Litter decomposition rate and soil organic matter quality in a patchwork heathland of Southern Norway

3

4 Norwegian heathland soils, although scant and shallow, are major reservoirs of carbon (C). 5 We aimed at assessing whether vegetation cover and, indirectly, its driving factor soil 6 drainage are good proxies for soil organic matter (SOM) composition and dynamics in a 7 typical heathland area of Southern Norway consisting in a patchwork of three different types 8 of vegetation, dominated by Calluna (Calluna vulgaris (L) Hull), Molinia (Molinia caerulea 9 (L) Moench), or Sphagnum (Sphagnum capillifolium (Ehrh.) Hedw.). Such vegetation covers 10 were clearly associated to microtopographic differences, which in turn dictated differences in 11 soil moisture regime, Calluna growing in the driest sites, Sphagnum in the wettest, and 12 Molinia in sites with intermediate moisture.

Litter decomposition was followed over a period of 1 year, by placing litterbags filled with biomass from each dominant species in each type of vegetation cover. The composition of the plant material and SOM were investigated by chemical methods and solid-state ¹³C nuclear magnetic resonance (NMR) spectroscopy.

Litter decomposition was faster for Molinia and Calluna, irrespective of the vegetation cover of the site where they were placed. Sphagnum litter decomposed very slowly, especially under Calluna, where the soil environment is by far more oxidising than under itself. In terms of SOM quality, Calluna covered areas showed the greatest differences from the others, in particular a much higher contribution from lipids and aliphatic biopolymers, apparently related to biomass composition.

Our findings showed that in the studied environment litter decomposition rate and SOM composition are actually dependent on vegetation cover and/or soil drainage. On this basis, monitoring changes in the patchwork of vegetation types in boreal heathlands could be a

- 26 reliable cost-effective way to account for climate change induced modifications to SOM and
- 27 its potential to last.

29	Litter decomposition rate and soil organic matter quality in a patchwork heathland of
30	Southern Norway
31	
32	Giacomo Certini ^{1*} , Live S. Vestgarden ^{2,3} , Claudia Forte ⁴ , Line Tau Strand ²
33	
34	¹ Dipartimento di Scienze delle Produzioni Agroalimentari e dell'Ambiente (DISPAA),
35	Università degli Studi di Firenze, Firenze, Italy
36	² Department of Environmental Sciences, Norwegian University of Life Sciences, Ås, Norway
37	³ Department of Environmental and Health Studies, Telemark University College, Bø, Norway
38	⁴ Istituto di Chimica dei Composti OrganoMetallici (ICCOM), UOS Pisa, CNR, Pisa, Italy
39	
40	* Corresponding author: certini@unifi.it
41	

43 **1. Introduction**

Heathland vegetation covers approximately 60% of Norway's land area. Norwegian 44 45 heathland soils, although scant and shallow, are so rich in organic matter that they represent a stock of carbon (C) at least one order of magnitude larger than the aboveground vegetation 46 47 they sustain (Rosberg et al., 1981). To predict the ecological effects of climate and land use changes, it is essential to understand the nature and environmental dependencies of soil 48 49 organic matter (SOM) in these widespread systems. In fact, any change influencing their 50 SOM stocks and dynamics may have major consequences for both C balance and the water 51 quality of lakes and rivers (Stuanes et al., 2008).

52 Following changes in SOM stocks is not a simple task, and several approaches have been 53 proposed for this purpose (e.g., Johnson and Curtis, 2001; Trumbore, 2009; Chiti et al., 54 2011). The current common belief is that environmental and biological factors predominate 55 on the molecular structure in controlling SOM stability (Kleber, 2010; Kleber et al., 2011; 56 Schmidt et al., 2011); however, in some environments, vegetation cover is a good proxy for 57 soil C dynamics, since it controls the input and quality of litter (De Deyn et al., 2008). In 58 turn, vegetation depends, among other factors, on soil drainage, which also influences litter 59 decay and SOM transformation (Wickland et al., 2010), so representing another possible 60 proxy for SOM storage.

Although present-day vegetation may be different from the one the underlying SOM originated from (Chambers et al., 1999; Hjelle et al., 2010), many studies have demonstrated that the most active part of SOM is the youngest (*e.g.*, Leavitt et al., 1996; Trumbore, 2000; Chiti et al., 2009). Trumbore (2000) found that the average age of the carbon dioxide (CO₂) released by decomposition processes in boreal forest soils is 30 years, and 50-60% of total soil respiration arises from SOM with mean residence time less than 1 year. The dominant 67 contribution of recently synthesized organic matter to soil respiration was also assessed by 68 Certini et al. (2003) for forests in temperate regions. Theoretically, the moister and colder the 69 pedoclimate, the better preserved the dead biomass in soil (Hobbie et al., 2000; Hicks Pries et 70 al., 2013). Hence, the wet boreal heathlands are environments where the investigation of a 71 possible relationship between vegetation covers and SOM dynamics is particularly 72 meaningful. Here, due to the intense leaching, lost dissolved organic C (DOC) may be much 73 older than the respired C (Karltun et al., 2005), rendering any possible relationship between 74 present day vegetation and bulk SOM quality less clear. Nonetheless, in the uppermost soil, 75 where SOM is younger and less degraded than below, such relationship is expected to be 76 strong enough.

77 In Southern Norway, heathland areas are in most cases characterised by the alternate 78 occurrence – essentially dictated by the soil drainage, in turn controlled by topography, 79 particle size distribution, and soil depth to bedrock – of three vegetation types, which are 80 dominated by the heather Calluna (Calluna vulgaris (L) Hull), the moor grass Molinia 81 (Molinia caerulea (L) Moench), and the peat moss Sphagnum (Sphagnum capillifolium 82 (Ehrh.) Hedw.). Such different vegetation types are cause and effect of the properties and 83 behaviour of the underlying soil. This is undoubtedly true for the soil profile morphology and 84 the sequence of horizons, generally ranging from the O-E-Bhs soil sequum of Callunasustaining podzols to multiple H horizons consisting histosols where Sphagnum grows 85 86 (Strand et al., 2008).

In this study we report an *in situ* investigation of the relationships between vegetation cover, initial litter decay rate and soil organic matter composition for a typical montane heathland area in Southern Norway where the alternation between Calluna, Molinia, and Sphagnum occurs on decametric scale. The objective of the study was to assess whether in this environment the current vegetation cover is a good proxy for SOM quality and dynamics. To 92 this end, litter decomposition was followed over a period of 1 year, by placing litterbags 93 filled with biomass from each dominant species under each type of vegetation cover, so as to 94 simulate the effects of possible climate change induced shift of vegetation on early stages of 95 litter decomposition. Furthermore, the composition of the aboveground biomass and the bulk 96 SOM were investigated by chemical methods and solid-state ¹³C nuclear magnetic resonance 97 (NMR) spectroscopy.

98

99 2. Materials and Methods

100 *2.1. Study site*

101 The study area, Storgama (59°02'47"N, 8°39'37"E), is located in the Telemark county, 102 southern Norway, at an elevation of 560-m above sea level. The mean annual precipitation in 103 Storgama for the period 1961-1990 was 994 mm, the mean annual air temperature for the 104 same period was 5.0 °C. Approximately 30 % of the area is barren granite bedrock and 105 boulders, and soil often occurs as pockets in depressions in the bedrock surface (Fig. 1a). The 106 average soil depth generally varies between 10 and 35 cm but greater thicknesses, up to 100 107 cm, do occur. According to the U.S. Soil Taxonomy (Soil Survey Staff, 2010) and moving 108 from drier to wetter locations, soils are Lithic Haplorthods, Lithic Udipsamments, Lithic 109 Endoaquents, and Lithic Haplosaprists. Although there are some scattered or vaguely 110 grouped Scots pines (*Pinus sylvestris* L) and Downy birch trees (*Betula pubescens* Ehrh), the 111 vegetation is largely dominated by heather (*Calluna vulgaris* (L) Hull) at well drained sites, 112 peat moss (Sphagnum capillifolium (Ehrh.) Hedw.) at poorly drained sites, and moor grass 113 (Molinia caerulea (L) Moench) at intermediately drained sites (Figs. 1a and 1b). These 114 dominant vegetation types are interspersed in the area, forming a patchwork dictated by 115 topography, which in turn is a driving factor of water supply. At the Calluna sites Calluna 116 vulgaris was virtually 100 % of the vegetation cover. At the Molinia sites some Calluna,

Erica (*Erica tetralix* L), and Nartecium (*Narthecium ossifragum* (L) Huds) were associated with *Molinia caerulea* but, on a visual basis, amounting to no more than 5 % of the total cover. At the Sphagnum sites, *Sphagnum capillifolium* covered the entire surface except for a few scattered individuals of Molinia, Erica and Calluna. Hereafter, we will refer to such vegetation assemblages simply as Calluna, Molinia and Sphagnum, respectively. Further pictures and information on vegetation and soils at Storgama are reported in Strand et al. (2008).

124

125 *2.2. Vegetation sampling and analysis*

126 Three sampling sites per dominant vegetation were chosen within an area of approximately 127 one hectare. At each location, three soil pits were dug down to bedrock, which was 35 to 50 128 cm deep. All the vegetation above the pit had been previously sampled and divided according 129 to species. In the case of Calluna, the woody stems and branches were separated from the 130 leaves and flowers. Capitula and the five upper centimetres were used to represent the whole 131 Sphagnum material. Visible roots were picked out from the soil samples and separated 132 according to species when possible. The aboveground biomass and the roots were analysed 133 for C and N by dry combustion on oven-dried (60 °C to constant weight) and finely ground samples using a LECO[®] CHN1000 Analyser. The aboveground biomass also underwent 134 NMR investigation. 135

- 136
- 137 *2.3. Soil sampling and analysis*

We focused our attention on the uppermost soil layer, where we expected the closest relationship between SOM quality and current vegetation. Two undisturbed soil samples, to be used for soil solution extraction, were taken by completely inserting 7.0 cm high and 4.6 cm in inner diameter, rigid cylinders at about five cm depth in each soil profile. The filled cylinders were carefully extracted from the soil and placed in a cooling box after sealing the ends with plastic lids. The samples were stored at 4 °C, for a maximum of one week, until they were processed further. Two standard disturbed soil samples were taken near the holes left by the cylinders and used for C, N, and pH determination, and NMR analysis. As for the plant material, soil C and N concentrations were measured by dry combustion on oven-dried and ground samples, while soil pH was determined potentiometrically in a 1:2.5 V/V distilled water suspension.

149 The soil-containing cylinders were inserted in two-compartment buckets and centrifuged at 150 4620 g for 20 min, following the method described by Giesler et al. (1996). The obtained 151 solution was filtered through a 0.45 µm membrane filter (Millipore). An aliquot of the filtrate 152 was analysed for total C (Shimadzu TOC-V element analyser) and, after oxidation by 153 peroxodisulphate (NS4743 1975), for total N (FiaSTAR, Tecator Spectrophotometer system). 154 Another aliquot of the filtrate was used to measure hydrophobicity, by determining the ratio 155 between the absorbances of the solution at 285 and 254 nm using an UV-VIS 156 spectrophotometer (UV-1201 Shimadzu). These two absorbances are, in fact, correlated to 157 hydrophobic C (π - π * electron transitions occur at ~285 nm for a number of aromatic substances, as described in Chin et al. 1994) and total C (Brandstetter et al., 1996), 158 159 respectively.

After centrifugation the soil was immediately passed through a 2 mm-mesh sieve. Two grams of the moist sieved soil was treated as in the second step of the procedure proposed by Ghani et al. (2003) to obtain hot-water extract (80 °C for 16 h). After centrifugation for 20 min at 2,000 g and filtration through 0.45 μ m membrane filters (Millipore), the extract was analysed for total C (HWC), total N (HWN), and carbohydrate C (Carb-C). HWC and HWN were determined by the same method as DOC and total dissolved nitrogen (TDN), while the analysis of Carb-C was done according to the "direct determination" method proposed by 167 Safarík and Santrucková (1992). In brief, 1 mL of the extract was combined in a polyethylene 168 tube with 1 mL 5% phenol solution and 5 mL concentrated sulphuric acid and immediately 169 shaken on a vortex mixer. The absorbance of the mixture was read after 1 h at 485 nm on a 170 UV-VIS spectrophotometer (UV-1201 Shimadzu). A calibration curve was built with the 171 following standards: 0.00, 0.05, 0.10, 0.25, 0.40 g L⁻¹ of α -D glucose (R²=0.9907).

172

173 2.5. Nuclear magnetic resonance spectroscopy

174 The chemical structure of the aboveground vegetation (one composite sample per dominant 175 species, after removal of stems and coarse branches in the case of Calluna) and SOM (one 176 composite sample per soil pit; hence, three samples per vegetation type) was investigated by solid-state ¹³C nuclear magnetic resonance (NMR) spectroscopy using the cross polarization 177 178 with magic angle spinning (CP MAS) technique. Prior to analysis, soil samples underwent 2 179 % HF treatment according to Skjemstad et al. (1994) in order to remove possible 180 paramagnetic oxides, which cause broadened resonances and signal loss. NMR spectra were 181 obtained by a Bruker AMX 300-WB spectrometer equipped with a 4 mm CP MAS probe. The operating frequencies were 300.13 and 75.47 MHz for ¹H and ¹³C, respectively; the $\pi/2$ 182 pulse was 3.4 µs on the ¹H channel. A contact time of 2 ms and a relaxation delay of 4 s were 183 184 used. The MAS speed was set to 8 kHz and the number of scans recorded ranged between 185 4800 and 40000, depending on the sample. The chemical shifts were referenced to 186 tetramethylsilane (TMS) using adamantane as external standard. Seven chemical-shift regions of the NMR-spectra, corresponding to the main C forms, were integrated and expressed as 187 188 percent contribution to total area subtended by the spectrum between 0 and 220 ppm. The 189 seven regions account for alkyl C (0-45 ppm, mainly comprising lipids, waxes, resins, 190 suberin), methoxyl and N-alkyl C (45-60 ppm, comprising the methoxy group of guaiacyl 191 and the two methoxy groups of syringyl lignin moieties at ~56 ppm), O-alkyl C (60-90 ppm,

192 carbohydrates, mainly cellulose and hemicellulose, with contributions from carbohydrate 193 carbons bonded to one oxygen), di-O-alkyl C (90-110 ppm, mainly from polysaccharides, 194 with contributions from anomeric carbons of carbohydrates, *i.e.*, bonded to two oxygens), H-195 and C-substituted aromatic C (110-140 ppm), O-substituted aromatic C (140-162 ppm, 196 mainly from lignin structures, tannins, polyphenols), and carboxyl C (162-190 ppm, esters, 197 acids and amides); no carbonyl intensity in the 190-220 ppm region, ascribable to aldehydes 198 and ketones, was detected.

199

200 2.6. Litter decomposition

201 Litter decomposition was determined in situ by the litterbag technique. Recently formed 202 aboveground biomass of Calluna, Molinia, and Sphagnum (approximately, the top 5 cm), 203 were collected at the end of the growing season in late September. This material was oven 204 dried (35 °C to constant weight) and used for filling 10 x 12 cm nylon mesh bags (0.5-1 mm mesh), with 3.0 g Calluna, 2.0 g Molinia, or 1.0 g Sphagnum. In November, 32 litterbags of 205 206 each vegetation type were installed on the surface of each sampling site, except Calluna 207 under Sphagnum, since a substitution of Calluna by Sphagnum was judged to be highly 208 improbable. Eight to ten litterbags per type of content were sampled from each site after 6, 9 209 and 12 months of decomposition. The removed litterbags were cleaned of plant remnants and 210 other minor foreign material by gently using a soft brush, oven dried (35 °C to constant 211 weight) and weighed for determining mass loss. Their content was thus ground and analysed 212 for carbon and nitrogen as described for the vegetation and soil samples.

213

214 2.7. Statistics

All statistical analyses were performed using the software program SAS (SAS Institute, Inc.,
1990, Cary, NC). After checking the dataset for normality and variance heterogeneity, effects

of vegetation and decomposition site on soil pH and SOM were tested by one-way analysis of variance (General Linear Model, GLM). Two separate one-way ANOVAs were performed on litterbags data: one assessing differences in decomposition rate between the three litter types disregarding the dominant vegetation effect, the other assessing differences between the vegetation types disregarding the litter effect. Pairwise comparisons were done by the Tukey's Simultaneous test.

223

3. Results

225 The experimental data set indicated marked differences in composition among the dominant 226 plant species. The C concentration in the aboveground biomass increased in the order 227 Sphagnum<Molinia<Calluna, whereas the C/N ratio increased the in order 228 Molinia<Sphagnum<Calluna, with most interspecific differences being significant (Table 1). 229 Belowground, Calluna, and Molinia also showed different composition, hence reflecting the 230 aboveground biomass (Table 1).

Concerning the soil, the measured pH values, all much below neutrality (Table 2), ensured that all C there present was in organic forms. The N content of Sphagnum was significantly higher than that of Calluna, although there were no differences in terms of C/N ratio. The latter, clustering around 20, was in all cases much smaller than the C/N ratio of the respective dominating plant species.

There was a large variability in soil DOC and TDN concentrations, and vegetation types did not show any significant difference with respect to these two variables (Table 2). On the contrary, the hydrophobicity index was significantly different in soils under the three types of vegetation, being highest for Calluna and lowest for Molinia. This difference indicates that a greater proportion of DOC under Calluna was hydrophobic. For the rest, the only other significant differences were in terms of HWN and HW-C/N ratio between Calluna, on the one hand, and Molinia and Sphagnum, on the other (Table 2).

243 The ¹³C CPMAS NMR spectra of the aboveground biomass and soil are shown in Fig. 2, and 244 the relative contributions of the different chemical shift regions are reported in Table 3. The 245 NMR spectra of the aboveground vegetation suggested more similar compositions for 246 Molinia and Sphagnum with respect to Calluna. The spectrum of the Calluna biomass was 247 dominated by signals between 60 and 104 ppm, characteristic of polysaccharides; the 248 relatively high intensity in the alkyl C region (0-50 ppm) was due to lipids and aliphatic 249 biopolymers. The spectrum also revealed the presence of lignin and tannins, as indicated by 250 the lignin methoxyl carbon signal at 56 ppm, and the distinct aromatic peaks at 145 and 155 251 ppm, typical of condensed tannins. The sharp peak at 172 ppm is normally assigned to the 252 carboxyl C of hemicellulose esters, but may also have contributions from amides (Forte et al., 253 2006). The spectra of Molinia and Sphagnum aboveground biomasses showed the same 254 dominant polysaccharide features of Calluna in the 50-110 ppm range, but a significantly 255 lower intensity of signal in the alkyl and aromatic C regions, which means lower contribution 256 of lipids and lignin/tannins, respectively. In the case of Molinia, the slightly narrower signals 257 in the 60-100 ppm region and the relatively smaller peak shoulder at about 103 ppm 258 compared with both Calluna and Sphagnum, suggested the occurrence of less hemicellulose 259 and some crystalline cellulose, respectively. Sphagnum did not show the typical lignin 260 signals, in agreement with the common lignin-free composition of bryophytes (Klavina et al., 261 2012). The only aromatic signals in the sphagnum spectrum were due to unsubstituted or C-262 substituted aryl C at 130 and 117 ppm, while the signal at 158 ppm was ascribable to 263 phenolic structures. In the case of Calluna, the two sharp tannin peaks at 145 and 155 ppm 264 observed in the aboveground biomass spectrum were totally absent in the SOM spectrum. In 265 the case of Molinia, differently from the other two vegetation types, the relative contribution 266 of aromatic C significantly increased in soil compared to that observed in the aboveground biomass. In the case of Sphagnum, no major changes occurred in the aromatic region, except
for the absence in the soil spectrum of the signal at 158 ppm detected for the aboveground
vegetation (Fig. 2). The alkyl C/O-alkyl C ratio increased for all vegetation types on passing
from the intact biomass to its decomposition products in soil (Table 3), with large differences
in absolute values between Calluna, on the one side, and Molinia and Sphagnum, on the other
side.

The *in situ* decomposition study using litterbags showed that the litter mass remaining after one year of decomposition varied between 62 and 66% in the case of Molinia and Calluna and 83 and 94% for Sphagnum (Fig. 3). The discrepancy between the couple Calluna-Molinia and Sphagnum was lower, although significant, in the intermediate stages of the experiment. After six months, Calluna showed significantly lower mass loss than Molinia under itself, while at the end of the experiment Calluna resulted to be better preserved than Molinia only under Molinia (Fig. 3).

280 In terms of relative C content of the residual litter, Calluna did not change throughout the 12 281 months of the experiment, while Molinia and Sphagnum experienced a marked decrease 282 compared to the original value (Fig. 4). Relative concentrations of N in the litter changed 283 more than the C ones. Except for Sphagnum under itself or under Molinia, all litters increased 284 their N content from November to May; later, all of them increased until August, with the 285 exception of Sphagnum under Calluna and Molina under itself; finally, in the period from 286 August to November, N concentration continued to increase in Calluna, whereas it decreased 287 in Molinia and showed an irregular trend in Sphagnum (Fig. 4). These C and N trends 288 implied progressive, although slight, decrease in C/N ratio for Calluna and Sphagnum, and a 289 sharper decrease for the same ratio for Molinia until August, after which it increased (Fig. 4). 290 Contrary to Calluna and Sphagnum, Molinia degraded maintaining significantly higher 291 values of C/N ratio under Sphagnum than under the other types of vegetation (Molinia and Calluna). At the end of the experiment, in November, the C/N ratio in Molinia underSphagnum was even higher than the original value.

294

295 **4. Discussion**

In the heathland environment of Storgama, the composition of SOM appeared to partly reflect that of the parent vegetation. Hence, for example, the abundance in alkyl C in the Calluna biomass relative to the other two vegetation types was transferred to the SOM. Nevertheless, SOM accumulated over a long period of time; as a consequence, it could be the result of multiple changes in vegetation cover in the area and, thus, be partly unrelated to the current vegetation cover. Actually, there was no direct or indirect evidence in this regard, but inputs of wind-blown or water-transported material cannot be excluded at any site.

303 Sphagnum showed a composition potentially more prone to decay than Calluna and Molinia. 304 Nevertheless, there were no significant differences in the SOM content of the topsoil of the 305 three vegetation covers. Evidently, the prevailing anoxic conditions limited decomposition at 306 the Sphagnum sites. This is in accordance with several studies that used the type of 307 vegetation cover as a proxy for carbon dynamics, based on the consideration that vegetation 308 chiefly reflects the soil moisture regime (Bridgham et al., 2008; Couwenberg et al., 2011; 309 Delarue et al., 2011), which is in turn a driving factor of litter decomposition (Hobbie et al., 310 2000; Laiho 2006). Large variability in DOC concentrations and no significant effect of 311 vegetation was observed (Table 2). It must be noted, however, that our study shows the 312 conditions only at one sampling occasion, *i.e.* at the end of the growing season, when DOC 313 concentrations are affected by a considerable contribution from senescing plant material. The 314 measured DOC concentrations were generally in agreement with those recorded in autumn 315 using zero tension lysimeters in soils at Storgama and other Norwegian heathland areas 316 (Strand et al., 2002; Vestgarden et al., 2010), although DOC concentrations in centrifuged and freely drained soil solutions are not directly comparable (Giesler et al., 1996). Similarly
to DOC, TDN showed a large variability and no apparent correlation with vegetation. The
relatively small amount of water extracted by centrifugation limited the number of possible
analyses, preventing N speciation. TDN therefore included both organic N and inorganic N,
the latter amounting to 25-50% of TDN in soil water from southern Norway (Austnes et al.,
2008; Kaste et al., 2008).

The hydrophobicity index of soil water differed significantly among vegetation types. Apparently, Calluna released DOC with the highest proportion of hydrophobic organic compounds, perhaps mostly arising from tannins and decomposition of lignin (Dilling and Kaiser, 2002), which are indeed important components of the Calluna litter (Fig. 2).

327 Hot water C approximately amounted to 4.5% of SOC in all samples, irrespective of 328 vegetation. This percentage is in the range reported by von Lützow et al. (2007). Significantly 329 lower amounts of HWN were extracted from the Calluna soils compared to the Molinia and 330 Sphagnum ones, which also implied significantly higher HWC / HWN-ratio for Calluna 331 (Table 2). We did not partition HWN, however Curtin et al. (2006) demonstrated that it is 332 mainly organic and, in suborder, NH₄-N generated by hydrolysis of heat-labile organic N. 333 The quality of the hot water extract rather well discriminated Calluna from Molinia and 334 Sphagnum. Some authors have proposed hot water extraction of SOM as a method to 335 measure the labile SOM pool (Chodak et al., 2003; Ghani et al., 2003; Curtin et al., 2006); 336 however, other authors consider this method not selective enough for this purpose (Landgraf 337 et al., 2006; von Lützow et al., 2007). In our case, approximately half the C extracted by hot 338 water belonged to carbohydrates.

The NMR spectra showed clear structural differences in aboveground plant material (Fig. 2 and Table 3). In particular, Calluna was richer in alkyl C and poorer in O-alkyl C than Molinia and Sphagnum, as well reflected in the alkyl C / O-alkyl C ratio. The richness in 342 alkyl C has been correlated to slow decomposition rates in heathland ecosystems (van Vuuren 343 and van der Eerden, 1992; van Vuuren and Berendse, 1993). However, in our litterbags 344 experiment there were little and variable differences between the mass losses of Calluna and 345 Molinia, and both of them were much higher than the one in Sphagnum wherever the latter 346 was placed (Fig. 3). A possible explanation for such short-term resistance of Sphagnum to 347 degradation could be that this type of vegetation is particularly rich in sphagnan pectin-like polysaccharides, which, unlike the other types of polysaccharides, induce processes that 348 349 prevent organic matter decay (Hájek et al., 2011; Ballance et al., 2012). Moreover, it must be 350 considered that Sphagnum might have experienced a "non-additive" pattern of mass loss, *i.e.* 351 a decomposition behaviour sometimes observed in litter mixes that deviates from the 352 response predicted for the individual species because of the influence of the other species 353 present in the mix (Gartner and Cardon, 2004). In this case, the necromasses of Molinia and, 354 in particular, Calluna could have partly inhibited the decomposition of the Sphagnum in the 355 litterbags.

356 In addition to a "vegetation effect", the litterbags experiment showed some "site effect", *i.e.* more rapid decomposition when litter was placed beneath the parent vegetation rather than 357 358 beneath other species (Ayres et al., 2009; Perez et al., 2013; Wang et al., 2013). In fact, for 359 Sphagnum the mass loss was significantly lower when it decayed under Calluna than under 360 Molina or Sphagnum (Fig. 3). Calluna was better preserved under Molinia than under itself at 361 the end of the trial, while, after six months only, Molinia litter showed significant 362 environment-induced advantage under Sphagnum compared to under itself (Fig. 3). 363 Unexpectedly, the well drained Calluna soils preserved Sphagnum and Molinia from decay 364 better than the moister soils where they were growing, perhaps as an effect of a seasonal 365 drought.

366 The enclosure of litter inside mesh bags may actually change its overall decomposition rate 367 and its C and N percent concentrations compared to non-bagged substrate (Berhe, 2013). 368 However, we must confidently assume that in our experiment such a bagging effect is 369 uniform through the samples, also because the bags are very similar. On this basis, Molinia showed an initial C/N ratio much higher than the ones of Calluna and, especially, Sphagnum 370 371 (Fig. 4), which suggested a more marked intrinsic resistance of Molinia to decay. Noteworthy 372 is the difference in C/N ratio between the aboveground Molinia biomass analysed for basic 373 characterisation (data of Table 1) and the Molinia used in the litterbags experiment (30 vs. 374 circa 80). Actually, Molinia is a grass that wilts at the end of the growing season, when we 375 sampled the material to be inserted in the bags, while the Molinia sampled for basic 376 characterisation was still with active photosynthesis, when the C/N ratio is relatively low 377 (Taylor et al., 2001). On the other hand, Calluna is an evergreen and no great seasonal 378 changes in C and N concentrations occur, while Sphagnum, although it is not an evergreen, 379 does not wilt and its C/N ratio is rather constant throughout the year. Our litterbags 380 experiment showed that the C/N ratio is a poor predictor of decay in this environment. The 381 anoxic conditions imposed by prolonged water saturation, commonly occurring in the 382 Sphagnum soils and expected to have considerable influence in slackening litter 383 decomposition, on the contrary appeared to be irrelevant in preserving organic residues 384 during a 1-year long experiment (Fig. 3). In this regard, during a 3-year study in heathlands 385 on Molinia caerulea and Erica tetralix, van Vuuren and Berendse (1993) did not find any site 386 effect and litter quality appeared to be the sole driving factor. Also Scheffer et al. (2001), 387 studying the decomposition process in fens dominated by Sphagnum species or without 388 Sphagnum, concluded that decomposition was controlled more by intrinsic differences in 389 litter quality than by the environment.

The NMR investigation revealed that soils, compared to the litter they receive, showed significantly higher contribution of alkyl C and lower contribution of O-alkyl C (Fig. 2 and Table 3), most probably as a result of a faster decay of carbohydrates than of other C forms and the synthesis of lipids from the biodegradation of carbohydrate and aromatic fractions (Baldock et al., 1992). The alkyl C / O-alkyl C ratio, which generally increases as decomposition proceeds, was significantly higher in the Calluna soil than under Molinia and Sphagnum.

- 397
- 398

399 **5.** Conclusions

400 We found that in the varied heathland of Storgama there were many significant differences in 401 terms of SOM composition between the Calluna dominated areas and the interspersed 402 Sphagnum-covered areas. Most differences were clearly due to the litter quality. A 403 "vegetation effect" on the early stage of litter decomposition rate was clear, Sphagnum 404 remnants being much more stable independently of the environmental conditions they 405 underwent, which differed especially in terms of soil drainage. Hence, overall, vegetation 406 appeared to be a good proxy for SOM quality. On this basis, monitoring the distribution of 407 vegetation types in heathlands of Norway and elsewhere could be of particular interest for 408 assessing the consequences of climate change on SOM stocks and dynamics. In the plausible 409 scenario of a less continuous rainfall supply and a consequent contraction of Sphagnum-410 covered areas, the Sphagnum-released litter seems to have good short-term ability to resist 411 decomposition under the two replacing types of vegetation, Molinia and Calluna. Long-term 412 experiments addressing this issue are needed.

- 413
- 414

415 Acknowledgements

We thank Irene Eriksen Dahl, Grete Bloch, and Ivan Digernes for laboratory assistance at the
Department for Plants and Environmental Sciences, Norwegian University of Life Sciences.
We also thank Silvia Pizzanelli of ICCOM-CNR, for performing part of the NMR analyses.
The study was carried out in close cooperation with the CLUE project (NFR 155826/S30).
This specific investigation was made possible by a grant from the Research Council of
Norway (NFR 164903/S30) enabling the first author to cooperate with researchers from the
Norwegian University of Life Sciences.

423

424 References

- 425 Austnes, K., Kaste, O., Vestgarden, L. S., and Mulder, J.: Manipulation of snow in small
- headwater catchments at Storgama, Norway: Effects on leaching of total organic carbon
 and total organic nitrogen. Ambio 37: 38–47, 2008.
- 428 Ayres, E., Steltzer, H., Simmons, B. L., Simpson, R. T., Steinweg, J. M., Wallenstein, M. D.,
- Mellor, N., Parton, W. J., Moore J. C., and Wall, D.H.: Home-field advantage accelerates
 leaf litter decomposition in forests. Soil Biol. Biochem. 41: 606–610, 2009.
- 431 Baldock, J. A., Oades, J. M., Waters, A. G., Peng, X., Vassallo, A. M., and Wilson, M. A.:
- 432 Aspects of the chemical structure of soil organic materials as revealed by solid-state ${}^{13}C$
- 433 NMR spectroscopy. Biogeochemistry 16: 1–42, 1992.
- 434 Ballance, S., Kristiansen, K. A., Skogaker, N. T., Tvedt, K. E., and Christensen, B. E.: The
- 435 localisation of pectin in Sphagnum moss leaves and its role in preservation. Carbohyd.
- 436 Polym. 87: 1326–1332, 2012.
- 437 Berhe, A. A.: Effect of litterbags on rate of organic substrate decomposition along soil depth
- 438 and geomorphic gradients. J. Soils Sediments 13: 629–640, 2013.

- Brandstetter, A., Sletten, R. S., Mentler, A., and Wenzel, W. W.: Estimating dissolved
 organic carbon in natural waters by UV absorbance (254 nm). Z. Pflanz. Bodenkunde
 159: 605–607, 1996.
- 442 Bridgham, S. D., Pastor, J., Dewey, B., Weltzin, J. F., and Updegraff, K.: Rapid carbon
- response of peatlands to climate change. Ecology 89: 3041–3048, 2008.
- 444 Certini, G., Corti, G., Agnelli, A., and Sanesi, G.: Carbon dioxide efflux and concentrations
- in two soils under temperate forests. Biol. Fert. Soils 37: 39–46, 2003.
- 446 Chambers, F. M., Mauquoy, D., and Todd, P. A.: Recent rise to dominance of Molinia
- 447 *caerulea* in environmentally sensitive areas: new perspectives from palaeoecological data.
- 448 J. Appl. Ecol. 36: 719–733, 1999.
- 449 Chin, Y. P., Aiken, G., and Oloughlin, E.: Molecular-weight, polydispersity, and
- 450 spectroscopic properties of aquatic humic substances. Environ. Sci. Technol. 28: 1853–
 451 1858, 1994.
- 452 Chiti, T., Certini, G., Perugini, L., Mastrolonardo, G., Papale, D., and Valentini, R.: Soil
- 453 carbon dynamics in a Mediterranean forest during the Kyoto Protocol commitment
- 454 periods. Reg. Environ. Change 11: 371–376, 2011.
- 455 Chiti, T., Neubert, R. E. M., Janssens, I. A., Certini, G., Curiel Yuste, J., and Sirignano, C.:
- 456 Radiocarbon dating reveals different past managements of adjacent forest soils in the
- 457 Campine region, Belgium. Geoderma 149: 137–142, 2009.
- 458 Chodak, M., Khanna, P., and Beese, F.: Hot water extractable C and N in relation to
- 459 microbiological properties of soils under beech forests. Biol. Fert. Soils 39: 123–130,
 460 2003.
- 461 Couwenberg, J., Thiele, A., Tanneberger, F., Augustin, J., Barisch, S., Dubovik, D.,
- 462 Liashchynskaya, N., Michaelis, D., Minke, M., Skuratovich, A., and Joosten, H.:

463 Assessing greenhouse gas emissions from peatlands using vegetation as a proxy.

464 Hydrobiologia 674: 67–89, 2011.

- 465 Curtin, D., Wright, C. E., Beare, M. H., and McCallum F. M.: Hot water-extractable nitrogen
 466 as an indicator of soil nitrogen availability. Soil Sci. Soc. Am. J. 70: 1512–1521, 2006.
- De Deyn, G. B., Cornelissen, J. H. C., and Bardgett, R. D.: Plant functional traits and soil
 carbon sequestration in contrasting biomes. Ecol. Lett. 11: 516–531, 2008.
- 469 Delarue, F., Laggoun-Defarge, F., Disnar, J. R., Lottier, N., and Gogo, S.: Organic matter
- 470 sources and decay assessment in a Sphagnum-dominated peatland (Le Forbonnet, Jura
- 471 Mountains, France): impact of moisture conditions. Biogeochemistry 106: 39–52, 2011.
- 472 Dilling, J. and Kaiser, K.: Estimation of the hydrophobic fraction of dissolved organic matter
- in water samples using UV photometry. Water Res. 36: 5037–5044, 2002.
- Forte, C., Piazzi, A., Pizzanelli, S., and Certini, G.: CP MAS C-13 spectral editing and
 relative quantitation of a soil sample. Solid State Nucl. Mag. 30: 81–88, 2006.
- 476 Gartner, T. B. and Cardon, Z. G.: Decomposition dynamics in mixed-species leaf litter. Oikos
 477 104: 230–246, 2004.
- 478 Ghani, A., Dexter, M., and Perrott, K. W.: Hot-water extractable carbon in soils: a sensitive
- 479 measurement for determining impacts of fertilisation, grazing and cultivation. Soil Biol.
- 480 Biochem. 35: 1231–1243, 2003.
- 481 Giesler, R., Lundstrom, U. S., and Grip, H.: Comparison of soil solution chemistry
- 482 assessment using zero-tension lysimeters or centrifugation. Eur. J. Soil Sci. 47: 395–405,
 483 1996.
- 484 Hájek, T., Ballance, S., Limpens, J., Zijlstra, M., and Verhoeven, J. T. A.: Cell-wall
- 485 polysaccharides play an important role in decay resistance of Sphagnum and actively
- 486 depressed decomposition in vitro. Biogeochemistry 103: 45–57, 2011.

- 487 Hicks Pries, C. E., Schuur, E. A. G., Vogel, J. G., and Natali, S. M.: Moisture drives surface
 488 decomposition in thawing tundra. J. Geophys. Res.-Biogeo. 118: 1133–1143, 2013.
- 489 Hjelle, K. L., Halvorsen, L. S., and Overland, A.: Heathland development and relationship
- 490 between humans and environment along the coast of western Norway through time.
- 491 Quatern. Int. 220: 133–146, 2010.
- Hobbie, S. E., Schimel, J. P., Trumbore, S. E., and Randerson, J. R.: Controls over carbon
 storage and turnover in high-latitude soils. Glob. Change Biol. 6 (Suppl. 1): 196–210,
 2000.
- 495 Johnson, D. W. and Curtis, P. S.: Effects of forest management on soil C and N storage:
- 496 meta-analysis. Forest Ecol. Manag. 140: 227–238, 2001.
- Karltun, E., Harrison, A. F., Alriksson, A., Bryant, C., Garnett, M. H., and Olsson, M. T.: Old
 organic carbon in soil solution DOC after afforestation evidence from C-14 analysis.
 Geoderma 127: 188–195, 2005.
- 500 Kaste, O., Austnes, K., Vestgarden, L.S., and Wright, R. F.: Manipulation of snow in small
- 501 headwater catchments at Storgama, Norway: Effects on leaching of inorganic nitrogen.
- 502 Ambio 37: 29–37, 2008.
- 503 Kļaviņa, L., Bikovens, O., Šteinberga, I., Maksimova, V., and Eglīte, L.: Characterization of
- chemical composition of some bryophytes common in Latvia. Environ. Exp. Biol. 10: 27–
 34, 2012.
- 506 Kleber, M.: What is recalcitrant soil organic matter? Environ. Chem. 7: 320–332, 2010.
- 507 Kleber, M., Nico, P. S., Plante, A., Filley, T., Kramer, M., Swanston, C., and Sollins, P.: Old
- and stable soil organic matter is not necessarily chemically recalcitrant: implications for
- 509 modeling concepts and temperature sensitivity. Glob. Chang. Biol. 17: 1097–1107, 2011.
- 510 Laiho, R.: Decomposition in peatlands: Reconciling seemingly contrasting results on the
- 511 impacts of lowered water levels. Soil Biol. Biochem. 638: 2011–2024, 2006.

- 512 Landgraf, D., Leinweber, P., and Makeschin, F.: Cold and hot water-extractable organic
- matter as indicators of litter decomposition in forest soils. J. Plant Nutr. Soil Sci. 169: 76–
 82, 2006.
- 515 Leavitt, S. W., Follett, R. F., and Paul, E. A.: Estimation of slow- and fast-cycling soil
- organic carbon pools from 6N HCl hydrolysis. Radiocarbon 38: 231–239, 1996.
- 517 Perez, G., Aubert, M., Decaens, T., Trap, J., and Chauvat, M.: Home-Field Advantage: A
- 518 matter of interaction between litter biochemistry and decomposer biota. Soil Biol.
- 519 Biochem. 67: 245–254, 2013.
- 520 Rosberg, I., Ovstedal, D. O., Seljelid, R., Schreiner, O., and Goksoyr, J.: Estimation of
- 521 carbon flow in a Calluna Heath System. Oikos 37: 295–305, 1981.
- Safarík, I., and Santrucková, H.: Direct determination of total soil carbohydrate content. Plant
 Soil 143: 109–114, 1992.
- 524 Scheffer, R. A., van Logtestijn, R. S. P., and Verhoeven, J. T. A.: Decomposition of Carex
- and Sphagnum litter in two mesotrophic fens differing in dominant plant species. Oikos
 92: 44–54, 2001.
- Schmidt, M. W. I., Torn, M. S., Abiven, S., and 11 others: Persistence of soil organic matter
 as an ecosystem property. Nature 478: 49–56, 2011.
- 529 Skjemstad, J. O., Clarke, P., Taylor, J. A., Oades, J. M., and Newman, R. H.: The removal of
- 530 magnetic-materials from surface soils A solid-state C-13 Cp/Mas nmr-study. Aust. J.
- 531 Soil Res. 32: 1215–1229, 1994.
- 532 Soil Survey Staff: Keys to Soil Taxonomy. 11th Edition. USDA Natural Resources
- 533 Conservation Service. Washington, DC., 2010.
- 534 Strand, L. T., Abrahamsen, G., and Stuanes, A. O.: Leaching from organic matter-rich soils
- by rain of different qualities: I. Concentrations. J. Environ. Qual. 31: 547–556, 2002.

- 536 Strand, L. T., Haaland, S., Kaste, Ø., and Stuanes, A. O.: Natural variability in soil and runoff
 537 from small headwater catchments at Storgama, Norway, Ambio 37: 18–28, 2008.
- 538 Stuanes, A. O., de Wit, H. A., Hole, L. R., Kaste, O., Mulder, J., Riise, G., and Wright, R. F.:
- Effect of climate change on flux of N and C: Air-land-freshwater-marine links: Synthesis.
 Ambio 37: 2–8, 2008.
- 541 Taylor, K., Rowland, A. P., and Jones, H. E.: *Molinia caerulea* (L.) Moench. J. Ecol. 89:
 542 126–144, 2001.
- 543 Trumbore, S.: Age of soil organic matter and soil respiration: Radiocarbon constraints on
 544 belowground C dynamics. Ecol. Appl. 10: 399–411, 2000.
- 545 Trumbore, S.: Radiocarbon and soil carbon dynamics. Annu. Rev. Earth Pl. Sci. 37: 47–66,
 546 2009.
- van Vuuren, M. M. I. and Berendse, F.: Changes in soil organic-matter and net nitrogen
 mineralization in heathland soils, after removal, addition or replacement of litter from

549 *Erica-tetralix* or *Molinia-caerulea*. Biol. Fert. Soils 15: 268–274, 1993.

- van Vuuren, M. M. I. and van der Eerden, L. J.: Effects of 3 rates of atmospheric nitrogen
- deposition enriched with n-15 on litter decomposition in a heathland. Soil Biol. Biochem.
 24: 527–532, 1992.
- Verhoeven, J. T. A. and Toth, E.: Decomposition of carex and sphagnum litter in fens effect
 of litter quality and inhibition by living tissue-homogenates. Soil Biol. Biochem. 27: 271–
 275, 1995.
- 556 Vestgarden, L. S. and Austnes, K.: Effects of freeze-thaw on C and N release from soils
- below different vegetation in a montane system: a laboratory experiment. Glob. Chang.

558 Biol. 15: 876–887, 2009.

559	Vestgarden, L. S., Austnes, K., and Strand, L. T.: Vegetation control on DOC, DON and DIN
560	concentrations in soil water from a montane system, southern Norway. Boreal Environ.
561	Res. 15: 565–578, 2010.

- von Lützow, M., Kögel-Knabner, I., Ekschmitt, K., Flessa, H., Guggenberger, G., Matzner,
- 563 E., and Marschner, B.: SOM fractionation methods: Relevance to functional pools and to
- stabilization mechanisms. Soil Biol. Biochem. 39: 2183–2207, 2007.
- Wang, Q. K., Zhong, M., He, T.: Home-field advantage of litter decomposition and nitrogen
 release in forest ecosystems. Biol. Fert. Soils 49: 427–434, 2013.
- 567 Wickland, K. P., Neff, J. C., and Harden, J. W.: The role of soil drainage class in carbon
- 568 dioxide exchange and decomposition in boreal black spruce (*Picea mariana*) forest
- 569 stands. Can. J. Forest Res. 40: 2123–2134, 2010.
- 570

571 **Figures captions:**

- Figure 1. a) A general view of the study area, Storgama, showing soil occurring in pockets 573 574 and small depressions at the bedrock surface; note that close up vegetation at the bottom right is dominated by Molinia caerulea (L), the understorey of pines beyond is Calluna vulgaris 575 576 (L) Hull, the basin in the background is covered by *Sphagnum* spp. **b**) A rare coalescence of the three dominant species, Calluna vulgaris, on the left, Sphagnum spp. L, at the bottom, 577 578 and Molinia caerulea, on the right. 579 Figure 2. ¹³C CPMAS NMR spectra of the aboveground biomass of the dominant plant 580 581 species and soil. 582 Figure 3. Residual mass in the litterbags as a function of time for different combinations of 583 584 litter and vegetation cover. Cal in Cal means Calluna litter decomposing under Calluna, Cal 585 in Mol means *Calluna* litter decomposing under *Molinia*, and so on. Error bars are standard 586 deviations, while lower case letters indicate significant differences (p<0.05) between same 587 litters decomposing under different types of vegetation. The trial was one year long. 588 589 Figure 4. Carbon and nitrogen concentrations and C/N ratio in decaying biomass in the 590 litterbags as a function of time for different combinations of litter and vegetation cover. Cal 591 in Cal means Calluna litter decomposing in soil under Calluna, Cal in Mol means Calluna 592 litter decomposing under Molinia, and so on. Upper case letters indicate significant
- differences (p<0.05) between different litters, whereas lower case letters indicate significant
 differences between same litters decomposing in soils covered by different types of
 vegetation.



Fig 1 a, b





Molinia



Sphagnum

250 200 150 100 50 0 ppm

600









Fig. 4

609	Table 1. Carbon and nitrogen concentrations and C/N ratio of the dominant plant species
610	in the Storgama area. Values in parentheses are standard deviations of six independent
611	replicates. Lower case letters indicate significant differences (p<0.05), with above and
612	below ground vegetation treated separately.
(1.2	

Vegetation	С	Ν	C/N ratio	
-	g kg ⁻¹	g kg ⁻¹		
Aboveground:				
Calluna leaves and flowers	536.0 (3.2) a	9.3 (1.0) b	58.3 (5.8) b	
Calluna stems and branches	522.6 (4.8) b	3.9 (0.2) c	135.6 (6.2) a	
Molinia	487.0 (0.9) c	16.8 (4.1) a	30.3 (7.2) d	
Sphagnum	473.9 (9.3) d	11.6 (1.2) b	41.4 (5.0) 0	
Belowground:				
Calluna roots	528.6 (15.6) a	9.2 (2.6) b	60.4 (13.6) a	
Molinia roots	499.9 (5.4) b	15.6 (5.2) a	36.4 (15.5) t	

Table 2. Selected properties of the topsoil according to dominant vegetation. Values in

617 parentheses are standard deviations of six independent replicates. Lower case letters indicate

	1
618	significant differences (p<0.05).

		Calluna		Molinia		Sphagnum	
pН	_	4.2 (0.2)		4.3 (0.1)		4.3 (0.0)	
SOC	g kg ⁻¹	373.6 (140.9)		436.8 (101.4)		459.1 (7.3)	
SON	g kg ⁻¹	16.3 (7.0)	b	22.3 (3.1)	ab	25.7 (5.0)	a
C/N-ratio soil		24 (5)		20 (2)		19 (4)	
DOC	mg L ⁻¹	86.0 (49.2)		174.5 (138.3)		53.5 (47.5)	
TDN	mg L^{-1}	4.4 (3.7)		9.9 (9.9)		3.4 (3.0)	
C/N-ratio soil solution		23 (7)		27 (19)		16 (3)	
Hydrophobicity index		0.772 (0.014)	а	0.692 (0.021)	b	0.740 (0.01:	c
HWC	g kg ⁻¹	16.3 (6.5)		20.7 (8.4)		20.7 (7.8)	
HWN	g kg ⁻¹	0.68 (0.34)	а	1.43 (0.64)	b	1.54 (0.31)	b
HWC/N-ratio		26 (7)	а	15 (3)	b	13 (4)	b
HWcarbC	g kg ⁻¹	8.6 (3.9)		11.1 (5.0)		10.1 (3.8)	
HWcarb-C/HWC	%	52 (4)		52 (11)		49 (7)	

SOC = soil organic carbon; *SON* = soil organic nitrogen; *DOC* = dissolved organic carbon;

 $TDN = total \, dissolved \, nitrogen; \, Hydrophobicity \, index = hydrophobicity \, index \, of \, soil$

621 solution; HWC = carbon in the hot-water extract; HWN = nitrogen in the hot-water extract;

HWC/N-ratio = carbon to nitrogen ratio in the hot-water extract; HWcarb-C =

carbohydrate carbon in the hot-water extract; HWcarb–C/HWC = percent carbohydrate

carbon to total carbon in the hot-water extract.

Table 3. Relative intensities, expressed as per cent of total area between 0 and 190 ppm, of seven main chemical shift regions of CP MAS ¹³C NMR spectra and the ratio between the Alkyl C and the O-alkyl C related signals for aboveground vegetation and topsoil.

Vegetati on	Sample	Alkyl C (0-45 ppm)	Methoxy l and N- alkyl C (45-60 ppm)	O-alkyl C (60-90 ppm)	di-O- alkyl C (90-110 ppm)	H-, C- substitute d aromatic C (110-140 ppm)	O- substitute d aromatic C (140-162 ppm)	Carboxyl C (162-190 ppm)	Alkyl C/O- alkyl C (0-45 ppm/60- 110 ppm)
Calluna	Abovegr ound	24	6	39	11	7	7	6	0.5
	Top 10- cm soil*	42 ± 1	8 ± 1	29 ± 3	7 ± 1	8 ± 1	2 ± 1	4 ± 1	1.2 ± 0.1
Molinia	Abovegr ound	12	8	51	14	8	3	4	0.2
	Top 10- cm soil*	28 ± 5	9 ± 1	33 ± 4	8 ± 1	11 ± 1	5 ± 1	6 ± 1	0.7 ± 0.2
Sphagnu m	Abovegr ound	13	7	51	14	8	2	5	0.2
	Top 10- cm soil*	30 ± 10	7 ± 2	36 ± 10	8 ± 2	9 ± 1	3 ± 1	7 ± 2	0.7 ± 0.3

* These values are means of three independent samples±standard deviation