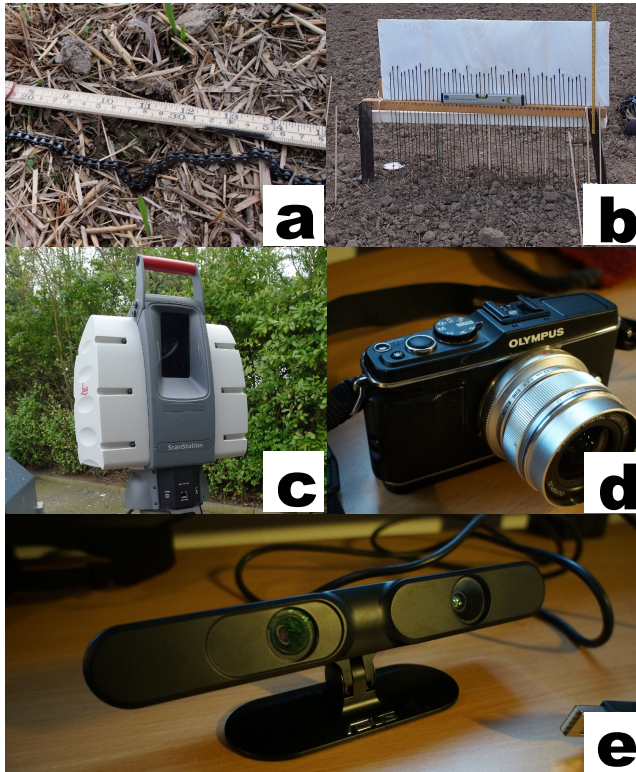
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**Figure 2.** Setup of the measurement plot. The markers delineate an area of  $1\text{ m}^2$ . The patterns on the markers were used for the stereophoto technique.



**Figure 3.** Five devices used for measuring soil surface roughness, **(a)** roller chain; **(b)** pin-board; **(c)** Leica ScanStation2 terrestrial laser scanner; **(d)** the digital camera used for creating stereophotos; and **(e)** the Xtion Pro. **(a)** and **(b)** are contact-methods, while **(c)**, **(d)** and **(e)** are non-contact methods.

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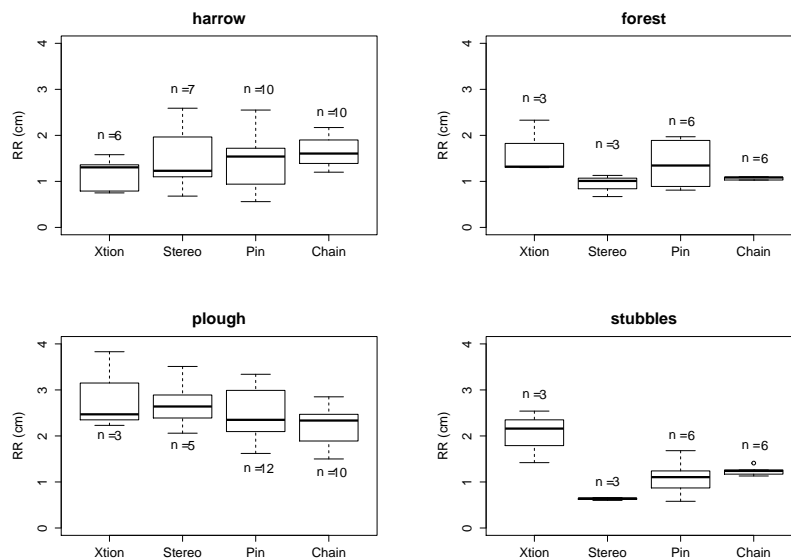
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**Figure 4.** Boxplot of RR values measured with four different methods on four different land units.

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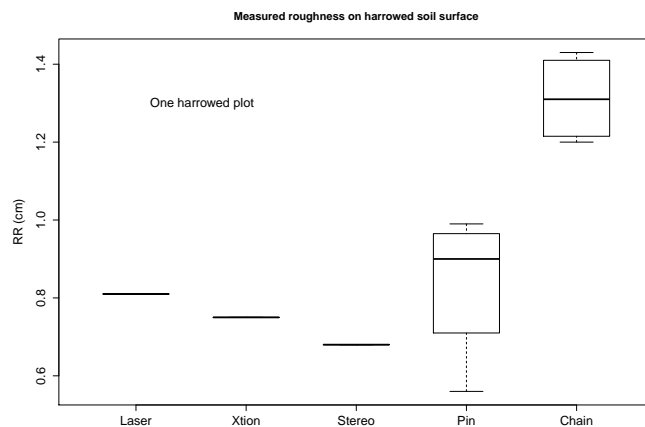
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**Figure 5.** Random roughness (RR) values for one harrowed plot based on five different measurement methods.

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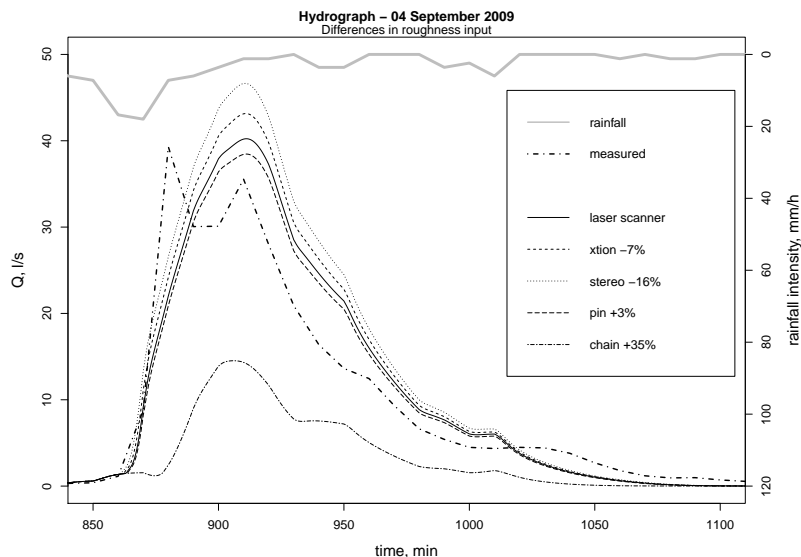
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**Figure 6.** Model output as a function of mean RR values per method for harrowed fields.

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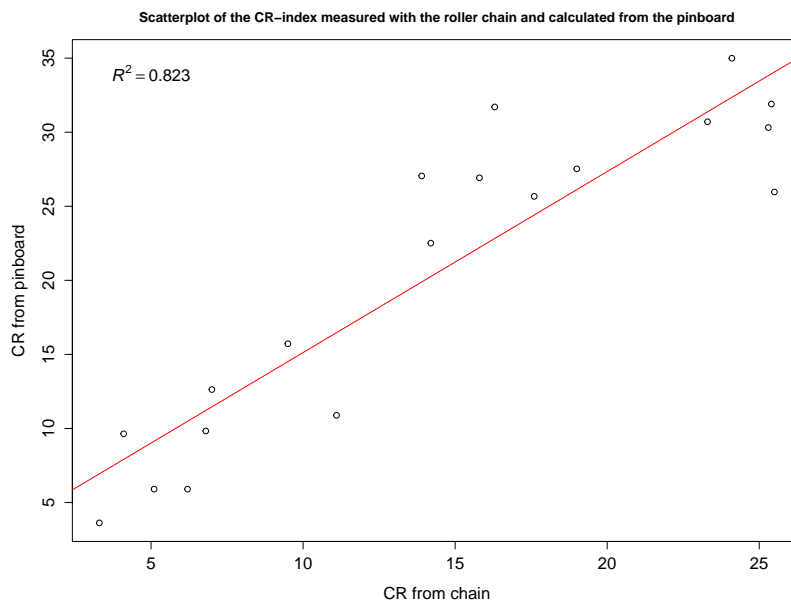
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**Figure A1.** Linear regression between the Cr index measured with the roller chain and the pinboard.

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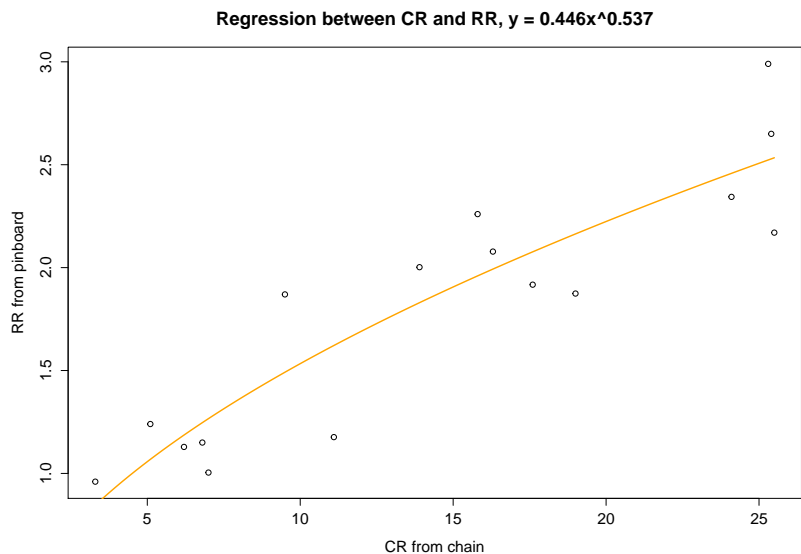
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**Figure A2.** Non-linear regression used to convert the Cr index to RR values.